THE RELATIONSHIPS BETWEEN MORPHOSTRATIGRAPHY, ROCK STRATIGRAPHY, AND ASPECTS OF TILL FABRIC IN CENTRAL ILLINOIS

> Thesis for the Degree of Ph. D. MICHIGAN STATE UNIVERSITY DAVID ALAN CASTILLON 1972





This is to certify that the

thesis entitled

The Relationships Between Morphostratigraphy Rock Stratigraphy, and Aspects of Till Fabric In Central Illinois presented by

David Allan Castillon

has been accepted towards fulfillment of the requirements for

Ph.D. degree in Geography

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ABSTRACT

THE RELATIONSHIPS BETWEEN MORPHOSTRATIGRAPHY, ROCK STRATIGRAPHY, AND ASPECTS OF TILL FABRIC IN CENTRAL ILLINOIS

By

David Alan Castillon

This thesis is concerned with the relationship between selected morphostratigraphic units and aspects of rock stratigraphy including pebble orientation in central Illinois. Investigations were made to determine the nature of the areal relationships between rock stratigraphy and morphostratigraphy of the Woodfordian end moraines in central Illinois. Included in these investigations was an analysis of long-axis pebble orientation of the till fabric. The nature of the sedimentary contact zone deposited by two lobes of Woodfordian glacial ice was studied to determine if morphostratigraphic units in an interlobate area could be delineated by the use of pebble orientation data or rock stratigraphy. Analysis of the pebble lithology of elongate stones was made to determine if a discernible difference in lithologic composition exists between delineated morphostratigraphic units or between delineated rock-stratigraphic units.

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David Alan Castillon

Thirty-eight analysis sites within the till of the Bloomington, Champaign, Eureka, and Normal Morphostratigraphic Units were selected. At each of these sites a random fifty-pebble sample was examined by established methods of till-fabric analysis. The pebbles' long-axis orientation was measured to the nearest five degrees and the dip strength was recorded. These same pebbles were classified lithologically in the laboratory. A sample of the till matrix surrounding the fifty pebbles was collected and analyzed by the laboratory of the Illinois State Geological Survey providing data on grain size, clay mineral composition, and carbonate content. Data on the color of the till matrix were provided by the author. These data were used to identify and correlate rock stratigraphic units.

The findings of this study indicate that the Eureka and Normal Morphostratigraphic Units should be mapped as a single unit identified as the Eureka-Normal Moraine (Drift). The physical characteristics of the till within the Eureka-Normal Moraine (Drift) indicate that this morphostratigraphic unit has a single rock-stratigraphic unit surface till which can be correlated with the Malden Till of the Wedron Formation. Pebble-orientation data from analysis sites along the trend of the Eureka-Normal Moraine (Drift) reveal that the long axis of elongate pebbles within the Malden Till prefer a perpendicular alignment to the trend of the moraine. These orientations indicate a Lake Michigan Lobe source. Pebble

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lithology data show that a lower percentage of crystalline pebbles and a larger percentage of carbonate pebbles are embedded in this till than were found in tills from the other morphostratigraphic units of this study.

The data from analysis sites within the till of the Champaign Moraine (Drift) indicate that this morphostratigraphic unit is well defined topographically and that the surface till of this unit can be correlated with the Snider Till of the Wedron Formation. This same rock-stratigraphic unit was correlated with the till on the eastern flank of the interlobate region. This portion of the interlobate region adjoins and can be associated with the Champaign Moraine (Drift) of this study. The long-axis pebble orientations of the Champaign Moraine (Drift) tend to be orthogonal to the trend of the morphostratigraphic unit outside of the interlobate, then become parallel to the trend of the landform in the associated interlobate region. An Erie Lobe source can be postulated for the Snider Till from these orientation findings. Pebble lithology data from the Champaign Moraine (Drift) and associated interlobate area indicate that pebbles within till from an Erie Lobe source have approximately 9 per cent fewer carbonate pebbles when compared with pebble lithology data from the Eureka-Normal Moraine (Drift) which has a Lake Michigan Lobe source.

The findings of this study reveal that the Bloomington Moraine (Drift) is very complex. Three distinct surface

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David Alan Castillon

tills within this morphostratigraphic unit can be delineated: the Tiskilwa, Malden, and Shamrock Till Members. The latter was named and defined in this study. As was the case with the other units investigated in this study, the long axis of pebbles embedded in these tills preferred a perpendicular alignment to the trend of the moraine outside of the interlobate area. In the interlobate region associated with the Malden Till of the Bloomington Moraine (Drift) pebble-orientation data once again reveal a strong tendency for pebbles to align themselves parallel to the trend of the landform.

Because the relationships between morphostratigraphy, rock stratigraphy, and till fabric did not complement one another within the area of the Bloomington Moraine (Drift) as they did on the other units of this study, an alternate interpretation of the morphostratigraphy was investigated. An alternate interpretation must account for three distinct rock-stratigraphic units and also account for an ice source region that will explain its fabric. An analysis of the topography of the Bloomington Moraine (Drift) indicates that it is not necessarily one morphostratigraphic unit. Instead. it may represent a situation where three units of different age converge on one another. The oldest unit identified by the presence of Tiskilwa Till, which is the surface till west of Bloomington, Illinois, and exists over a wide area in the subsurface, was deposited by an ice advance from the north. The Shamrock Till, whose fabric orientations suggest it was

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deposited by a western advance, possibly of the Erie Lobe, was deposited and overlies the Tiskilwa Till. A third and younger till unit which is the surface till within the moraine everywhere east of the Kickapoo Creek Valley and also the surface till on a part of the proximal slopes to the west of this stream was deposited contemporaneously with the Snider Till of the Champaign Moraine (Drift) forming the interlobate moraine. The time-stratigraphic correlation of Malden and Snider Tills suggested by the fabric orientations in the interlobate region, which are parallel to the ice-contact zone, and the similar physical characteristics of these two till members suggest that the landform associated with this region be named the Champaign-Bloomington Interlobate Moraine (Drift).

Of general interest to the area of glacial geomorphology is the finding of this study that morphostratigraphic units exist and can be useful geomorphically when sedimentary parameter data complement the landform unit.

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By

David Alan Castillon

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1972

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Dedicated to

Joanne Cindy, Kim, Linda, Laura, and David Alan II

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Chapter 1

INTRODUCTION

Statement of Problem

This thesis is concerned with the relationships between selected morphostratigraphic units and aspects of rock stratigraphy including pebble orientation in central Illinois. Several specific matters, listed below, were investigated in an effort to determine these relationships. (1) The nature of the relationship between rock stratigraphy and morphostratigraphy of the Woodfordian end moraines in central Illinois was investigated. (2) The nature of the relationship between the pebble orientations of the till fabric and morphostratigraphy of the Woodfordian end moraines in central Illinois was investigated. (3) Investigations of the contact zone between two glacial lobes of the Woodfordian glacial ice were made to determine if morphostratigraphic units in an interlobate area could be delineated by the use of pebble-orientation data or rock stratigraphy. And (4) investigations were made to determine if a discernible difference in pebble lithology of the collected orientated stones exists between delineated morphostratigraphic units, or between delineated rockstratigraphic units.

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¹The study area is defined on page two and illustrated in Figure 1.

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Study Area

The study area selected in central Illinois contains one of the best examples in North America of arcuate-lobate end moraines with good topographic expression that merge into what appears to be an interlobate area. The study area has an approximate east-west extent of 40 miles (73 kilometers) and north-south extent of 30 miles (48 kilometers), and includes parts of Champaign, Ford, McLean, Piatt, and Woodford Counties in Illinois. The study area is delineated in Figure 1, a map of Illinois.

The literature describing various aspects of all or part of this area is extensive. An early work, very comprehensive for its time and entitled "The Illinois Glacial Lobe," was written in 1899 by Frank Leverett. This monograph describes the Pleistocene landforms of the state of Illinois and has served since its publication as a primary reference for glacial geology studies in the state. Leverett's interpretation of the glacial landforms of the study area was modified only slightly during the first thirty years following publication. He revised some of his interpretations in 1929 and again in 1932. These revisions included minor changes in both the delineation of moraines (Leverett, 1932) and the interpretation and terminology regarding substages of the Wisconsin glacial stage (Leverett, 1929). From 1932 to the present, changes in the interpretation of the Wisconsin glacial stage have continued. Two of the more

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Figure 1 Map of Illinois illustrating study area

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important works concerned with the chronology of the Wisconsin glacial stage in Illinois were made by Leighton (1958) and Frye and Willman (1960).²

Many other publications describing various characteristics of the Wisconsinan glacial drift of Illinois were published in the 1960's. To cite a few of the more important works: Frye, Glass, and Willman (1962) published a study on the stratigraphy and mineralogy of the Wisconsinan loess; Willman, Glass, and Frye (1963, 1966) wrote on the mineralogy and weathering profiles of glacial till; and Piskin and Bergstrom (1967) described the thickness and character of the glacial drift of Illinois.

Two recent publications describing the study area were written by Willman and Frye (1970) and Kempton, DuMontelle, and Glass (1971). The former, Bulletin 94 of the Illinois State Geological Survey entitled "Pleistocene Stratigraphy of Illinois," is general in nature and covers the entire state whereas the latter is specific and focuses on till units in McLean County.

Bulletin 94 is useful because Willman and Frye review the principles of stratigraphic classification for Pleistocene deposits. Four different types of stratigraphy may be recognized: (1) time stratigraphy, (2) morphostratigraphy, (3) rock stratigraphy, and (4) soil stratigraphy. The study area for this thesis is a part of the Woodfordian time

4

²Hereafter in this study the term "Wisconsinan" will be used for Wisconsin when referring to the glacial stage as suggested by Frye and Willman (1960).
<u>zine</u>(ie De **n**:-5 *i*it (1 Ţ ssti 2012S A ut i ii nay ⊐its; : ye Hite p in:f ::::i 0.1¢.11 ≓gla Cea: Reat te te hìch stratigraphic unit, the morphostratigraphic units broadly defined by the Peoria Sublobe of the Lake Michigan Lobe and the Decatur Sublobe of the Erie Lobe, the Wedron Formation rock-stratigraphic unit, and the Jules Soil soil-stratigraphic unit (see Figure 2). This study is concerned with certain relationships between morphostratigraphy and rock stratigraphy.

The Kempton, DuMontelle, and Glass (1971) work on the stratigraphy of Woodfordian tills in McLean County is useful because it defines five till units, three of which are surface tills in the study area of this thesis.

Morphostratigraphy

A morphostratigraphic unit is defined "as a body of rock that is identified primarily from the surface form it displays; it may or may not be distinctive lithologically from contiguous units; it may or may not transgress time throughout its extent" (Frye and Willman, 1960, 1962). Morphostratigraphic units were proposed in an effort to explain the relationship and significance of topographic forms to rock, soil, and timestratigraphic units. The term "Drift" is attached to a named morphostratigraphic unit and is used to connote the deposit of glacial till and outwash associated with a moraine and traceable from it into the ground moraine, outwash apron, and beneath younger drifts (Willman and Frye, 1970). In this study the term "Moraine" will be retained to include the landform which is discernible by its topographic expression and the term "Drift" when included in parentheses following Moraine



Figure 2 Stratigraphic classification of the Pleistocene deposits of Illinois (after Willman and Frye, 1970)

25 :)] 93V ī.e xe 1.... n 3 i) έx 2017 .he Fig: æ ŧЦ iy W ÷.s iora 3 . T See is interpreted to be at least partially representative of the morphostratigraphic unit as defined by Willman and Frye (1970).

During the Woodfordian Substage two lobes of glacial ice covered approximately one-third of the northeastern sector of the state of Illinois (see Figure 3). The Erie Lobe had only one sublobe in Illinois called the Decatur Sublobe (Frye and Willman, 1970). The Lake Michigan Lobe had six sublobes (Frye and Willman, 1970). This study is concerned with some of the deposits of both the Peoria Sublobe of the Lake Michigan Lobe and the Decatur Sublobe of the Erie Lobe. These two sublobes deposited a sequence of Drifts that form the basis of the morphostratigraphy of the study area (Frye and Willman, 1970). The map entitled "Woodfordian Moraines of Illinois" (see Figure 4, in pocket) is included in this dissertation by permission of the Illinois State Geological Survey and shows all mapped Woodfordian morphostratigraphic units in the state.

The individual morphostratigraphic units defined as Drifts by Willman and Frye (1970) and selected for investigation in this study are based on part of the Bloomington and Eureka Moraines and all of the Normal Moraine of the Peoria Sublobe and part of the Champaign Moraine of the Decatur Sublobe. The drifts of this study from the Peoria Sublobe merge with the drift of the Decatur Sublobe to form the interlobate area (see Figure 4, in pocket).



Figure 3 Woodfordian lobes and sublobes in Illinois (after Willman and Frye, 1970)

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Rock Stratigraphy

The American Commission on Stratigraphic Nomenclature Code (1961, page 649) defines a rock-stratigraphic unit as follows:

A rock-stratigraphic unit is a subdivision of the rocks in the earth's crust distinguished and delimited on the basis of lithologic characteristics....Rock-stratigraphic units are recognized and defined by observable physical features rather than by inferred geologic history.

The study area for this thesis lies entirely in the rockstratigraphic unit named the Wedron Formation (see Figure 5). The Wedron Formation consists of glacial tills, outwash sands and gravels associated with the glacial tills, and silt deposited in Illinois during the Woodfordian Substage of the Wisconsinan Stage (Frye et al., 1968; Willman and Frye, 1970). The formation is bounded below by the Morton Loess or older deposits and above by the Richland Loess (Frye et al., 1968). Its type section is the Wedron Section, Wedron Silica Company pit, SE SW Sec. 9, T.34 N., R.4 E. (Willman and Frye, 1970).

Willman and Frye (1970) differentiated eight till members in the Wedron Formation in Illinois. Kempton, DuMontelle, and Glass (1971) defined five till units in the Wedron Formation in the McLean County region. And Johnson, Follmer, Gross, and Jacobs (1972) recognized four till members in the Wedron Formation in east-central Illinois. Figure 5 gives the areal distribution of the named till members as published in Guidebook Series 9 of the Illinois State Geological



Figure 5 Areal distribution of till formations and members in Illinois (after Johnson et al., 1972, and others)

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On the basis of works by Willman and Frye (1970); Kempton, DuMontelle, and Glass (1971); Johnson, Gross, and Moran (1971); and Johnson et al. (1972) the defined or correlated surface units for the study area of this thesis are (1) the Tiskilwa and Malden Till Members of Willman and Frye (1970); (2) Till Units 1, 2, and 4 of Kempton, DuMontelle, and Glass (1971); and (3) the Glenburn, Snider, and Batestown Till Members of Johnson, Gross, and Moran (1971).

According to Willman and Frye (1970) the Tiskilwa Till Member is the uppermost till of the western part of the Bloomington Moraine (Drift) in the study area. The western part of the Bloomington Moraine (Drift) is here defined as the drift west of Bloomington, Illinois. The type section for the Tiskilwa Till Member is a roadcut in Bureau County known as the Buda East Section. It is located in the SE SE SW Sec. 31, T.16 N., R.8 E., 5 miles (8 kilometers) northwest of Tiskilwa (Frye and Willman, 1965, page 96, unit 1). In the opinion of the author the Tiskilwa Till Member can be correlated areally with Unit 4 of Kempton, DuMontelle, and Glass (1971). Unit 4 is correlated with the Glenburn (Johnson, Gross, and Moran, 1971).

The Glenburn Till Member, which has a possible correlation with the Tiskilwa Till, is named for the town of Glenburn and

37 ; 20 I Ш, y) HET.C .per Te s Ċe ! x. ÷ V e;... ad I .pe 1 1 5. 17 Ē 37 HS. 0, Ŷ its type section is the Emerald Pond Section NE SW SW Sec. 33, T.20 N., R.12 W., Vermillion County (Johnson, Gross, and Moran, 1971, page 211). Radiocarbon dates indicate that this unit may be pre-Woodfordian (Johnson et al., 1972).

The Malden Till Member is the uppermost till of the eastern part of the Bloomington Moraine (Drift) and is the uppermost till of the Eureka and Normal Moraines (Drifts) in the study area (Willman and Frye, 1970). The type section for the Malden Till Member is in the Malden South Section, SW SE SE Sec. 5, T.16 N., R.10 E., roadcuts 2 miles (3 kilometers) south of Malden (Willman and Frye, 1970, pages 185-86). The Malden Till Member is differentiated into two units (Kempton, DuMontelle, and Glass, 1971). Unit 2, broadly defined, represents the uppermost till of the eastern part of the Bloomington Moraine (Drift), and Unit 1 is the uppermost till of the Eureka and Normal Moraines (Drifts) (Kempton, DuMontelle, and Glass, 1971).

The Batestown Till Member is correlated with the uppermost till of the Champaign Moraine (Drift) in the study area (Johnson et al., 1972). The type section for the Batestown Till Member is the Emerald Pond Section in Vermillion County, NE SW SW Sec. 33, T.20 N., R.12 W., (Johnson, Gross, and Moran, 1971, page 211). Unit 3 of Kempton, DuMontelle, and Glass (1971) was not shown to exist in the study area of this thesis. However, it has been suggested that the Batestown Till Member correlates with Unit 3 (Johnson, Gross, and Moran, 1971; Johnson et al., 1972). Previously the Batestown Till was called

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The Snider Till Member was not previously recognized or correlated with any of the tills in the study area of this thesis, but the data collected in this study suggests a possible correlation (see Chapter 4). In the Danville region the Snider Till is the youngest till and is separated from the Batestown Till Member below by stratified drift and is named for the town of Snider, Illinois (Johnson, Gross, and Moran, 1971). The Emerald Pond Section in NE SW SW Sec. 33, T.20 N., R.12 W., Vermillion County is the type section (Johnson, Gross, and Moran, 1971, page 211).

Unit 5 of Kempton, DuMontelle, and Glass (1971) is reported to exist in the study area only as a subsurface till and was correlated by Johnson et al. (1972) with the Oakland Till Member of Ford (in preparation).

Figure 6 summarizes the relationship between Till Members or Till Units with suggested possible correlations.

Eve land (1952)	Ekblaw Willman (1955)	Willman Frye (1970)	Kempton DuMontelle Glass (1971)	Johnson Gross Moran (1971)	Johnson et al. (1972)
			1	Snider	Snider
	Champaign	Malden	2	Batestown	Batestown
Tazewell	Cerro Gordo		3		
		Tiskilwa	4	Glenburn	Glenburn
			5		Oakland

Figure 6 Correlated till members and till units

Till Fabric

Holmes (1939) defined till fabric as the space relations among the component rock and mineral fragments in undisturbed till and used a quantitative interpretation of these space relationships in an attempt to explain the direction of glacial ice movement. The technique established by Holmes to determine fabric has not been modified significantly over the past thirty years, but refinements in methods of analysis have been described during this period of time. Examples include Harrison (1957a), Kauranne (1960), and Andrews and Shimizu (1966). Practical uses for till-fabric analysis have expanded to include determination of the speed of glacial ice movement (Harris, 1968), stratigraphic variation in tills (Rains, 1969), reorientation of till by readvancing ice (MacClintock and Dreimanis, 1964), and correlation with landforms such as drumlins (Wright, 1957) and moraines (Hoppe, 1952). Also variability of till fabrics both horizontally and vertically has been studied by West and Donner (1957) and Young (1969).

The only till-fabric studies performed near the study area of this thesis are included in papers by Harrison (1957b), Smith (1970), and Lineback (1971). The results of the studies by Smith and Lineback are summarized in Guidebook Series 9 of the Illinois State Geological Survey (Johnson et al., 1972). Orientation findings for the Glenburn and Batestown Till Members of the Wedron Formation in the Danville, Illinois region show the Glenburn Till Member to have a southwest

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The till-fabric analysis for this study utilizes the technique of Holmes (1941). Both dip and strike were measured while the pebbles were in place in the till. Pebbles were then removed, measured to the nearest millimeter, and classified according to shape and lithology. The field and laboratory procedure of the till-fabric analysis is described in Chapter 2.

Till Matrix Composition

The till matrix was analyzed for grain size, clay mineral composition, and carbonate content by the laboratory of the Illinois State Geological Survey in Urbana, Illinois. Their help and cooperation is gratefully acknowledged. The laboratory procedures used to analyze the till matrix are described in Chapter 2.

The Illinois State Geological Survey classifies gravel as particles larger than 2.0 millimeters, sand as particles between 2.0 and 0.062 millimeters, silt as particles between 0.062 and 0.004 millimeters, and clay as particles smaller than 0.004 millimeters.

The clay mineral composition is determined by X-ray diffraction of oriented aggregates for clay-size particles. Three categories--expandable clays, illite, and chlorite plus kaolinite--are separated by this technique. The X-ray diffraction process is also used to determine the carbonate

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content. Carbonates are measured in counts per second giving either calcite or dolomite depending on the angle of analysis.

The till matrix was also described according to color. Till samples were moistened and a Munsell Soil Color Number was determined, and a verbal notation of the dry color was determined.

Data and Analysis Sites

In this study till-fabric and till-matrix composition analyses were performed on samples collected at 4 to 5 mile intervals (6 to 8 kilometers) along the trend of the moraines and at approximately 2 mile (3 kilometers) intervals in the interlobate area. A discussion on analysis site selection is presented in Chapter 2.

A total of thirty-eight analysis sites were selected with thirteen of these in or near the interlobate area. For each analysis site data collected includes (1) location of the site; (2) elevation; (3) type of exposure (roadcut, pipeline trench, etc.); (4) depth of sample; (5) topography of the area including location of site with regard to the outline of the moraine (proximal, distal, or central); (6) till-pebble data which includes the size, shape, long-axis orientation, and lithology of each pebble; and (7) till-matrix composition data which includes the till color, grain-size distribution, clay mineral composition, and carbonate content. The data for all thirty-eight analysis sites are included in the Appendix of this thesis.

The relationship between landforms and sedimentary parameters can be found extensively in the literature of glacial geomorphology. For example, Hoppe (1952) carefully described the location of each of his till fabrics in terms of its location with respect to the outline of the moraine or hummock from which it was taken. Such descriptions made it possible for him to state certain relationships between sediments and landforms. Rains (1969) used the sedimentary parameters of till-fabric analysis to confirm the existence of two till units stratigraphically at a particular location. Anderson (1955) used the sedimentary parameter of pebble lithology and reports that pebble lithology is determined largely by the position and distance of the source area relative to the direction of ice movement and can be used to distinguish tills from different ice lobes. Willman and Frye (1970) used sedimentary parameters to differentiate the rock-stratigraphic units defined in Illinois.

The purpose of this thesis is to analyze and describe the relationships that exist between certain sedimentary parameters and landforms. The landform unit in this thesis is the morphostratigraphic unit (end moraine). The sedimentary parameters are pebble orientation, pebble lithology, and till-matrix composition. This thesis will analyze and describe the relationships between four morphostratigraphic units in central Illinois and certain sedimentary parameter data gathered from thirty-eight analysis sites along these units.

Justification for the Study

Till-fabric analysis and till-matrix composition analysis have proved to be important and meaningful field and laboratory techniques in glacial geomorphology. In addition to adding to knowledge of the glacial geomorphology of an area, they may reveal especially significant relationships between landforms and sediments. The writer knows of no till-fabric studies that apply the concept to interlobate areas. Such application may make it possible to determine at least in part ice behavior in interlobate areas and to better understand the nature of deglaciation in such regions. Finally, the area selected in Illinois is especially favorable for this analysis because the topography has been carefully mapped and the landforms are a part of one of the best understood Midwestern glacial landscapes.

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Chapter 2

FIELD WORK AND LABORATORY ANALYSIS

Selection of the Study Area

Introduction

The primary objective of this study is to determine relationships between morphostratigraphy and certain aspects of rock stratigraphy including pebble orientation, pebble lithology, and till-matrix composition. Such relationships would probably be most revealing (1) where the end moraines have good topographic expression both vertically and laterally, (2) where the proximal and distal boundaries of end moraines are reasonably well defined, (3) where the end moraines were formed by two distinct lobes of glacial ice, and (4) where the end moraines of two lobes merge into an interlobate area. An area that satisfies all these criteria is located in central Illinois, where the Eureka, Normal, and Bloomington Moraines (Drifts) merge with the Champaign Moraine (Drift) in what appears to be an interlobate area.

Study Area

Willman and Frye (1970) defined and delineated the Bloomington, Champaign, Eureka, and Normal Moraines (Drifts) (see Figure 4, in pocket). The study area for this thesis is shown on Figure 7 and can be delineated as follows: The Bloomington Moraine (Drift) was studied from the McLean County line on the west to the interlobate area in eastern



Figure 7 Morphostratigraphic units of the study area with general location of analysis sites

McLean County. The Champaign Moraine (Drift) was studied in the area between its two points of dissection by the Sangamon River approximately marked by analysis site locations C6 and C1. The Eureka Moraine (Drift) was studied from the area of its dissection by the Mackinaw River near analysis site EB2 to the interlobate area in eastern McLean County. And the Normal Moraine (Drift) was studied throughout its entire extent.

Selection of Analysis Sites

Introduction

The Map "Woodfordian Moraines of Illinois" (Figure 4, in pocket) summarizes more than seventy years of refinement in the mapping of the surficial landforms of the study area of this thesis. Leverett (1897, 1899) first mapped the Pleistocene landforms of this area, and the Bloomington Morainic System and Champaign Morainic System were named by him. Leighton and Ekblaw (1932) and Ekblaw (1941, 1959) revised and updated the mapping of Pleistocene deposits in the study area, adding the Normal Moraine. Further revisions by M. M. Leighton and J. A. Brophy (1961) altered the extent of the Normal Moraine. Willman and Frye (1970) defined the Eureka Moraine and established the extent of the Bloomington, Champaign, and Normal Moraines to their present positions.

Analysis Sites

Analysis sites were spaced 4 to 5 miles (6 to 8 kilometers) apart along the trend of the Bloomington and Champaign Moraines (Drifts), but because of the close and problematical

relationship between the Eureka and Normal Moraines (Drifts) the 4 to 5 mile (6 to 8 kilometers) interval was alternated. The series of alternating sites appears justified in this area because according to Willman and Frye (1970) the Eureka and Normal Moraines (Drifts) are both part of the same rockstratigraphic unit, the Malden Till Member. In the interlobate area where greater definition and refinement of relationships is sought, the interval between analysis sites was reduced to a 1 to 2 mile (1 to 3 kilometers) interval. A total of thirty-eight general analysis locations was selected along the moraines (see Figure 7).

Once a desirable general location for an analysis site was selected, two factors were considered before the specific location was determined. First only an exposure (such as a railroad cut, roadcut, pipeline trench, or borrow pit) within the moraine deep enough to reveal till below the zone of calcium carbonate (CaCO₃) leaching was considered to be a suitable location for analysis (see Figure 8). The surface till of the study area lies stratigraphically below the Richland Loess (Willman and Frye, 1970). The thickness of loess on top of the moraines of the study area ranged from less than 1 foot (0.3 meters) to approximately 8 feet (2.3 meters) in depth. Three feet (1 meter) of loess was very common. The surface of the calcareous zone was usually found to be about 4 feet (1.3 meters) and never exceeded 5 feet (1.5 meters) below the top of the cut or excavation. Figure 8 Examples of analysis sites

Figure 8A Pipeline trench Six foot deep trench in Malden Till at analysis site E7; (See Figure 7 for location)

Figure 8B Borrow pit Borrow pit exposure of Shamrock Till along I-74 at analysis site B4; (See Figure 7 for location)



Figure 8A



Figure 8B

Figure 8 Examples of analysis sites

Figure 8C Roadcut with dig Twelve foot roadcut exposure showing Richland Loess overlying Malden Till at analysis site N2

Figure 8D Close up of dig illustrating dip measurement Brunton compass measuring dip of pebble embedded in Snider Till at analysis site BD7



Figure 8C



Figure 8D

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The average depth of till analysis was 8 feet (2.3 meters) below the original surface but this figure ranged from 4 to 25 feet (1.3 to 7.5 meters). A second factor considered in the selection of an analysis site was its location relative to the surrounding topography. Glen, Donner, and West (1957) suggest that, if care is taken to select deposits that do not appear to have been seriously disturbed since deposition, reorientation by gravitational movement can be minimized. When an exposure was selected as suitable, an attempt was made to locate the dig⁵ in such a way that reorientation of the fabric by slides or mechanical movement was minimized. This was accomplished in most cases by selecting satisfactory exposures in the higher part of the moraine in the area of the analysis. By selecting such a location, the till under analysis could be assumed to have most likely remained unmoved or unaffected in any way by material that had moved as a result of gravitational mass wasting since its time of deposition. Thus, the fabric of the till would represent its original depositional environment.

Field Technique

Introduction

Most of the field data for this thesis was collected during the summer of 1971. The field procedure was initiated

A "dig" is here defined to be the excavation into the side of an exposure necessary to create a suitable working surface for a till-fabric analysis (see Figures 8C and 8D).
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by the selection of an analysis site. Each location was identified by a field note number which was to identify the analysis site. The letter B was used to identify analysis sites along the trend of the Bloomington Moraine (Drift), C was used for Champaign Moraine (Drift), E for Eureka Moraine (Drift), and N for Normal Moraine (Drift). The letters EB refer to an analysis site located areally on the Eureka Morphostratigraphic Unit but in a lower rock-stratigraphic unit than the top till of the Eureka Moraine (Drift). In the interlobate area, which was defined by Willman and Frye (1970) as Bloomington Moraine (Drift), the letters BP and BD were used to identify analysis sites. BP denotes defined Bloomington Moraine (Drift) on the Peoria Sublobe side of the interlobate, and BD denotes Decatur Sublobe side of the same unit. The boundary between Peoria Sublobe side and Decatur Sublobe side of the interlobate is defined as a line connecting the upper limits of two intermittent streams as they bisect the interlobate (see Figure 9). Along with each letter designation is a number which identifies a specific location on the morphostratigraphic unit.

The identifying number, legal land description, elevation, type of exposure, depth of sample below the top of the exposure or cut, and topography were recorded for each analysis site (see the Appendix). Topography was described by slope development in the area of the analysis site and location with respect to the outline of the moraine. Slope development was



Figure 9 Interlobate map

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given by letter designation according to the scheme of the U.S.D.A. Soil Soncervation Service Soil Surveys (see Figure 10).

A BOARDEN	month of	Descriptive Terms									
Slope	% Grade	USDA-SCS	(General)								
A	0-2	Nearly level	(Level-flat)								
В	2-6	Gently sloping	(Undulating)								
С	6-12	Moderately sloping	(Rolling)								
D	12-18	Strongly sloping	(Hilly)								
E	18-25	Steep	(Steep)								
F	25-35	Very steep	(Very steep)								
G	Over 35	Very steep	(Rough)								

Figure 10 Slope development classification

Location of the analysis site with respect to the outline of the moraine was described as proximal, distal, or central.

Till-Fabric Analysis

At each site selected a horizontal surface approximately 3 feet square (2500 square centimeters) was excavated (see Figure 8D). The surface was cleared and leveled with a mattock and brushed with a small broom to expose the top surface of any pebbles at that level. The matrix of till around each pebble was then carefully removed with a pocket knife in an effort to determine the long axis of the pebble. Approximately 4 pounds (1.8 kilograms) of the excavated till matrix was collected and bagged for later laboratory analysis as the pebbles were uncovered. While the pebble remained in place within the till matrix, the direction of the long axis was marked with a pencil line on the pebble. The Brunton compass was used to determine the direction of the pebble's

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Appendix

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long axis to the nearest 5 degrees in the 180-degree sector to the north of east-west. The clinometer measured the amount of dip of each pebble to the nearest 5 degrees (see Figure 8D). The measurement of pebbles to the nearest 5 degrees in the field was utilized because it appears to provide an acceptable degree of accuracy for till-fabric studies (Harris, 1969). Both measurements were recorded in the field notes (see the Appendix). The pebble was then removed, placed in an envelope, and filed for later laboratory analysis. After all pebbles of one layer had been measured, recorded, and filed, a second lower level was scraped clean and the process repeated until a random sample of fifty pebbles had been collected (see Figure 11). Harris (1969) illustrates statistically that a fifty-pebble sample is usually sufficient to show preferred orientation.

On occasions when the long axis of a pebble was vertical, the pebble's long axis orientation measurement (strike) is actually a measurement of the orientation of the pebble's second longest axis.⁴ Pebble orientation data is presented in the Appendix and summarized graphically on Figure 14 (in pocket).

Laboratory Analysis of Pebbles

Introduction

Each pebble collected was analyzed in the laboratory. The laboratory procedure included (1) measuring the three

⁴An example of a study that discusses pebbles with a vertical long axis orientation is Hoppe (1952).

Figure 11 Random sample of fifty pebbles The fifty-pebble sample from analysis site N1 serves as an example to illustrate size and shape of pebbles for an analysis site location of this study.

Figure 12 Representative pebble shape categories 1 and 2 -- Tabular 3 and 4 -- Rhombohedroid 5, 6, and 7 -- Wedge-form 8 and 9 -- Ovoid -



Figure 11

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Figure 12

dimension (2) class and (3) i These lab summarize For ategorie forms and of his pe woid, ta The round slightly Holmes' ((1923, 19 Wentworth to descr lake reg illustrat of form a ^{tabular} tis reas the sharp ^{shar}ply a ^{and} Varih ^{study}. F dimensions of each pebble to the nearest millimeter, (2) classifying the pebble according to shape (see Figure 12), and (3) identifying and classifying each pebble's lithology. These laboratory data are presented in the Appendix and summarized in Tables 2, 5, and 7.

Pebble Shape

For this study pebble shape was divided into nine categories (see Figure 12). Holmes (1941) used six major forms and four categories of roundness to describe the shape of his pebbles. The six major forms were called discoid, ovoid, tabular, wedge-form, rhombohedroid, and varihedroid. The roundness classifications were called sharply angular, slightly rounded, moderately rounded, and well rounded. Holmes' (1941) categories were based on the work of Wentworth (1923, 1936), Wadell (1932, 1936), and Tester (1931). Wentworth (1936) used thirteen descriptive categories of shape to describe glacial pebbles from the Baraboo and Devil's Lake region of Wisconsin. Holmes (1941, plate 1, page 1355) illustrates seventeen of his twenty-four possible categories of form and roundness. Holmes (1941, page 1307) states: "tabular and discoid stones should be grouped together." For this reason and because it is very difficult to differentiate the sharply angular to slightly rounded varihedroid from the sharply angular to slightly rounded wedge-form, the discoid and varihedroid forms of Holmes (1941) were not used in this study. Four forms--tabular, rhombohedroid, wedge-form, and

ovoid--we he first in form, stapes we four. Th Three wed six, and five or s same shap The numbe roundness and the n the round sharply a assigned ratio of elongated included · tiesis. Class ^{jeer} a res Milthers (in Norway ^{stud}ied va ovoid--were expanded into nine shape categories (see Figure 12). The first and second are tabular shapes, number one arrowhead in form, number two less pointed, but both flat. Rhombohedroid shapes were classified in two groups and numbered three and four. The number four was more elongated than number three. Three wedge-form shapes were recognized and numbered five, six, and seven, with the number six shape more elongated than five or seven. Numbers five and seven have basically the same shape but were differentiated on the basis of roundness. The number five shape would include all pebbles with a roundness classified as moderately rounded or well rounded, and the number seven shape would be more angular and include the roundness categories classified as slightly rounded or sharply angular (Holmes, 1941). The ovoid forms were assigned the numbers eight and nine and differ only in the ratio of the two longest dimensions, the latter being more elongated. The data on pebble size and pebble shape are included in the Appendix but will not be analyzed in this thesis.

Pebble Lithology

Classification of the lithology of glacial pebbles has been a research tool in glacial geomorphology for many years. Milthers (1942) studied the lithology of glacial pebbles in Norway and the Baltic region of Denmark; Holmes (1952) studied variations in glacial pebble lithology in New York;

and Anderson (1955) published an important work on the pebble lithology of the Marseilles till sheet in northeastern Illinois. More recently Ancin and Jacobs (in Johnson et al., 1972) classified the lithology of glacial pebbles from tills in east-central Illinois.

In this study only the elongate pebbles measured for orientation were classified lithologically. Thus, the sample was not a random sample. Since each sample of fifty pebbles in this study was biased by the same collection technique, they can effectively be compared lithologically. However, future investigators using the sedimentary parameter of pebble lithology should be aware of the restrictions placed on the pebble samples of this study which were classified lithologically.

Classification of each pebble's lithology was accomplished by fracturing each pebble to expose a fresh surface. The exposed fresh surface was then examined by hand lens and if necessary through a binocular microscope. Each pebble was identified and classified (see the Appendix). Five lithologic categories were recognized: crystalline, carbonate, noncarbonate clastic, chert-flint, and other. The crystalline pebbles were dominantly granites, basalts, and quartzites, but the category includes all igneous and metamorphic crystalline pebbles. Carbonate pebbles include limestone, dolomite, and all gradations between these such as limey dolomite, dolomitic limestone, and cherty dolomite. The

noncarbonate clastic pebbles were dominated by shales and siltstones but included sandstones and all gradations between these three categories. The chert-flint group of pebbles includes only those lithologies. The lithologic classification of the remaining pebbles was diverse and statistically insignificant. These were placed together in the category termed other.

Laboratory Analysis of Till Matrix

Introduction

Laboratory analysis of the till-matrix samples collected at each analysis site was performed to characterize certain physical and compositional properties of the sediment. The laboratory analysis provided numerical data on grain size, percentage of sand, silt, and clay; clay mineral composition, percentage of expandable clay minerals, illite, and chlorite plus kaolinite; and carbonate content, counts per second of calcite and dolomite. Clay mineral composition and carbonate content was determined by X-ray diffraction analysis. The color of a moist sample of the till matrix for each analysis site was identified using the Munsell Soil Color Charts (1971), and in addition a descriptive color was recorded for a dry sample. The data from the laboratory analysis for each analysis site are given in the Appendix and summarized in Tables 1, 4, and 6.

Gross (1969, pages 33-37) has described the grain-size analysis and X-ray diffraction analysis techniques used by the Illinoi (oral comm wified v sections o X-ray Diff work of D. The g was determ State Geo. amalysis. Geologica from othe Samp the labor subsample analysis, were used The 25 milli: vere add liter the to a 100 1000 mil Vigorous 30 degre the Illinois State Geological Survey and has given permission (oral communication, June 1972) for use of a slightly modified version of his description in this thesis. The sections of this chapter, entitled Grain-Size Analysis and X-ray Diffraction Analysis which follow are largely the work of D. L. Gross.

Grain-Size Analysis

The grain-size distribution of the till-matrix samples was determined in the sedimentation laboratory of the Illinois State Geological Survey by a combined hydrometer and sieve analysis. This technique has been standardized by the Geological Survey so that the data may be compared with data from other parts of the state of Illinois.

Samples from each analysis site were first air-dried in the laboratory. The original sample was split into three subsamples, one for grain-size analysis, one for X-ray analysis, and one to be stored. Fifty-five grams of sample were used for grain-size analysis.

The sample was transferred to a milkshake mixer, 25 milliliters of 4 per cent sodium hexametaphosphate (Calgon) were added, and the mixer was filled with distilled water. After thirty minutes of mixing, the sample was transferred to a 1000 milliliter settling cylinder. After filling to 1000 milliliters with distilled water, the cylinder was vigorously shaken and placed in a water bath heated to 30 degrees Centigrade. After three hours, if no flocculation

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was observed, the sample was reshaken for five minutes. A single hydrometer reading was taken after one hour and fifty-eight minutes. The hydrometers have been specially calibrated for this temperature so that they can be read directly in grams of clay present in the sample.

The sample was then wet-sieved through a 53-micron sieve to remove the sand. The sand was dried and sieved on a rotap machine for fifteen minutes to remove the greater-than-2.00-millimeters and the greater-than-62-micron fractions.

The original 55-gram sample weight was divided by the weight of the greater-than-2.00-millimeters-size fraction to give the percentage of gravel. The weight of the less-than-2.00-millimeters-size fraction was divided by the weight of the sand (2.00 millimeters to 0.062 millimeters) and the weight of the clay (less than .004 millimeters) to give the percentages of sand and clay respectively. The percentage of silt (0.062 to .004 millimeters) was obtained by subtracting the sum of the sand and clay percentages from 100 per cent. The sand, silt, and clay percentages total 100 per cent and are measurements of the less-than-2.00-millimeters fraction.

X-ray Diffraction Analysis

The composition of the clay mineral fraction and an indication of the relative abundance of calcite and dolomite were determined on oriented aggregates of the less-than-2-micron clay fraction by X-ray diffraction procedures. These analyses

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were performed by Dr. H. D. Glass in the clay mineralogy laboratory of the stratigraphy section of the Illinois State Geological Survey.

Approximately 35 grams of the whole sample were placed in a 100-milliliter beaker and distilled water added. The beaker was vigorously stirred, and a small portion of the suspension was immediately decanted into another beaker. This procedure was repeated several times and the sand and gravel fractions discarded. The clay fraction was then dispersed. If flocculation occurred, the water was carefully decanted off, more distilled water added, and the beaker restirred. This was usually sufficient to disperse the sample, but occasionally a little sodium hexametaphosphate was added.

After settling for twenty minutes, a portion of the less-than-2-micron suspension was removed with a pipette, placed on a glass slide, and air-dried.

The oriented aggregate slides were analyzed with a General Electric XRD-5 diffraction unit coupled with a horizontal goniometer and a recording diffractometer. The slides were run after exposure to ethylene glycol vapor. Quantitative analysis of the percentages of expandable clay minerals, illite, and chlorite plus kaolinite, and relative abundance of calcite and dolomite has been standardized by the Geological Survey so that these data may be compared with data from other parts of the State of Illinois.

The expandable clay minerals are identified by the presence of a peak at approximately 17 Angstroms after treatment with ethylene glycol. These minerals include montmorillonite but are here principally formed by expandable vermiculite and chlorite as well as their mixed lattice intermediates.

Illite is identified by the presence of a 10 Angstrom peak, and is defined as a nonexpanding and noncollapsing 10 Angstrom clay mineral.

Chlorite plus kaolinite is calculated by the intensity of the 7.2 Angstrom reflection. Chlorite is the dominant mineral and is characterized by a 14 Angstrom periodicity. A trace of kaolinite was present in some samples. The first-order kaolinite reflection and the second-order chlorite reflection both occur at about 7 Angstroms, but the second-order kaolinite peak and the fourth-order chlorite peak could be observed separately and the presence of kaolinite thus noted.

To calculate the clay mineral percentages, a base line was drawn on the log-normal-diffraction chart. The differences in intensity, measured in counts per second, between this base line and the 7, 10, and 17 Angstrom peaks were then recorded. The counts noted for illite were multiplied by three, the counts noted for chlorite plus kaolinite were multiplied by two, and the counts for the expandable clay minerals were left unchanged in order to equalize the scattering difference between the different clay minerals. These ratios were then converted to 100 per cent.

A dividi second in cou iniex in the first to 14 clay m inis 1 the 7 The Di sectio not be 1 Deasu count. ceasu the p size, the m resul duiy deter calci (Drie

A diffraction intensity (DI) ratio was calculated by dividing the intensity of the illite peak, in counts per second, by the intensity of the chlorite plus kaolinite peak, in counts per second. This DI ratio is used as a numerical index of weathering. Illite is not appreciably weathered in the Woodfordian Substage of Illinois, but one of the very first effects of weathering is the alteration of chlorite to 14 Angstrom vermiculite and eventually to an expandable clay mineral at about 17 Angstroms (swelling chlorite). This loss of chlorite is reflected by a loss in intensity of the 7 Angstrom peak and causes the DI ratio to increase. The Diffraction Intensity ratio is most useful in a vertical section to compare samples within a profile. However, it should not be used as an "index fossil."

The relative abundance of calcite and dolomite was measured from the oriented slide. Calcite is measured in counts per second at 29.4 degrees 2-theta and dolomite is measured at 30.9 degrees 2-theta. These measurements reflect the percentage of the minerals present and also the particle size, crystallinity, and orientation of the minerals. Thus, the measurements are not true percentages. Preliminary results of work by Moran and Gross (oral communication, July 1972) indicate that calcite counts per second as determined by these X-ray methods correlate very well with true calcite percentages as determined with the Chittack apparatus (Driemanis, 1962). The correlation coefficient, based on

134 sets of duplicate analyses, for these two parameters is .90, which indicates a very high degree of correlation. The calcite counts per second should be divided by 3.6 (the slope of the regression line for the correlation) to get a true percentage of calcite. The correlation of dolomite counts per second to true per cent dolomite is not as good, with a .56 correlation coefficient. This relationship is almost one to one; the slope of the regression line is .97, which means that dolomite counts per second are approximately equal to true per cent dolomite.

This rapid clay mineral and carbonate X-ray procedure is particularly well suited for working with large numbers of samples. It is true that the numbers reported may not represent actual clay mineral percentages, but these are reproducible data which can be used for stratigraphic correlations.

Quantitative estimation of the amount of each clay mineral component present is difficult and an approximation at best. The procedure described above is based on the assumption that the greater the quantity of a particular mineral, the more intense the diffraction peaks of that mineral will be. However, other factors such as particle size, crystallinity, and the degree of orientation on the slide also influence the peak intensities and thus limit the preciseness of quantitative estimates. Therefore, it must be realized that the clay mineral data presented in this study are useful for stratigraphic correlation, but not necessarily as precise measurements of clay mineral content.

Color Analysis

The color of the till-matrix samples was determined by comparing a moist sample of the till matrix with the chips of Munsell Soil Color Charts. The Munsell Soil Color Book (1971) has seven charts containing 196 chips. The color of each chip has a three-variable description. The three variables are hue, value, and chroma. The hue notation of a color indicates its relation to red, yellow, green, blue, or purple; the value notation indicates its lightness; and the chroma notation indicates its strength or departure from a neutral of the same lightness (Munsell Color Company, 1971). An example of a Munsell color notation would be 7.5YR 5/6, with the 7.5YR representing the hue as a combination of yellow and red, 5 as the value, and 6 as the chroma.

The technique previously described in this study for X-ray diffraction analysis required the use of a slide that contains a dry sample of the clay fraction of the till-matrix sample. This slide was used to indicate a dry sample color because the slides can be compared for color easily. The descriptive words used included the following: yellow tan, gray brown, violet, peach, salmon, green gray, gray, olive, gray tan, tan, buff tan, and red brown.

Chapter 3

CHARACTERISTICS OF THE

EUREKA-NORMAL MORPHOSTRATIGRAPHIC UNIT

Introduction

The individual morphostratigraphic units considered in Chapters 3, 4, and 5 of this thesis were delineated on the basis of definitions by Willman and Frye (1970). Portions of the Eureka, Champaign, and Bloomington Moraines (Drifts) and all of the Normal Moraine (Drift) were studied (see Figure 7). This chapter is concerned with the characteristics of the Eureka and Normal Moraines (Drifts).

By definition a morphostratigraphic unit is identified by the surface form it displays (Frye and Willman, 1960, 1962). A topographic boundary which delineates and separates the Eureka and Normal Moraines (Drifts) as two separate morphostratigraphic units is not apparent through field observation or map study except possibly in the area around Arrowsmith, Illinois. Therefore, these two morphostratigraphic units will be considered as one unit called the Eureka-Normal Moraine (Drift). Figure 13 (in pocket) is a Peoria 1:250,000 topographic contour map that may be referred to when altitudes associated with the study area are described. The distal base of the Eureka-Normal Moraine (Drift) in the study area is approximately marked by the 800 foot (243 meters) contour west of the city of Normal and by the 825 foot

(247 meters) contour east of the city. The proximal base of the Eureka-Normal Moraine (Drift) closely parallels the 800 foot (243 meters) contour except in the area where the Fletchers Moraine (Drift) exists (see Figure 4, in pocket). Here the proximal base of the moraine ranges from 800 feet (243 meters) to over 920 feet (276 meters) where the Eureka and Fletchers Moraines (Drifts) are adjoining.

The maximum relief in the study area is found in sections 7 and 18, T.23 N., R.6 E., where the crest of the Eureka-Normal Moraine (Drift) is 150 feet (45 meters) higher than the distal margin. At an altitude of approximately 950 feet (285 meters) this crest is the highest elevation in the study area. The rolling distal slopes are steeper than the more gradual gently sloping proximal slopes and the tops of the moraines are undulating. Within the moraine the more rugged topography is generally located in regions where streams have dissected or eroded the moraine. Thus, the present topography is the result of both glacial deposition and stream dissection. Possibly the best example of an area within the Eureka-Normal Moraine (Drift) that shows such effects is the Mackinaw River Valley (see Figure 13, in pocket). The Appendix contains a statement on the topography at each analysis site.

The trend of each moraine was determined to be a line connecting the crests of each morphostratigraphic unit as illustrated on Figure 14 (in pocket) and called the peak of the moraine. Fifteen analysis sites were selected along the trend of the Eureka-Normal Moraine (Drift) (see Figure 7). Six (El, E3, E5, E6, E9, and EB2) were located in a central position on the moraine; two (E7 and Ell) were located on the proximal slope; and seven (E4, E8, E10, N1, N2, N3, and N4) were located on the distal slope. The average altitude of the crests of the exposures used for analysis is 830 feet (248 meters). The average depth of analysis below these crests was 6 feet (2 meters). One analysis site (EB2) was stratigraphically below the top till of the Eureka Moraine (Drift). The stratigraphic relationships between the till at EB2 and the till at E10 will be discussed where the Congerville Section is described later in this chapter. Data from these fifteen analysis sites will be presented and analyzed in this chapter and later compared with comparable data from the other morphostratigraphic units under consideration so that the basic problem of relationships can be examined.

Physical Characteristics of the Till Matrix

The Eureka-Normal Moraine (Drift) represents some of the deposits of the Peoria Sublobe of the Lake Michigan Lobe (see Figure 3) deposited during the Woodfordian Substage of the Wisconsinan Stage in central Illinois (Willman and Frye, 1970). The Eureka-Normal Moraine (Drift) is a part of the Malden Till Member of the Wedron Formation (Willman and Frye, 1970) (see Figure 5). Willman and Frye (1970) describe the physical characteristics of the Malden Till Member as a silty, locally sandy, yellow gray to gray tan till bounded above by the Yorkville Till and below by the Tiskilwa Till.

This till can be correlated areally in McLean County with Unit 1 of Kempton, DuMontelle, and Glass (1971). These authors summarized the physical characteristics of Unit 1 as a tan or yellowish brown oxidized till or a gray or olive gray unoxidized till with average grain-size distribution of 15 per cent sand, 49 per cent silt, and 36 per cent clay, with average clay mineral composition of 2 per cent montmorillonite, 82 per cent illite and 16 per cent chlorite and kaolinite, and with a carbonate content of 10 counts per second calcite and 16 counts per second dolomite. The percentages given in the above description are based on thirty samples from McLean County. Willman and Frye (1970, Table 3, pp. 169-170) reported the results of similar analyses for four samples, two from the Eureka Moraine (Drift) and two from the Normal Moraine (Drift). Averages of the results of these four analyses show a grain-size distribution of 15 per cent sand, 41 per cent silt and 44 per cent clay, a clay mineral composition of 6 per cent expandable clays, 81 per cent illite, and 13 per cent kaolinite and chlorite, and carbonate content of 7 counts per second calcite and 23 counts per second dolomite. Location of the type section for the Malden Till Member can be found in Chapter 1 (see page 12).

⁵Montmorillonite has been used by some I.S.G.S. authors in place of expandable clays. Data produced by X-ray diffraction analysis at the Illinois State Geological Survey on clay minerals is comparable for both categories. (H. D. Glass oral communication).

The physical characteristics of the Malden and Unit One Tills are described in terms of the same variables investigated in this study. Table 1 summarizes the data, on the physical characteristics of the till matrix for the Eureka-Normal Moraine (Drift), found in the Appendix. Data on color, grain size, clay mineral composition, and carbonate content for the fifteen analysis sites arranged from east to west along the trend of the Eureka-Normal Moraine (Drift) are included in this table (see Figure 7). Analysis site EB2 is included here because its location is in the area of the Eureka Moraine (Drift), but the till lies stratigraphically below the top till unit. The till overlying EB2 is described by the data for analysis site E10. A discussion of the stratigraphic relationship between these two till members is presented later in this chapter where the Congerville Section is described.

All till samples collected along the Eureka-Normal Moraine (Drift) were calcareous and oxidized. The average depth of leaching on well-drained level upland sites is about 4 feet (1.3 meters). The oxidation is reflected in the tan to buff tan color (2.5Y 5/4) that may grade into unoxidized till at a depth ranging from 12 to 15 feet (3.6 to 4.5 meters). Average depth of analysis for Eureka-Normal till samples was approximately 6 feet (2 meters) and ranged from 5 to 10 feet (1.5 to 3 meters) below the crest of the exposure.

The averages reported in Table 1 for all Eureka-Normal analysis sites, exclusive of EB2, are remarkably consistent with the averages for Eureka and Normal Moraine (Drift)

Till
Eureka-Normal
Matrix
Till
the
g
Characteristics
Physical

Table l

				_															
Carbonate Content	er Second	Dolomite	6	15	19	13	15	27	21	29	27	16	22	17	13	26	26	19	
	Counts Pe	Calcite*	<i>c</i> •	I	8	9	,	22	9	10	7	וו	12	7	¢•	10	22	7	
nposition	%Chlor.	£ Kaol.	15	16	17	16	16	12	16	13	14	14	13	18	1	13	16	15	
lineral Con	%Illite		81	81	80	78	79	81	81	83	82	83	83	78	84	82	2	81	
Clay M	%Exp.	Clay	4	m	м	9	S	7	m	4	4	M	4	4	ю	ы	14	4	
bution	%Clay		34	37	34	33	18	30	29	34	38	42	38	40	42	29	32	34	o i i to
ze Distr	%Silt		38	39	38	48	45	53	43	46	40	41	45	40	38	55	41	44	tanna of
Grain Si	%Sand	848952993658255005		16	27	22	ahla avie												
	Color		2.5Y5/4	2.5Y5/4	2.5Y5/4	2.5Y5/4	2.5Y5/4	10YR5/2	2.5Y5/4	7.5YR5/4	2.5Y5/4	2 due tion							
Anal.	Site		E7	E6	ES	IN	EI	N3	E8	N4	63	N2	E4	E3	EIO	EII	EB2	Avg.#	*

/ questionable existence of calcife - analyzed for calcife but none existed both are considered zero for purpose of averaging calcife # Data for analysis site EB2 are not **inc**luded in these averages

samples for the Malden Till Member of Willman and Frye (1970) and with the averages for Unit 1 of Kempton, DuMontelle, and Glass (1971). The only notable difference exists in the sand content of the grain-size distribution which shows sand content somewhat higher than the 15 per cent reported by both Willman and Frye (1970) and Kempton, DuMontelle, and Glass (1971). However, Willman and Frye (1970) described the Malden Till as a "silty locally sandy till" and the locally sandy condition appears to exist at analysis sites El, E5, E6, E7, and E8 (see Figure 7). The till-matrix composition at analysis site N3 is somewhat anomalous, but the color difference and high calcite content are within the ranges for the Malden Till Member reported for these variables by Willman and Frye (1970). The averages of the numerical data in Table 1 for the Eureka-Normal till samples are also comparable to the numerical data presented for the Snider Till Member in the Danville, Illinois region (Johnson, Gross, and Moran, 1971; Johnson et al., 1972, Table 2, page 7). The relationship between Malden and Snider Till will be discussed in Chapters 4 and 5 of this thesis. Thus, the rock-stratigraphic unit to which the Eureka-Normal Moraine (Drift) in the study area of this thesis is assigned is the Malden Till Member of the Wedron Formation. This conclusion, which is consistent with the findings of Willman and Frye (1970), is based on the high degree of correlation between the physical characteristics of the till matrix reported in Table 1 and the description of the Malden Till.

Characteristics of Till Pebbles

Introduction

Data on pebble lithology, long-axis pebble orientation, size, and shape were collected for a fifty-pebble sample at each of the fifteen analysis sites along the trend of the Eureka-Normal Moraine (Drift) (see the Appendix and Figure 7). This study is concerned with the relationship between long-axis pebble orientation, pebble lithology, rock stratigraphy, and morphostratigraphy. Pebble size and shape data are included in the Appendix but will not be analyzed in this thesis.

Holmes (1941) was the first in this country to correlate long-axis pebble orientation with the direction of ice movement associated with an area of ground moraine in New York. He recognized both parallel and transverse patterns of longaxis pebble orientation and associated these with the direction of ice movement. Hoppe (1952) correlated the parallel and transverse patterns of long-axis orientation to ridge trends of hummocky moraines in Sweden but found little or no pebble orientation in the hollows between ridges. Wright (1962) correlated patterns of pebble alignment with the axes of drumlins in Minnesota and found them to parallel one another. Today most if not all investigators using the techniques of till-fabric analysis would agree that there is a strong tendency for the long axis of an elongated pebble within glacial till to be aligned either parallel to or perpendicular to the direction of ice flow if the till was deposited in association

with active ice. Thus, in instances where associated landforms also bear a specific areal relationship to an active glacier pebble alignment and landform trend may be directly related. Possible realignment or reorientation of pebbles in glacial till resulting from events subsequent to deposition were reported by MacClintock and Dreimanis (1964) and by Ramsden and Westqate (1971). In Alberta, Canada, Ramsden and Westqate (1971) present evidence to support a till-fabric orientation change produced by the advance of a later glacier. Harrison (1957b) found that the clay till fabric of the Marseilles Moraine in northeastern Illinois (see Figure 4, in pocket) showed a greater tendency toward perpendicular than parallel orientation to the trend of the moraine but many fabrics appeared to have been reoriented slightly to an unpredictable alignment. It is worth noting at this time that the lobate form of the Marseilles Moraine is similar to that of the Eureka-Normal Moraine (Drift).

Another study which reports the findings of till-fabric analyses on lobate form moraines was written by Andrews and Smithson (1965). These authors describe four morainal forms called linear, hooked, s-shaped, and asymmetrical. The latter is a parabolic lobate form. The results of the Andrews and Smithson (1965) study reveal that the orientation and dip strength in asymmetrical form moraines is poorer than the orientation and dip strength in the other three forms. They also concluded the orientation strength was no different from proximal to distal slope and that dip strength was only slightly higher on the distal slope in asymmetrical form moraines.

These conclusions have special relevance to the discussion of pebble-orientation data described later in this chapter and in Chapters 4 and 5.

Till Pebble Orientation

The horizontal long-axis orientation of a fifty-pebble sample was determined in the field to the nearest five degrees (see Chapter 2 for a discussion of the field technique). Such data have been represented and analyzed in various ways. Holmes (1941) used rose and contoured diagrams in his analysis. Harrison (1957b) used rose diagrams and maps with preferred orientations indicated by arrows. Krumbein (1939), Kauranne (1960), and Andrews and Shimizu (1966) have all suggested statistical techniques for the analysis of orientation data but they also used orientation diagrams in these works to show associations. In this study it was concluded that pebble orientation could best be analyzed and compared by using rose diagrams centered at the location of each analysis site on a map of the study area showing the morphostratigraphic units (see Figure 14, in pocket).

Long-axis pebble orientation for the fourteen analysis sites in Malden Till along the Eureka-Normal Moraine (Drift) ^{appear} consistent with the findings of both Harrison (1957b), and Andrews and Smithson (1965). These works will be used for ^{comparison} in the remainder of this section as the findings of the till-fabric analyses of this study are reported.
Comparison of data from site E10 on the distal slope with that of site E11 on the proximal slope does show the distal orientation strength is somewhat better than the proximal orientation strength, but the consistency of this result is not demonstrated by the orientations at the remainder of the analysis sites on this moraine. The tendency for longaxis orientations to be parallel to the trend of the moraine exists at analysis sites E8, N4, E9, and N2 but is not as pronounced as the parallel tendency in the region near the interlobate area at analysis sites E5, E6, and E7 (see Figure 14, in pocket). A secondary mode orthogonal to the primary mode is present in the orientations at analysis sites N2 and N4. At all the remaining sites except N3, which is trimodal, the tendency of the long-axis orientation is toward a perpendicular to the trend of the moraine.

Both long-axis orientation and dip orientation and strength are considered in most till-fabric studies. These variables have a high degree of correlation and when used in combination strengthen the fabric representation. Holmes (1941), Harrison (1957a), West and Donner (1957), and others have used both measurements. Dip orientation and strength will be discussed in this study only when the rose diagram of the long-axis orientation shows a poor or unoriented pattern.

At analysis site N3, which has a trimodal long-axis orientation, and at analysis sites Ell, E8, and N1 where the long-axis orientation is poor, dip orientation is considered an alternate fabric determinant. At N3, which has little or

n0 to 10 st CI to to Đê : 01 t(Ca sÌ ť. i 2 . ţ à 5 no orientation, there exists a moderately strong dip orientation to the south-southeast. At N1, where there is a weak unimodal long-axis orientation to the north-northwest, the same moderately strong dip orientation to the south-southeast exists. The dip orientation is moderately strong to the west at E8 and weak to the northeast at E11. The dip orientation shows a tendency toward a perpendicular to the trend of the moraine except at E8 where it tends toward a parallel alignment to the trend of the moraine.

Figure 14 (in pocket) shows a greater tendency of fabric orientation toward a perpendicular than parallel alignment to the trend of the Eureka-Normal Moraine (Drift). However, careful examination reveals that a majority of the fabrics show a discrepancy between a perpendicular to the trend of the moraine and the actual pebble orientation. Harrison (1957b), in investigating the very arcuate Marseilles Moraine (see Figure 4, in pocket), and Andrews and Smithson (1965), in investigating the asymmetrical form moraines on Baffin Island, found their fabrics to be slightly askew from the perpendicular alignment to the moraine trend they expected. Harrison (1957b) suggested that these differences were the result of a till deformation process which altered the original depositional environment of the till. Andrews and Smithson (1965) associated the asymmetrical form of moraines to a process of pushing and overriding which altered and weakened the fabric strength. If Harrison (1957b) and Andrews and Smithson (1965) are correct regarding the relationship of the long-axis

pebble orientation to the trend of a lobate form moraine, then the data of this study supports the contention that the original depositional environment of the pebbles in Eureka-Normal Moraine (Drift) was altered slightly in the formation of this moraine. As suggested by Glen, Donner, and West (1957), great care was taken to avoid locations for analysis sites in this study where alteration of till fabric as a result of gravitational forces would be a factor (see Chapter 2, selection of analysis sites). The processes by which such reorientation might take place can only be speculated on the basis of available information. However, it would appear that if a reorientation did take place it is most likely that the alteration was accomplished by a pulsation or slight readvance of the Woodfordian ice sheet after the original formation of much of the Eureka-Normal Moraine (Drift).

Till Pebble Lithology

The fifty pebbles measured for long-axis orientation at each analysis site were classified lithologically in the laboratory. Chapter 2 contains an account of the laboratory procedure and the resulting data can be found in the Appendix. The purpose of the investigation of pebble lithology was to show if a relationship exists between either pebble lithology and morphostratigraphy or pebble lithology and rock stratigraphy.

At this point it is appropriate to mention some of the deficiencies of the pebble lithology analysis of this study.

f:e vt.e ļ1 50 101 цт(of 00 bu sp je зр 10 t ħ CT pe <u>c]</u> Ŋ (] CC E: in tr <u>}</u>e The sample size of fifty pebbles appears too small, especially when pebbles without a long axis were not considered. Anderson (1955) used 100 pebbles for each sample location when studying the pebble lithology of the Marseilles Drift in northeastern Illinois (see Figure 4, in pocket) and found a great deal of variability both from proximal to distal side of the moraine as well as laterally. He also found local concentrations of particular lithologic types and his distribution of sample sites was more dense than the 1 to 5 mile spacing of analysis sites for this study. In spite of these deficiencies slight variations in pebble lithology, which appear worthy of consideration, did exist between both rock-stratigraphic units and morphostratigraphic units.

To accentuate these differences and facilitate analysis, the various pebbles were grouped into lithologic categories. The procedure involved first dividing the pebbles into crystalline and noncrystalline groups. Then the noncrystalline pebbles were separated into four groups: carbonate, noncarbonate clastic, chert-flint, and other (see Chapter 2 for lithologic types included in each group).

Lithologic characteristics of Eureka-Normal Moraine (Drift) pebbles are summarized in Table 2. Two general conclusions can be drawn about the lithologic composition of Eureka-Normal Moraine (Drift) pebbles in comparison to pebbles from till associated with other morphostratigraphic units treated in this study. First, the average number of crystalline pebbles is very low, and, second, the average number of

Table 2

Pebble Lithology--Eureka-Normal Till

Analysis	%Crystalline	od sei	NonCrystallin	ne	e a 	_
Site	re 1. Riden Gelin tion) this	%Carbonate	%NonCarbonate Clastic	%Chert Flint	%Other	1
E7	9 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0 0	64	16	14	0	1
E6	12	52	24	12	0	-
ES	10	54	26	80	2	
TN	26	42	14	14	4	
EI	8	68	20	4	0	
N3	9	72	18	4	0	
E8	9	50	26	18	0	
N4	24	54	12	10	0	-
E9	4	80	9	10	0	-
N2	9	64	22	8	0	
E4	80	56	10	24	2	-
E3	10	70	14	9	0	-
EIO	8	62	14	16	0	
ELL	80	64	16	12	0	
EB2	20	62	12	4	2	
Avg.#	10	61	17	11	1	
#Aver	are does not include	data from analvsi	s site ER2.	-		
-> ··· II	and a state of the	and a state when a state				

carbonates, the majority most probably Silurian dolomites, is very high. These conclusions were drawn after the data in Table 2 were compared with the data in Tables 5 and 7. However, an effective comparison of these data cannot be made until the characteristics of the Bloomington and Champaign Moraines (Drifts) have been discussed. Therefore, a more detailed discussion of pebble lithology will be presented in Chapters 4 and 5 where the lithology data for pebbles from these units are discussed.

Congerville Section

The Congerville Section is located areally on the Eureka Moraine (Drift) and is included in this chapter because of its location. The importance of this stratigraphic section is that it establishes the relationship between three Woodfordian units: the Richland Loess, the Malden Till, and the Tiskilwa Till.

Data on physical characteristics of this section can be found in Table 3. The upper two tills in this section are defined as Malden and Tiskilwa Till. The data for samples from this section are comparable to averages for samples from analysis sites defined to have these till units. A sharp line of demarcation between the Malden Till and Tiskilwa Till can be found in this section. A boulder line at the base of a thin zone of outwash clearly establishes the contact and stratigraphic relationship of these two till members.

Physical Characteristics of the Congerville Section Samples

Table 3

		<u> </u>	_				_			
Long	Axis Orient		N30E		N20W					
te Content	Per Second Dolomite	•	23	19	32	12	20	10	19	20
Carbona	Counts Calcite	1	10	14	24	7	13	6	12	13
mposition	%Chlor. Kaol.	6	14	12	15	13	14	17	14	17
/ Mineral Co	%Illite	55	82	81	67	81	81	78	れ	65
CLay	%Exp. Clay	36	4	7	18	9	S	S	15	18
bution	%Clay	40	31	22	36	15	39	26	32	¥
e Distril	%Silt	53	50	33	36	41	41	64	4	37
Grain Siz	ßand	2	19	45	28	\$	20	10	24	29
	No.	6	8(E10)	7	6(EB2)	2	4	м	8	Ч

Table 3

Physical Characteristics of the Congerville Section Samples

. •

	Grain Siz	se Distri	bution	Clay	y Mineral Co	mposition	Carbona	te Content	Long
	%Sand	%silt	%Clay	%Exp. Clay	%Illite	%Chlor. Kaol.	Counts Calcite	Per Second Dolomite	Axis Orient
Γ	6	53	40	36	55	6		1	
6	19	50	31	4	82	14	10	23	N30E
	45	33	22	7	81	12	14	19	
5)	28	36	36	18	67	15	24	32	N20W
	4	41	15	9	81	13	7	12	
	20	41	39	S	81	14	13	20	
	0T	64	26	S	78	17	თ	10	
	24	4	32	15	1 2	14	12	19	
	29	37	4	18	65	17	13	20	

Congerville Section

	Me	asured in	roadcuts	in	NE	SE	NE	Sec	. 3,	T.25	Ν.,
R.1	W.,	Woodford	County,	111:	inoi	8,	197	71,	1972	2.	·

		Thickness (ft)
Pleistocene Series Wisconsinan Stage Woodfordian Substage Richland Loess		
9. Loess; tan to tan brown, leached; contains modern surface soil		2.0
Wedron Formation Malden Till Member Eureka Drift		
 Till, silty, upper part leached; buff tan to light brown; oxidized 		3.0
 Sand and gravel, calcareous, dark tan, shape contact boulder line at base 		0.5
Tiskilwa Till Member		
 Till, sandy, pink to red brown, calcareous, compact angular blocky in upper part grading to unconsolidated near base, contains silt lenses 		9.5
5. Sand and gravel, calcareous, yellow		2.0
can		2.0
Illinoian Stage Undifferentiated		
4, Till, silty, calcareous, buff tan to tan		2.0
3. Silt, yellow tan to lemon		0.5
 Till, silty, olive to olive tan to pale green 		1.0
1. Till, sandy, dark tan to gray to brown		9.5
	Total	30.0

Long-axis pebble orientation data for the top two till units is also included in Table 3. The fabric of the Malden Till has a northeasterly orientation in this section and the fabric of the Tiskilwa Till has north-northwest orientation. The original depositional mode of Tiskilwa Till is believed to be represented by the fabric orientation at analysis site Bl5 (see Figure 14, in pocket). The long-axis pebble orientation at analysis site EB2 shows the reorientation effect that the glacial ice, which deposited the Eureka Drift, had on the Tiskilwa Till in the till plain behind the Bloomington Moraine (Drift) (see page 88 for a more detailed discussion).

The Woodfordian Units overlie a sequence of undifferentiated units believed to be Illinoian in age. The important aspect of the Congerville Section is that it clearly establishes the stratigraphic relationship between the Tiskilwa and Malden Till Members of this study.

Conclusions

The major conclusions of this chapter are the following: (1) The Eureka-Normal Moraine (Drift) is one morphostratigraphic unit. (2) Physical characteristics of the surface till within this unit indicate that it belongs to the Malden Till Member of the Wedron Formation. (3) Pebble orientation findings reveal that the long axis of elongate pebbles within this till prefer a perpendicular alignment to the trend of the moraine. These orientations indicate a Lake Michigan Lobe

source. And (4) pebble lithology data indicates that a lower percentage of crystalline pebbles and a larger percentage of carbonate pebbles are embedded in this till than was found in tills from the other units of this study.

Chapter 4

CHARACTERISTICS OF THE CHAMPAIGN MORPHOSTRATIGRAPHIC UNIT AND ASSOCIATED INTERLOBATE AREA

Introduction

The Champaign Moraine (Drift) as delineated by Willman and Frye (1970) was studied between the two valleys of the Sangamon River which are approximated by analysis site locations Cl and C6 (see Figures 4 and 7). Near C6 the Sangamon River Valley separates the Champaign Moraine (Drift) from the interlobate section to the northeast, and near Cl the Sangamon River traverses the Champaign Moraine (Drift) and separates it into two segments.

The base of both the proximal and distal slope of the Champaign Moraine (Drift) in the study area is clearly defined topographically by the 750 foot (225 meters) contour (see Figure 13, in pocket). However, the topography of the Champaign Moraine (Drift) is more subdued than that of either the Bloomington or Eureka-Normal Moraines (Drifts). The distal slopes tend to be slightly steeper than the proximal slopes but both could be categorized as gently sloping to rolling (see Figure 10). In the study area the total local relief on the Champaign Moraine is approximately 80 feet (24 meters), but the average relief is less than 50 feet (15 meters) per square mile. A maximum altitude of about 830 feet (248 meters) is attained in a number of locations

between Bellflower and Saybrook (see Figure 13, in pocket). Though the moraine is the least prominent in the study area, it is the best defined topographically in terms of both extent and boundaries.

Five analysis sites were selected along an 18 mile (29 kilometers) extent of Champaign Moraine (Drift) within the study area. Two analysis sites (C2 and C3) were in the central part of the moraine and three (C1, C4, and C6) were on the distal slope (see Figure 7). The average altitude of the exposures is 790 feet (237 meters) and the excavations average about 11 feet (3.3 meters) below the present land surface.

The associated interlobate area of the Champaign Moraine (Drift) is defined as the area of BD analysis-site locations, and the criterion used to establish this area is the same as that used to separate it from the area of BP analysis-site locations (see page 28 and Figure 9). The base of this area is approximated by the 750 foot contour (225 meters), and the altitude in the central part of this region is slightly more than 900 feet (270 meters). For purposes of discussion and analysis, the data from the three BD analysis sites **are** considered along with the data from the Champaign Moraine (Drift) so that the relationship between the interlobate area, which Willman and Frye (1970) refer to as Bloomington Moraine (Drift), and the Champaign Moraine (Drift) can be established.

Physical Characteristics of the Till Matrix

The Erie Lobe has been divided into sublobes including the Decatur which deposited a sequence of drifts (see Figure 4, in pocket) in east-central Illinois during the Woodfordian Substage of the Wisconsinan Stage (Willman and Frye, 1970) (see Figure 3). The Champaign Moraine (Drift) deposited by the Decatur Sublobe is a part of the Wedron Formation rockstratigraphic unit that was undifferentiated by Willman and Frye (1970). Johnson, Gross, and Moran (1971) have studied the previously undifferentiated Wedron Formation in the Danville, Illinois region and have recognized three till members that they correlate with proposed units in the McLean County region of the study area.

Figure 6 illustrates the relationship between the Batestown, Unit 2, and Snider Tills. This figure also correlates these three units with tills that were named prior to the works in which the Batestown, Unit 2, and Snider Tills were defined. A discussion of the physical characteristics of the Batestown, Unit 2, and Snider Tills follows so that their relationships to the till of the Champaign Moraine (Drift) and associated interlobate area can be established.

The Batestown Till Member is described as a gray or dark gray often silty till which oxidizes to a characteristic light olive brown (Johnson et al., 1972). It is known to exist stratigraphically between the Snider and Glenburn Till Members in the Danville region (Johnson, Gross, and Moran, 1971). Some physical characteristics of the Batestown Till

Member are (1) a grain-size distribution of 28 per cent sand, 38 per cent silt, and 34 per cent clay; (2) a clay mineral composition of 3 per cent expandable clay, 80 per cent illite, and 17 per cent kaolinite and chlorite; and (3) a carbonate content of 5 per cent calcite and 19 per cent dolomite (Johnson, Gross, and Moran, 1971; and Johnson et al., 1972).

Unit 2 is described as a gray silty till, but a pinkish tint has been noted in some samples on the outer margin of the till (Kempton, DuMontelle, and Glass, 1971). The averages for 150 samples of Unit 2 till showed (1) a grain-size distribution of 27 per cent sand, 45 per cent silt, and 28 per cent clay; (2) a clay mineral composition of 3 per cent montmorillonite (see footnote 5, page 48), 79 per cent illite, and 18 per cent chlorite and kaolinite; and (3) a carbonate content of 14 counts per second calcite and 21 counts per second dolomite (Kempton, DuMontelle, and Glass, 1971).

The Snider Till Member is a gray brown to light olive brown till where oxidized and a dark gray where unoxidized (Johnson, Gross, and Moran, 1971). In comparison with the Batestown Till it contains less sand and slightly more illite (Johnson, Gross, and Moran, 1971; Johnson et al., 1972). Johnson, Gross, and Moran (1971) reported the average physical characteristics of the Snider Till Member to be (1) grain-size distribution 19 per cent sand, 45 per cent silt, and 36 per cent clay; (2) clay mineral composition 3 per cent expandables, 85 per cent illite, and 12 per cent kaolinite and chlorite; and (3) carbonate content 6 per cent calcite and 19 per cent dolomite.

With some of the physical characteristics of the Batestown, Unit 2, and Snider Tills established, a discussion of the data provided by analyses of till-matrix samples from the Champaign Moraine (Drift) and the associated interlobate area may be presented. These data are contained in the Appendix and summarized in Table 4. Averages for the BD locations and C locations are figured separately so that they may be compared. The data in the table are arranged in order from north to south along the trend of the moraine (see Figure 7).

Till samples from analysis sites C2 and C4 were from a depth of 15 feet (4.5 meters) and were an unoxidized gray. The till-matrix samples from locations C1 and C3 came from a depth of 9 feet (2.7 meters) and 5 feet (1.5 meters) respectively. Their colors were a greenish gray to olive gray, the characteristic oxidation colors for both the Snider and Batestown Till Members (Johnson et al., 1972). All the remaining samples listed in Table 4 came from shallow depth analysis sites, (see the Appendix for depths) were oxidized but calcareous, and have a tan color.

Comparison of the physical characteristics of the till samples of Champaign Moraine (Drift) with those from the BD locations in the interlobate area reveal no major consistent differences. From the numerical data in Table 4, it can be concluded that the till from BD locations on the interlobate and till from the analysis sites on the Champaign Moraine (Drift) are the same till.

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Physical Characteristics of the Till Matrix--Champaign Till and Associated Interlobate Till

-					_
e Content	er Second Dolomite	16 16 21 17 20 23 23	18	18	
Carbonat	Counts P Calcite*	8 7 9 11	7	5	
nposition	%Chlor. & Kaol.	13 12 12 13 15 18	16	12	
Mineral Com	%Illite	82 85 85 79 79 79	82	84	
Clay 1	%Exp. Clay	50 M M A A A A M	2	4	
ribution	%Clay	35 38 37 37 37 36 42 43 36	43	37	
ze Distr	%Silt	36 44 35 39 39 38	37	40	
Grain Si	%Sand	29 22 19 18 18 26	20	23	
	Color	2.5Y5/4 2.5Y5/4 2.5Y5/4 2.5Y5/4 2.5Y5/4 2.5Y5/2 2.5Y5/2 2.5Y4/2	Variable#	2.5Y5/4	
Anal.	Site	BD8 BD7 BD5 C6 C3 C4 C1 C1	Avg.C	Avg.BD	

Due primarily to various degrees of oxidation
* ? questionable existence of calcite
- analyzed for calcite but none existed
both are considered zero for purpose of averaging calcite

A comparison of the numerical data for physical characteristics of the Batestown Till Member and its correlative Unit 2 with the numerical data of Table 4 reveals two differences: (1) the Batestown Till and Unit 2 have a slightly sandier texture, and (2) the Batestown Till and Unit 2 have a little less illite. However, these differences are the same differences that separate the Batestown Till Member from the Snider Till Member. On this basis a correlation, though somewhat uncertain, may be drawn between the Snider Till Member and the various till samples whose physical characteristics are summarized in Table 4. This correlation of numerical data is strong enough that it can be concluded that the till on the Decatur Sublobe side of the Bloomington interlobate (see Figure 9) and the till of the Champaign Moraine (Drift) in the study area of this thesis are a part of the Snider Till Member of the Wedron Formation rock-stratigraphic unit. This conclusion is not inconsistent with the correlation drawn by Johnson, Gross, and Moran (1971), who correlated the Batestown Till with both Unit 2 and upper till in the Champaign-Urbana region of Kempton, DuMontelle, and Glass (1971), because the latter authors did not define the till in the section of Champaign Moraine (Drift) studied in this thesis. The correlation between the till described in Table 4 and the Batestown Till still remains a possibility on the basis of data from analysis sites BD8, C6, and C1 because of the sandier texture of the till at these locations. However, the high illite content in the clay mineral fraction at BD8 and C6

agree with the Snider Till correlation. Therefore, the till samples from analysis site Cl are the only ones which might be correlated with the Batestown Till. These findings are not inconsistent with the findings of Willman and Frye (1970) because the till of the Champaign Moraine (Drift) was not differentiated by them and because there still remains the possibility that the Malden Till Member, to which they assigned the interlobate region, can be correlated with the Snider Till Member (this correlation will be suggested later in this study).

Characteristics of Till Pebbles

Introduction

The purpose of this section is to discuss and analyze first, the relationships between long-axis pebble orientation, and both rock stratigraphy and morphostratigraphy; and, second, the relationships between the lithologic composition of pebbles and both rock stratigraphy and morphostratigraphy.

Data on pebble orientation, lithology, size, and shape for a fifty-pebble sample from each analysis site on the Champaign Moraine (Drift) and from each BD analysis site in the associated interlobate area can be found in the Appendix. Data on pebble size and shape were not analyzed in this study.

Till Pebble Orientation

Long-axis pebble orientation data for the analysis sites along the Champaign Moraine (Drift) and the BD locations in the associated interlobate region are illustrated on

Figure 14 (in pocket). Analysis sites Cl, C4, and C6 are located on the distal side of the moraine while C2 and C3 are near the crest. The three analysis sites in the associated interlobate region would be on the proximal slope if this till was deposited by the Decatur Sublobe. However, Willman and Frye (1970) have mapped the drift of this segment as having been deposited by the Peoria Sublobe which, if correct, would place the analysis sites on the distal side of the Bloomington Moraine (Drift). Regardless of the interpretation selected, assuming Andrews and Smithson (1965) are correct, the position of analysis site with respect to the outline of the moraine makes very little difference in pebble orientation strength when asymmetrical (parabolic-lobate) form moraines are being considered. Therefore, the pebble orientation data from both the Champaign Moraine (Drift) and associated interlobate area will be discussed together and consideration of location with respect to outline of the moraine will not be included in a discussion of the interlobate area orientations.

Reasonably strong long-axis orientation patterns exist at seven of the eight locations studied (see Figure 14, in pocket). The one exception is at analysis site Cl where the long-axis orientation is bimodal and weak. The alternate measurement of dip strength is strong to the southwest for this location. Such an orientation at location Cl tends perpendicular to the trend of the moraine. The tendency towards a long-axis pebble orientation which is orthogonal to the trend of the moraine is more prevalent along the Champaign

Moraine (Drift) than along the Eureka-Normal Moraine (Drift) previously described. In the associated interlobate section, however, the tendency of the primary pebble fabric is to become parallel to the morainal unit with the secondary mode tending toward the perpendicular. This important relationship is also shown in the interlobate area associated with the Bloomington Moraine (Drift) and is elaborated upon in Chapters 5 and 6. Comparison of the actual long-axis pebble orientations as represented by the rose diagrams on Figure 14 (in pocket) with the true perpendicular to the trend of the moraine once again reveals a discrepancy. The orientations are slightly askew from a true perpendicular indicating, as was the case with the Eureka-Normal Moraine (Drift), that a pulsation or readvance of a retreating ice sheet may have reoriented the fabrics slightly from their original depositional environment. This possibility is based on the long-axis orientation findings of this study and the previous work of Harrison (1957b) and Andrews and Smithson (1965).

Till Pebble Lithology

The lithologic classification of each pebble from analysis sites along the Champaign Moraine (Drift) and the BD locations of the interlobate area can be found in the Appendix. These data are categorized and summarized in Table 5. A comparison of the data for C locations with the data for BD locations reveals a small difference in percentage of crystallines and a slightly larger difference for percentage of noncarbonate clastics. The latter difference can be

Table 5

Pebble Lithology--Champaign Till and Associated Interlobate Till

Analysis	%Crystalline		NonCrystalline		
Site		%Carbonate	%NonCarbonate Clastic	%Chert Flint	%Other
ង្គីម៉ី ឌីងូនភូល្អ ទទួម ខ្លួនភូលូទ	16 12 14 18 18 18	48 66 55 66 57 52 53 53 53 54 55 55 55 55 55 55 55 55 55 55 55 55	20 28 12 12 28 20 20 20 20 20 20 20 20 20 20 20 20 20	14 16 102 24 102 24 24 24 24 24 24 24 24 24 24 24 24 24	0000000
Avg.BD	12	51	23	13	ч
Avg.C	18	52	15	14	-1

attributed to a local concentration of shale pebbles in the area of the interlobate. Anderson (1955) believes that the sedimentary parameter of pebble lithology can be used to distinguish deposits from different lobes. Within the study area the only discernible differences in pebble lithology between the samples from the Peoria and Decatur Sublobes are reflected by the 9 per cent difference in carbonates. It appears that when the sample size is large (700 Eureka-Normal pebbles versus 400 Champaign-BD pebbles) that this degree of variation may be significant enough to provide a basis for determining primary source.

Conclusions

The major conclusions of this chapter are the following: (1) The till within the Champaign Moraine (Drift) and the till within the associated interlobate area are both correlated with the rock-stratigraphic unit named the Snider Till Member of the Wedron Formation. (2) The long-axis pebble orientations of the Champaign Moraine (Drift) tend to be orthogonal to the trend of the morphostratigraphic unit outside of the interlobate then become parallel to the landform in this area. These orientations indicate an Erie Lobe source for this till. And (3) the pebble lithology data from the Champaign Moraine (Drift) and associated interlobate indicate that pebbles within till from an Erie Lobe source have approximately 9 per cent fewer carbonate pebbles than pebbles within till from a Lake Michigan Lobe source.

Chapter 5

CHARACTERISTICS OF THE BLOOMINGTON MORPHOSTRATIGRAPHIC UNIT AND ASSOCIATED INTERLOBATE AREA

Introduction

The Bloomington Moraine (Drift) as defined by Willman and Frye (1970) was studied from the Tazewell-McLean County line on the west to the interlobate area just east of the Ford-McLean County line on the east (see Figure 4, in pocket). This section of the Bloomington Moraine (Drift) has the most massive topographic form of any of the morphostratigraphic units under consideration in this study (see Figure 13, in pocket). The local relief repeatedly exceeds 100 feet (30 meters) and the moraine varies in width from 1.5 to 4.5 miles (2.4 to 7.2 kilometers). The areal plan of the Bloomington Moraine (Drift) does not show the consistent arcuate lobate form that is exhibited by the Eureka-Normal and Champaign Moraines (Drifts) (see Figure 4, in pocket).

Topographically, it can be divided into three sections: (1) a section west of Bloomington, Illinois, which has its own small lobate form; (2) a central section which is almost linear in form; and (3) an interlobate section which has a form similar to the other moraines that merge to form this area. The western section is separated from the central section by the valley of the Sugar Creek (see Figure 13, in pocket). This western section has a distal base which is approximated

by the 750 foot (225 meters) contour and a maximum altitude of 883 feet (265 meters). The central section is separated from the interlobate section by the valley of the Sangamon River. The central section has a distal base of approximately 800 feet (243 meters), and crest altitudes of over 900 feet (265 meters) are common in this part of the Bloomington Moraine (Drift) giving it a massive and prominent appearance (see Figure 13, in pocket). Here the proximal slopes are more gently rolling than the steeper distal slopes due in part to the higher altitude of the proximal base. This base is somewhat variable but generally parallels the 800 to 850 foot (243 to 265 meters) contour depending on location. The third section is the associated interlobate area of the Bloomington Moraine (Drift). This section is defined by the BP analysis site locations and the areal extent is illustrated on Figure 9. The base of this area is approximately 825 feet (260 meters) and the crest altitude exceeds 900 feet (265 meters).

At four places stream erosion has resulted in extreme dissection of the Bloomington Moraine (Drift) (see Figure 13, in pocket). Sugar Creek, Kickapoo Creek, and the Sangamon River traverse the Bloomington Moraine (Drift) while the North Fork of Salt Creek partially dissects this same moraine. The topography near their associated valleys is more rugged and very steep slopes are common (see Figure 10).

Fifteen analysis sites were selected for the purpose of collecting samples from the Bloomington Moraine (Drift) (see Figure 7). Two sites were on the proximal slope, four

were in the central area near the crest, five were on the distal slope, and four were situated along the west flank of the interlobate section (see the Appendix for specific locations). Average depth of analysis below the top of the soil profile was approximately 9 feet (3 meters) with excavations ranging from 4 feet (1.3 meters) to 25 feet (7.5 meters) below the surface of the Bloomington Moraine (Drift). The surface of this drift above the analysis sites had an average altitude of 834 feet (250 meters). Data on physical characteristics of the till matrix, long-axis pebble orientation, and the lithology of pebbles collected from the fifteen analysis sites on the Bloomington Moraine (Drift) are contained in the Appendix. These data will be compiled, summarized, and presented in this chapter so that the basic question of relationships between morphostratigraphy, rock stratigraphy, pebble lithology, and pebble orientation can be answered.

Physical Characteristics of the Till Matrix

The Bloomington Moraine (Drift) in the study area of this thesis does not have the consistency found in the Eureka-Normal and Champaign Moraines (Drifts). Both Willman and Frye (1970) and Kempton, DuMontelle, and Glass (1971) recognized two till units in this part of the study area. Willman and Frye (1970) mapped the till on the Bloomington Moraine (Drift) west of Bloomington, Illinois, as belonging to the Tiskilwa Till Member and the till east of that same city as a part of the Malden Till Member (see Figure 5). The latter authors describe the till on the Bloomington Moraine (Drift) in McLean County to

the west of Danvers,⁶ Illinois, as Unit 4 Till and the till in McLean County to the east of the Danvers area as Unit 2 Till. These authors also postulated and mapped a third till named Unit 1 overlying the proximal slope of this moraine in places. The Malden Till and Unit 1 were previously described in Chapter 3, and Unit 2 was described in Chapter 4.

The Tiskilwa Till was described by Willman and Frye (1970) as a sandy, pink tan to reddish tan brown till (usually called pink) that is overlain by the Malden Till and may rest on the Delavan, Esmond, or Lee Center Till Members. Some of the physical characteristics of the Tiskilwa Till Member are represented by data from two samples of Bloomington Moraine (Drift) reported by Willman and Frye (1970) to contain this unit. Averages show (1) a grain-size distribution of approximately 29 per cent sand, 37 per cent silt, and 34 per cent clay; (2) a clay mineral composition of 18 per cent expandable clays, 68 per cent illite, and 14 per cent kaolinite and chlorite; and (3) a carbonate composition of 23 counts per second calcite and 36 counts per second dolomite. Kempton, DuMontelle, and Glass (1971) gave the following description of unit 4 Till in McLean County:

It is normally a reddish brown to pinkish gray till. Locally, this till grades to brownish gray. It can nearly always be differentiated from the upper three units in the region on the basis of color and clay mineral composition, particularly when one or more of the upper units are found in sequence with Unit 4. This till unit is slightly silty, averaging 30 per cent

⁶The location of Danvers is approximated by analysis site B5 on Figure 7.

sand, 40 per cent silt, and 30 per cent clay. The average clay mineral content is 8 per cent montmorillonite, 70 per cent illite, and 22 per cent chlorite plus kaolinite. No significant change in clay mineral composition or grain size of the unit occurs with changes in color. Unit 4 contains the greatest amount of carbonate of all five units.

Numerical data describing certain physical characteristics of the till matrix for the fifteen analysis sites on the Bloomington Moraine (Drift) are summarized in Table 6 and arranged in this table in order from east to west along the trend of the moraine (see Figure 7). Because previous workers have recognized more than one rock-stratigraphic unit within the Bloomington Moraine (Drift) (Willman and Frye, 1970; Kempton, DuMontelle, and Glass, 1971), it may be more revealing if the samples are divided into as many groups as is meaningful on the basis of available data. The result is the recognition of three distinct till units within this part of the study area. For purposes of this discussion these are named Units A, B, and C to avoid confusion with the Kempton, DuMontelle, and Glass number system. Unit A is defined on the basis of data for till samples from analysis sites B13, B14, B5, and B15 (east to west); Unit B by till samples from sites B4 and B3; and Unit C by till samples from all the remaining analysis site locations whose data is summarized in Table 6 (see Figure 7 for locations).

Comparison of the averages of Unit A data reported in Table 6 with the average numerical physical characteristic data of the Tiskilwa Till Member (Willman and Frye, 1970) and also Unit 4 (Kempton, DuMontelle, and Glass, 1971) reveals

Q
Table

Physical Characteristics of the Till Matrix--Bloomington Till and Associated Interlobate Till

_		_	_	_	_	_		-	_	-		-	_	-		-	-	-		_	•		
e Content	er Second	Dolomite	14	5 T	15	17	21	14	22	20	21	35	23	32	27	R	28	30	29	18			
Carbonat	Counts P	Calcite"	10	12	\$	7	\$	æ	\$	ۍ	11	26	11	9	18	24	15	T 6	19	S			calcite
mposition	%Chlor.	<pre>\$ Xaol.</pre>	13	13	15	12	17	10	14	13	15	24	23	16	17	14	14	15	24	13			averaging
Mineral Co	XILLITE		18	81	78	83	78	87	81	81	76	71	Ľ	67	67	72	66	6 8	れ	81		sted	urposes of
Clay	XEXP.	Clay	9	9	7	S	S	m	S	9	6	5	9	17	16	14	20	11	S	9	calcite	ione ext	o for p
ribution	% Clay		33	38	38	40	41	48	37	38	36	33	38	34	35	36	37	35	35	39	tence of	ite but n	ed`as zer
ize Dist	%Silt		42	42	40	36	35	36	43	40	40	38	34	46	38	37	37	40	36	39	le exis	for calc:	consider
Grain Si	%Sand		25	20	22	24	24	16	20	22	24	29	28	20	27	27	26	25	29	22	lestional	alyzed 1	oth are o
	Color		2.5Y5/4	2.5Y5/4	2.5Y5/4	2.5Y5/4	2.5Y5/4	2.5Y5/4	2.5Y5/4	2.5Y5/4	2.5Y5/4	2.5Y4/2	2.5Y4/2	10YR5/4	7. SYR5/4	10YR5/3	7.5YR5/4	Ante	Gray	Tan	nb : .	- ar	д
	Anal.	Site	BP8	BP11	BP10	BP7	B6	B12	B9	Bl	B2	B4	B3	B13	B14	BS	B15	Unit A	Unit B	unit q			
						ວ	7	ŢU	n			8.	นก	A	1	Ţ	'n	•	6/	٨			

a high degree of similarity. This is especially well shown when comparing the low illite content, high carbonate content, and pink color. On this basis it is reasonable to conclude that the surface till on the Bloomington Morphostratigraphic Unit to the west of the city of Bloomington belongs to the Tiskilwa Till Member of the Wedron Formation rock-stratigraphic unit. This correlation is commensurate with the mapping of this unit by Willman and Frye (1970), but does not agree with the tentative boundary drawn between Unit 2 and Unit 4 by Kempton, DuMontelle, and Glass (1971). The findings of this study would suggest that this tentative boundary of Kempton, DuMontelle, and Glass (1971) be placed near the Kickapoo Creek Valley and that the distribution of Unit 1 till be restricted to the proximal slope of the Bloomington Moraine (Drift).

Unit B, the second of the three groups under discussion, can be differentiated from Units A and C on the basis of color and clay mineral composition. This unit is violet gray when unoxidized and a gray brown or tan when oxidized. The clay mineral composition of this unit has less expandable clay than Unit A and less illite than Unit C, producing a greater percentage of chlorite and kaolinite than either of the other units. Less significant differences exist in both texture, Unit B a little sandier than both Units A and C, and carbonate content, slightly more calcite in Unit B than Unit A and a great deal more calcite in Unit B than Unit C (see Table 6). Willman and Frye (1970) included the area defined by analysis sites B3 and B4 in the Malden Till Member,

but findings of this study indicate that Unit B is clearly different from the Malden Till. Furthermore, Unit B cannot be correlated with Unit 4 Till, which was recognized and mapped in this area by Kempton, DuMontelle, and Glass (1971), because the differences between Unit B and Unit 4 Till are the same as the differences already described between Unit B and Unit A of this study. Thus, in the opinion of the author, Unit B should be treated as a separate and distinct till that has not been correctly correlated with any other rock-stratigraphic unit to this date. It should be pointed out, however, that the conclusion that Unit B Till has not been defined by previous investigators is based on the physical characteristic data from two analysis sites. Because data from only two sample sites may be judged insufficient for the purpose of naming a new till member, additional data were gathered from drill borings in McLean County. Data from drill borings McLean 10, 12, 13, 17, 18, 20, 21, and 25 support and extend the existence of this till unit. Figure 15 gives the location and elevation of these drill borings. At each of these drill-boring locations analyses of the clay fraction of split spoon samples reveal the existence of a till which is violet gray in color and has a clay mineral composition of approximately 7 per cent expandable clay, 70 per cent illite, and 23 per cent chlorite plus kaolinite. These data are on file in the geologic records section of the Illinois State Geological Survey in Urbana, Illinois. With the addition of this evidence the author

proposes that Unit B be hereafter referred to as the Shamrock Till Member of the Wedron Formation. The suggested name is for the town of Shamrock in McLean County. The type section is an exposure in a borrow pit (see Figure 8B) on the north side of I-74 1.5 miles (2.4 kilometers) west of Shamrock, SE SE SW, Sec. 25, T.23 N., R.2 E. The name Shamrock Till will be used in the remainder of this paper to refer to Unit B Till, and the type section will be described later in this chapter.

McLean No.	Location	Elevation
10	NW SE Sec.22 T.23 N. R.2 E.	785
12	SE SE NW Sec.26 T.23 N. R.2 E.	849
13	NE NW Sec.36 T.23 N. R.2 E.	803
17	NW SW NE Sec.33 T.24 N. R.2 E.	774
18	SW NW NW Sec.5 T.23 N. R.2 E.	730
20	SE NW Sec.5 T.22 N. R.3 E.	747
21	NW NW Sec.5 T.22 N. R.3 E.	784
25	SE NW Sec.5 T.22 N. R.3 E.	748

Figure 15 McLean County drill boring locations

Unit C Till has very different physical characteristics from either Unit A or Unit B Till. The most noticeable differences are (1) finer texture, (2) high illite content, and (3) fewer carbonates. The averages reported in Table 6 for Unit C are from sample sites within both the Bloomington Moraine (Drift) and the associated interlobate area. The data

in Table 6 indicates that the till samples from all analysis sites included in the Unit C group probably belongs to the same rock-stratigraphic unit. Comparisons of the numerical averages which represent the physical characteristics of Unit C Till described in the area shows that a strong similarity exists between Unit C Till and Malden Till. On the basis of this similarity it is probable that a correlation exists between Unit C and the Malden Till. This interpretation is consistent with the findings of Willman and Frye (1970). Kempton, DuMontelle, and Glass (1971) separated the Malden Till in the study area of this thesis into Units 1 and 2 mainly on the basis of texture. The results of the till-matrix investigations for this thesis do not provide a basis for the separation of the Malden Till into two units. There are some locally sandy places in the Unit C group, but, in the opinion of the author, the relatively small differences in texture do not provide an adequate basis for the recognition of more than one till member. Therefore, it is concluded that the till in Unit C is the equivalent of the Malden Till and is referred to as such in the remainder of this study.

Characteristics of Till Pebbles

Introduction

The purpose of this section is to present the data on long-axis pebble orientation and pebble lithology for 750 pebbles from analysis sites associated with Bloomington Moraine (Drift) (Willman and Frye, 1970). Orientation and

lithologic data are contained in the Appendix and are summarized in Figure 14 (in pocket) and Table 7 respectively. The discussion of these data will be directed toward a clearer understanding of the relationship between the sedimentary parameter and associated landform unit.

Till Pebble Orientation

Due to the complexity in morphostratigraphic and rock-stratigraphic units within the Bloomington Moraine (Drift), long-axis pebble orientation data will be discussed in association with Units A, B, and C as previously described because a definite relationship can be shown to exist. The fabric orientations of pebbles measured in Unit A Till (Tiskilwa Till) exhibit a consistent tendency toward a perpendicular alignment to the trend of the moraine (see Figure 14, in pocket). The strong north-south alignment of pebbles at analysis site B15 probably best represents the original depositional mode of Tiskilwa Till because the Malden Till does not overlie the proximal slope of the moraine in this area. However, it appears likely that the till at locations B5, B13, and B14 may have been deformed slightly by ice that eventually constructed the Eureka-Normal Moraine (Drift) because Kempton, DuMontelle, and Glass (1971) report Unit 1 Till on the proximal slope of the Bloomington Moraine (Drift) in this area.

The Congerville Section previously described in Chapter 3 shows an exposure of Tiskilwa Till stratigraphically below
the Malden Till. Pebble-orientation data from analysis site EB2 (see Figure 14, in pocket) reveal that a slight reorientation may have occurred in the ground moraine deposited in association with the Bloomington Moraine (Drift). The slight reorientation is illustrated by comparing the rose diagram associated with the pebble alignment at EB2 with the comparable diagram for B15 (see Figure 14, in pocket). A probable cause for this slight realignment is believed to be a result of subsequent ice movement that deposited the Malden Till.

Within the Bloomington Moraine (Drift) the Shamrock Till (Unit B), represented by analysis sites B3 and B4, can be separated topographically from the Tiskilwa Till on the west by the valley of the Sugar Creek, and from the Malden Till on the east by the valley of the Kickapoo Creek. These boundaries extend topographically to the south and incorporate parts of the Shirley and LeRoy Drifts (see Figures 4 and 13, in pocket). Data from drill boring McLean 17 (see Figure 15 for location) indicate that the Shamrock Till is stratigraphically overlain by the younger Malden Till and probably rests upon the older Tiskilwa Till in the subsurface. The long-axis pebble orientation data from Shamrock Till at analysis site B4 is the strongest orientation discovered in this study. Fifty-six per cent of the pebbles were oriented in a 30 degree sector around east-west and dip strength is to the east. Analysis site B3 from the same till also has an east-west fabric orientation with 28 per cent in the same east-west sector.

This strong east-west orientation of pebbles in Shamrock Till is clearly at variance to other nearby sites and suggests it formed in association with westward-moving ice. If this conclusion is correct, it is reasonable to suggest that the source of the Shamrock Till may well be from the Erie rather than the Lake Michigan Lobe. The Erie Lobe source for Shamrock Till is postulated under the assumption that certain pebbles tend to align themselves parallel to the direction of ice movement (Richter, 1932; Cailleux, 1938, and Lundquist, 1949), and supported by the fact that the Shamrock Till is unlike the drift in this area deposited by the Lake Michigan Lobe. However, because of the limited data and complexity of the drift, it is recommended that additional study be initiated regarding the Shamrock Till so that the conclusion of an Erie Lobe source for this till can be definitely confirmed or rejected.

Fabric orientations in Unit C (Malden Till) on the Bloomington Moraine (Drift) show a high degree of similarity to the fabrics of the Eureka-Normal Moraine (Drift) to the north, which has also been classified in this paper as Malden Till (see Chapter 3 and, Figure 14, in pocket). Strong perpendicular orientations to the trend of the moraine are found at analysis sites Bl, B6, and Bl2. Locations Bl and B6 are on the proximal side of the moraine and Bl2 is near the crest. Unimodal but weak fabric orientations were found at analysis sites B2 and B9 both near the crest of the moraine. Dip strength, an alternate measure of orientation, is to the

south at B9 and to the southeast at B2, giving both fabrics a weak preferred orthogonal alignment to the morainal trend.

The interlobate analysis sites on the Peoria Sublobe side of the interlobate section (see Figure 9) show a tendency toward a parallel alignment of pebble axis to trend of the topography. This tendency for pebbles to align themselves parallel to the landform is illustrated by the fabric diagrams at most BD and BP analysis sites as well as at the three E locations near the interlobate region (see Figure 14, in pocket) and is especially worthy of attention and explanation. Parallel alignment of pebble axis to landform trend is not unusual. The fabric of sediments associated with a drumlin often reveal such a relationship. Wright (1957), in his studies of drumlins in Minnesota, reports that almost all rose diagrams representing pebble alignment show a strong preference toward a parallel alignment with the landform trend. Glen, Donner, and West (1957) conclude that the mechanisms by which stones become oriented in till are most favorable to an alignment parallel to the direction of ice flow. Data from this study indicates that the long axis of pebbles prefer an orthogonal arrangement to the end moraine outside the interlobate area but changes to a tangental arrangement in the interlobate area. If Glen, Donner, and West (1957) are correct, then the pebble orientation data outside the interlobate area indicates ice flow perpendicular to the end moraine and the pebble orientation data in the interlobate area indicates that the ice flow associated with an interlobate

landform is moving parallel to the ice-contact zone. It is reasonable to assume that the pressures produced at the contact between two glacial ice lobes could force the ice and its associated sediments to flow parallel to the contact zone. Therefore, it is a conclusion of this study that the sedimentary parameter data of pebble orientation may reveal the nature of ice behavior in an interlobate moraine and that the till fabric in areas believed to be interlobate landforms could possibly be identified using the data of this parameter.

Till Pebble Lithology

Table 7 contains a summary of the lithologic classification of pebbles from all B and BP analysis sites. The original data can be found in the Appendix. A discussion of the limitations in the use of this data and a discussion on the lithologic categories used in this analysis can be found in Chapter 2. In this section an attempt will be made to analyze and compare the data in Table 7 with the data in Tables 2 and 5. Results for data on lithologic composition of pebbles in Tables 2, 5, and 7 are not consistent. Crystalline percentages are highly variable between analysis sites. The same is true of noncarbonate clastic percentages. The only patterns that have become apparent in the study of pebble lithology for all units in the study area are these: (1) the percentage of carbonate pebbles for the entire interlobate area and the Champaign Moraine (Drift) is consistently low; and (2) the percentage of chert and flint pebbles in the Table 7

Pebble Lithology--Bloomington Till and Associated Interlobate Till

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entire interlobate area and Champaign Moraine (Drift) is slightly higher than the comparable percentages for Eureka-Normal Moraine (Drift) pebbles and nearly double the percentage of chert and flint pebbles found in the Bloomington Moraine (Drift) exclusive of its associated interlobate analysis sites.

A comparison of lithologic composition between all interlobate analysis site locations (see Figure 9) reveals very little variation. The same is true in a comparison of the lithologic composition of pebbles from the C analysis sites with pebbles from both the BP and BD locations. Based on the data provided by analysis of the lithologic composition of elongate pebbles, it appears that the interlobate area of this study has a higher degree of similarity to the Champaign Moraine (Drift) (Snider Till) than to any other unit. However, the till-matrix composition data and pebble orientation data from the till within the interlobate area indicate the following: (1) The till found in the interlobate area associated with the Bloomington Moraine (Drift) (Malden Till) and the Champaign Moraine (Drift) (Snider Till) are very similar (see Tables 2 and 5); and (2) the pebble orientations within the till of the interlobate area indicate that the Malden and Snider Tills were probably deposited contemporaneously. Therefore, based on all the appropriate data, it is suggested that the landform defined in this study as the interlobate area, which was defined by Willman and Frye (1970) as Bloomington Drift, be named the Champaign-Bloomington Interlobate Moraine (Drift).

Shamrock Section

The Shamrock Section is the type section for a new suggested till member defined in this study. Analysis site B4 on Figure 7 marks the location of this section, and Figure 8B is a photo of the exposure. A description of this section follows.

Shamrock Section

Measured in a borrow pit in SE SE SW Sec. 25, T.23 N., R.2 E., McLean County, Illinois, 1971, 1972.

> Thickness (ft)

Pleistocene Series Wisconsinan Stage Woodfordian Substage Richland Loess

 Loess; thin layer of tan to tan brown to black loess, locally coarse textured, leached modern surface soil grades into unit one below
 2.5

Shamrock Till Member (type section)

1. Till, tan to gray brown in upper part where oxidized to violet gray where unoxidized in lower part, compact angular blocky, silty till zones with very few pebbles irregularly spaced in profile
14.5

Total 17.0

The Shamrock Till physical characteristics are (1) a grain-size distribution of 30 per cent sand, 38 per cent silt, and 32 per cent clay; (2) a clay mineral composition of 7 per cent expandable clays, 70 per cent illite, and 23 per cent chlorite and kaolinite; and (3) a carbonate content of 18 counts per second calcite and 30 counts per second dolomite. Drill boring McLean 17 (NW SW NE Sec. 33, T.24 N., R.2 E.) to the north of the type section shows the stratigraphic position of the Shamrock Till as between the older Tiskilwa Till and the younger Malden Till. Kempton, DuMontelle, and Glass (1971) map the area defined in this study to have Shamrock Till as Unit 4 Till, which is correlated with the Tiskilwa Till. The Shamrock Till can be differentiated from the Tiskilwa and Unit 4 Till on the basis of color and clay mineral composition. The Shamrock Till is a tan or yellow tan till when oxidized and a gray or violet gray when unoxidized with the zone of oxidation at approximately 13 feet (4 meters) in the type section. The Tiskilwa Till has a salmon or peach color when oxidized and is pink to red brown where unoxidized. An easily discernible difference in clay mineral content separates these two units. The Shamrock Till has an average of 10 per cent less expandable clay than the Tiskilwa Till, and this difference is nearly balanced by the 8 per cent greater chlorite plus kaolinite content found in the Shamrock Till. In the Kempton, DuMontelle, and Glass (1971) paper these variables were not considered. Only illite content was considered, which is approximately the same for both the Shamrock and Unit 4 Tills. These differences clearly establish the necessity for defining a new till member. The till fabric data in the area of this type section confirm the need for a new unit. The long-axis pebble orientation data of this study indicate that the Shamrock Till could possibly have come from an Erie Lobe source because the fabric has a

strong east-west orientation in the type section. The topography to the south of this type section indicates the possibility of a north-south trending landform which, if it did contain Shamrock Till, would strengthen the Erie Lobe source hypothesis. Analysis of the till and till fabric to the south of the Shamrock Till of this study is suggested for future investigation. The significance of the Shamrock Section is that it represents a significant exposure of till from the Bloomington Moraine (Drift) not previously defined.

Conclusions

The major conclusions of this chapter are the following: (1) The Bloomington Moraine (Drift) has three distinct surface tills in the study area: the Tiskilwa, Malden, and Shamrock Till Members. The latter was named and defined in this chapter. (2) Pebble-orientation data indicate that elongate pebbles prefer a perpendicular alignment to the moraine trend outside of the interlobate area but a parallel alignment to the landform is preferred in the interlobate area. And (3) the data of this study indicate that the interlobate area should be named the Champaign-Bloomington Interlobate Moraine (Drift).

Chapter 6

CONCLUSIONS AND IMPLICATIONS

Introduction

The purpose of this thesis is to establish the relationships that exist between certain sedimentary parameters of glacial till and particular glacial landforms. The sedimentary parameters investigated were long-axis pebble orientation, pebble lithology, and matrix composition. The glacial landforms studied were end moraines which represent recognized morphostratigraphic units. Two of the end moraines investigated merged into an interlobate moraine. Figure 16 summarizes the relationships between morphostratigraphic and rock-stratigraphic units established by this study. Also illustrated on this map is direction of ice flow as indicated by pebble-orientation data. Implications of the specific findings of this study have general application in geomorphology.

General Implications

The morphostratigraphic units of this study represent one interpretation of the glacial landforms in central Illinois that have been delineated mainly on the basis of topographic expression (see Figures 4 and 13, in pocket). Findings of this study indicate that morphostratigraphic units clearly exist, but a reevaluation of their forms and sedimentary associations may make them even more useful and meaningful as geomorphic



Figure 16 Relationships between morphostratigraphic units, rock stratigraphic units, and ice flow in the study area

units. Such reassessments were a major objective of this study, and the following are examples of findings or problems that have been recognized as a result: (1) the reinterpretation of the Eureka-Normal relationship as discussed in Chapter 3, (2) the variation in till associated with the Bloomington Moraine (Drift) which calls for a reassessment of that morphostratigraphic unit as discussed in Chapter 5 and to be discussed later in this chapter, and (3) the complexity and interrelationship within the interlobate area between the Bloomington and Champaign Moraines (Drifts) that indicate it was formed in association with an Erie-Lake Michigan Lobe ice contact as discussed in Chapters 4 and 5.

It is also important to recognize that certain problems are associated with any attempt to show relationships between morphostratigraphic units and sedimentary parameters. If two or more morphostratigraphic units merge, it may be impossible to understand their relationships without sedimentary parameter data and even that information may be insufficient for a reasonable interpretation. This problem is just one of the many difficulties that will present itself in studies that attempt to relate morphostratigraphy with associated sediments. However, the morphostratigraphic unit can be a very useful and meaningful unit when carefully mapped. The Champaign Moraine (Drift) of this study is a good example. This unit is easily delineated by its topographic expression. The surface till throughout its extent is a part of the same rock-stratigraphic unit. And the fabric complements its shape and depositional

mode. This situation demonstrates that it is possible to define morphostratigraphic units which have positive relationships to their sediments and related fabrics.

Evaluation of the relationships between morphostratigraphic and rock-stratigraphic units, which is one of the stated purposes of this thesis, has revealed evidence that supports a reinterpretation of the morphostratigraphy of central Illinois. A proposed alternate interpretation of the morphostratigraphy in central Illinois is a major conclusion of this study and is discussed below.

An Alternate Interpretation of the

Morphostratigraphy in Central Illinois

Figure 16 illustrates the relationship between morphostratigraphic units and the defined or correlated till members of this study. Ice-flow direction, which is suggested by the preferred long-axis orientation of each till member, is indicated with arrows on this map. A congruent relationship exists between the Eureka-Normal Moraine (Drift) and the Malden Till and also between the Champaign Moraine (Drift) and the Snider Till. Pebble-orientation data indicate that ice flow reflects a direct relationship to these end moraines. However, the relationships between the morphostratigraphic unit, rock-stratigraphic units, and ice flow for the Bloomington Moraine (Drift) do not complement one another.

Previous workers have recognized this situation but have not correlated rock stratigraphy with morphostratigraphy in

this area. Any reinterpretation of the glacial geomorphology associated with the Bloomington Moraine (Drift) must account for three separate and distinct rock-stratigraphic units within a feature that has been mapped as one morphostratigraphic unit and also account for an ice source region that will explain its fabric. An analysis of the topography of the Bloomington Moraine (Drift) indicates that it is not necessarily one morphostratigraphic unit. Instead, it may represent a situation where three morphostratigraphic units of different age converge on one another. This is substantiated by the existence of three distinct rock-stratigraphic units.

The data of this thesis and the topography said to be associated with the present east-west trending LeRoy and Shirley Moraines (Drifts) indicate that these units could equally as well be explained by north-south trending moraines which extend to join the Heyworth Moraine (Drift) and portions of the Shelbyville Moraine (Drift) (see Figures 4 and 13, in pocket). By interpolating between sections of land outlined by the 750 foot (225 meter) contour, two or three north-south trending landform units take shape which intersect the Bloomington Moraine (Drift) at nearly right angles (see Figure 17).

This alternate interpretation of the morphostratigraphy can account for variations in the rock stratigraphy of the Bloomington Moraine (Drift) by assigning the western Tiskilwa Till to an east-west trending moraine, the eastern Malden Till



Figure 17 An alternate interpretation of the morphostratigraphy of Central Illinois

to a relatively younger unconformable moraine more closely associated with the Eureka-Normal Unit which contains this same rock stratigraphic till member, and the Shamrock Till to a previously unrecognized north-south trending moraine named in this paper the Randolph. These interpretations are supported by the preferred long-axis orientation of pebbles within each respective till and are consistent with topographic expression. The suggested name Randolph for the westernmost north-south trending unit (see Figure 17) is for the town of the same name in McLean County, Illinois, which is near the crest of the postulated moraine (see Figure 13, in pocket).

The rock-stratigraphic correlation of the Shamrock Till with the tills of the Shirley Drift, which are defined as part of the Delavan Till Member of the Wedron Formation (Willman and Frye, 1970), has been suggested by H. D. Glass.⁷ Because the study area of this thesis was restricted to the boundaries of the morainal units under consideration, this suggested correlation of Shamrock Till and Delavan Till was not confirmed. It is suggested for future investigation that the tills of the LeRoy, Shirley, and Heyworth Drifts be studied and analyzed in an attempt to establish their relationship to the Shamrock Till.

⁷H. D. Glass, Clay Mineralogist of the Illinois State Geologic Survey, in oral communication suggested this possible correlation.

Conclusions on the Glacial Geomorphology

of Central Illinois

The morphostratigraphy, rock stratigraphy, and pebbleorientation data examined in this study suggest the following glacial chronology for central Illinois: (1) An ice advance, possibly even pre-Woodfordian from the north, deposited the Tiskilwa Till as a surface till on the Bloomington Moraine (Drift) to the west of the city of Bloomington, and this same till was deposited over a wide area because it is an extensive subsurface unit in the remainder of the study area. (2) An ice advance from the east or east-northeast overrode the Tiskilwa Till and deposited the Shamrock Till of the Randolph Moraine (Drift), with a north-south trending drift border. And (3) an ice advance from the north-northeast contemporaneous with an ice advance from the east-northeast deposited the Malden Till of the Bloomington Moraine (Drift) and Snider Till of the Champaign Moraine (Drift) respectively, forming an interlobate moraine named in this study the Champaign-Bloomington Interlobate Moraine (Drift) in the area of ice contact. The ice marginal retreat of the glacier to the north-northeast was marked by a slight readvance to form the Eureka-Normal Moraine (Drift).

Suggestions for Future Investigations

In a number of places in this thesis where data were limited or where there was doubt about findings and interpretations, recommendations were made for further investigation. The following is a summary of the suggestions for future

investigation of the central Illinois region as indicated by the results of this study. (1) Of primary importance for further study is an analysis of the relationships between topography, till-matrix composition, and till-fabric orientation to the south of the Bloomington Moraine (Drift) of this study. An alternate interpretation of the morphostratigraphy of this region has been proposed. Additional investigation may provide evidence to support this proposed alternate interpretation. (2) The high degree of correlation between the morphostratigraphic unit termed the Eureka-Normal Moraine (Drift) and the rock-stratigraphic unit called the Malden Till Member has been established by this study. This relationship also exists between the Champaign Morphostratigraphic Unit and the Snider Till Member. However, some of the relationships between morphostratigraphy and rock stratigraphy along the Bloomington Moraine (Drift) remain unclear. For example, the lines of demarcation between till members are questionable and there is a need for further investigations of the till and till fabric of the Bloomington Moraine (Drift) to establish clearly the boundaries of the rock-stratigraphic units. And (3) future studies of interlobate moraines might well include an analysis of the till fabric in order to confirm or reject the hypothesis of this thesis that the long axis of pebbles embedded in the till of these landforms tends to prefer a parallel alignment to the associated ice-contact zone.

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APPENDIX

Explanation of Data in the Appendix

Depth of Analysis -	In feet below the crest of the exposure
Type of Exposure -	See Figure 8
Topography -	See Figure 10
TILL PEBBLE DATA:	
Dimensions -	Longest dimension first, shortest dimension last; all dimensions in millimeters
Sha pe –	<pre>(See Figure 12) 1. Tabular-elongate 2. Tabular 3. Rhombohedroid 4. Rhombohedroid-elongate 5. Wedge-form-rounded 6. Wedge-form-elongate 7. Wedge-form-angular, subangular 8. Ovoid 9. Ovoid-elongate</pre>
Strike -	Orientation of the pebble's long axis
Dip -	Angle and direction of the pe bbl e's dip
Misc	S - pebble shows glacial striations W - pebble was deeply weathered LAV - long axis was vertical, there- fore strike data is for second longest axis
TILL MATRIX COMPOSITION DAY	FA :
Grain-Size Distribution -	Sand .062-2.0 mm. Silt .004062 mm. Clay less than .004 mm.
Carbonates -	? the existence is qu estio nable - none exists

Analysis Site Number Bl Norphostratigraphic Unit Bloomington

Location NE NE NE Sec. 25, T.33N., R.3E. Quadrangle LeRoy

Elevation 860' Depth of Analysis _____ Type of Exposure Roadcut (old)

Topography C and D slopes on proximal side of moraine

H11 P	ebble	Data:						_			_
No.	D	imensi	ons	Shape	Str	ike	Di	p	Li	thology	Misc.
	35	24	15	2	N25	h'n	0		[volor	nite	W
<u>z</u>	20	13	<u> </u>	5	1125		405		Dole	lite	
3	13	/	5	6	120)Ľ	15.	1	Dolo	uite	
4	21	16	14	5	1	1	20;	<u>.</u>	1:010:	site	W
5	17	12		E		1	59	;	P.150.	.t	1
6	15	H		1	1165		150	<u>_W_</u>	<u>Chile</u>	?	
7	16	10	ġ	3	NP4C	00	4	<u></u>	1:010	nite	
8	12	10	4	1	<u>N2:</u>	W	30%	!	isasa.	lt	4
9	16	15	10	6	1.5	<u> </u>	0		1.24601	tone	
10	16	6	<u> </u>	6	1 101	of;	5.	W	Cran	<u>tę</u>	
11	26	24	3	2	1120	<u>)(-</u>	55	11	Shale	<u>.</u>	+
12	15	10	10	5	NSC	J.N	1 10		<u>rotor</u>	nite	
13	28	91.	9	6	1 1161)F;	40	W	<u> </u>	lt	
14	19	15	3	2		<u> </u>	-0		Shale	` <u>````````````````````````````````````</u>	· · · · · · · · · · · · · · · · · · ·
15	23	16	10	6	N 50)F;	109	W_	Limor	<u>nite</u>	+
16	23	17	5	<u> </u>	N 31	11	1-10	<u>,</u>	L'olar	ate	+
	18	12	<u> </u>	6	+		<u>↓ </u>		Linor	ute	+
18	28	21	14	6	1.4	<u></u>	1 3 55	W	Lime	rone	
19	19	10	<u> </u>	6	+	<u> </u>	- <u>0</u>		Dolo	ni te	+
20	18	13			+ NCS	Wie	259	SE_	Lime	tone	+
- 21	18		8	2		of.	<u> </u>		Limor	iite	
	18	15	10	<u>_</u>	120		0	~ .	5110	tone	
- 34	22	12	<u> </u>		N2:	<u>)</u>	10	51	Basa.	1 C	ł
- 24	12	$-\frac{10}{00}$			+ <u>N1</u>	<u>.</u>	201	N	Lines	tone	
25	45			<u>-</u>	- NSV	/	- 25		<u>L1mo</u>	tone	+
		-10-		<u> </u>	NZC	<u></u>	0		Shall		+
- 20	35	- 21	20	<u> </u>	1 .15%	·	-15	1	Limes	tone	W
20	10	-13-1	<u>b</u>	<u>-</u>					Limes	tone	
29	$\frac{10}{10}$	<u> </u>				<u>w</u>		~ .	Red	imestone	+
- 30	10				H- 110		- 25	SW_	Dolor	lite	·····
-10-	- 20-		10		- N.S.				Dolor	nite	
- 32	1/	10	10	4		·	-10	2	0010	uite	w
33	18	- 15	8		H-100	1	<u> </u>	,	Lime:	tone	·
- 29	29	- 28	20	8	N/1	12.	- 304	<u>v</u>	0010	nire	
- 10		18	10	<u> </u>		N N	10	2	<u> </u>	te	
30	-15	- 8		6	- NI		103	<u>-W</u>	0010	nice	+
- 37	- 12		10		100		- 30	<u></u>		2100	
30	$\frac{10}{10}$	- 20	10		<u> </u>	<u>/₩</u>		<u></u>	<u>1010</u>	ille	↓
1 20						<u>.</u>		<u></u>	KCO C	<u>ancer e</u>	+
				<u> </u>	1 N 70	.E			<u>Limes</u>	to	t
42					- N/2) <u>r</u>		ů –		nito	+
1		12	14	<u> </u>	1 N 70		- 10	217		hite	1.34
44	10	16	15	<u>></u>	1 N20		401	VE	1.010"		+
45	15	12				<u></u>		16		nito	t
46	22	- <u>-</u>		<u>_</u>	N70		40	N		tone	1
47	23	12						·			<u> </u>
48	27	12		<u> </u>		1	100	1	Cb.1/	<u></u>	t
49	12	19		£	N DC		· •			tone	t
50	27	20	21	<u> </u>	N 70		600		iolor	nite	<u>+</u>
							<u> </u>	<u> </u>	0010		L
TILL	Matrix	Compos	sition	Data: Co.	lor Yel	llow t	an Mu	unse osit	ll Soll C	olor No. 2.	5Y 5/4
Gra	In-SIz	e Dist	ributio	n Exput	udable	1	^ -	C	hlorite	Carbonates	Cts./Sec
Sa	nd	Silt	Clay		lay	111	ite		plus	Calcite	Dolomite
2	2	40	38	Mine	erals			Ка	olinite	?	20
					6	8	1		13		

Figure 18 Data for analysis site Bl

Analysis Site Number _____ Morphostratigraphic Unit <u>Bloomington</u>____

Location NE NE SE Sec. 28, T.23N, R.3E. Quadrangle LeRoy

Elevation 820' Depth of Analysis 8' Type of Exposure Roadcut (old)

Topography _D and E slopes near top center of moraine

.

FILL P	ebble	Data:						*****		
No.	D	- Mimensi (ons	Shape	Stri	ke D:	10		thology	Misc.
	21	T 16	14	5	N50W	15	tw.	Dole	mite	+
2	15	6	4 1	4	1704	35	CE	Dolo	mite	1
3	23	22	10	2	NIOE	30)E	Cher	t	
	26	1 12	10	6	N15E	65)N	Silt	stone	
⊢ <u>}</u> ∣	23_	20			E NI OE	10)N	Dolo	nite	T
	22	<u> </u>	<u> </u>	<u>5</u>	N20W)	Gree	nstone	
	- 22-		┥╌┶╧┥	, <u> </u>	N/5L	10	<u>10</u>	0010	mite	
	1-12-1		<u> </u>	·	N ZOL		<u>is</u>	B-15-1	10	+
10		+ + + - +	<u> </u>		N254) 		niite	+
11	37	29	18		NECW		NN N		SLONE	+
12	23	18	12 1		NSOE	- 25	<u>.</u>	10010	mite	+
13	24	20	11	2	NSSW	40	NIN I	Dolo	mite	+
14	20	15	14	7	NPSE	15		Cher	r	+
15	22	14	11	3	N/SE	45	11	LO10	mite	+
16	14	11	8	2	NEOE	0)	Dolo	mite	1
17	12	10	E	5	NUSE	5	W	Basa	Jt	1
18	21		15	8	N/SE	15	W.	Lime	stone	1
19	12	8		6	N 3QE	70	W	Dolo	mite	T
20	20	14	10	5	N75E	0)	Silt	stone	Τ
21	22	1-15-1	9	2	N45E	40	NE	Dolo	mite	T
┝╬┽┥	ا_قد_ا	14	<u></u>	S	NSSW	20	W 30	Lime	stone	Ι
	22	1-14-1			N703	10	ISE .	Shal	P	Ι
	1-12-1	-21	<u> </u>	6	<u> </u>	25	<u>_33</u>	<u>Dolo</u>	<u>mite</u>	I
					NSUL	<u></u>	NE	Basa	<u>]t</u>	
+ + +					NI DW			0010	mite	
56-1		┝──╤┊╾╋					ATE -	Dolo	· · · · · · · · · · · · · · · · · · ·	+
54	+ + + +		;+		NASW		- INL	0010	mite	+
1 56 1					NVSF		<u> </u>			
1 11	- <u>16</u> +				N N	- ~~	~	0010	mile	
121	1 3 0 1	24			NOSE		N	Dolo	mita	+
1 33 1	48	121			N SOF	÷-+-ž5	NE		mitta	<u> </u>
34 1	18				N85W		11-1	Gran	aliorite -	ł
35	15		101	5 1	NS 5W		กังป	Dolo	mite	ł
36	10	8		<u>ş</u>	N45E	- 40	NE	Line	stone	
37 1	131	10	- 51	<u>5</u>	NGSW		50		mite	+
38	19	15			NSE		<u> </u>	Limo	nite	+
39]	12	91	5	5	T OE		,	5 11E	stone	+
40	10	8	5	5	N701		<u> </u>	Lime	stone	+ +
41	9	7	5	2	NBOE		5	Dolo	mite	t
42	17	13	10	S	NSSE	. 10	NE	Shal	c	1
43	9	6	5	6	NSOW	10	ISE	Dolc	mite	t
44	11	9	5	2	NUSE	. 0		Gran	ite	
45	13	91	7	9	N75W	5	5E	Shal	C	
46	12	8	8		N 5 5W	0		Polo	nite	
47 +	16	12	5	1	N20E	45	SE	Shal	e	
48	13	- 9 +	- 5+	6	N	10	IS	D010	inite	
49 +	-15	<u>10</u>	\$+	$\frac{1}{2}$	NIUW	20	<u>JE</u>	Dolo	nite	
50	<u> </u>	15	<u> </u>	<u> </u>	N SW	· 20	JE 1	Shai	e	
			*							
TILI r	ACT1X	Compos	ition u	ata: Corc	Jr Yello	<u>w tan</u> Mu	inse.	11 So11 C	olor No. 2.:	SY 5/4
-					ly Miner	al Compo	sit	Ion		
Grai	in-Size	2 Distr	ibution	Expand	lable		a	hlorite	Carbonates	: Cts./Sec
San	<u>×</u>	<u>5110</u>	Clay		¹ Y	Illite		plus	Calcite	Dolomite
L	<u></u>	40 1	00	Miner	ials		Ka	olinite		21
				³	,	/6				

Figure 19 Data for analysis site B2

Analysis Site Number B3 Morphostratigraphic Unit Bloomington

Location SE NE SE Sec. 21, T.23N., R.2E. Quadrangle LeRoy

Elevation 850' Depth of Analysis 25' Type of Exposure Roadcut (new)

Topography C and D slopes on distal side of moraine

	Di	mensi	ons	Shape	Strike	Dip	Lithology	Hisc
1	15	12	8	5	N85W	0	Dolomite	
2	22	21	13	5	£	111	Limestone	1
3	26	_17_	16	7	1800	20	Dolomite	LAV
4	19	15	12	3	NSOW	25E	Dolomite	
5]2	10	8	3	N3SE	0	Limestone	
6	10	6	1	2	1:40W	1008	Chale	
7	23	20	12	2	N 30E	100	Dolomite	
8	35	32	16	2	1i 3CW	401.1:	011	
9	10	5	4	6	N 35W	30111	Dolomite	
0	13	8	6	4	NBSE	35N	Polonite	
1	15	11	8	6	NGON	Ő	unert	
2	22	16	10	2	NSON	406E	Muscovite Gneis:	5
3	17		11	5	N15#	20:5	bolouite	
4	21	16	14	6	N SOW	15:W	Limestone	
5	15	13	3	1	N2OE	302	Shale	
6	13	10	8	5	£	103	Dolonite	
7	13	8	1	1	N40E	205'W	Shale	
8	27	18	18	5	1/55E	3584	Quartzite	
9	11	10	4	1	1155E	55'IN	Dolomite	I
0	16	8	6	6	NIOW	2 0S	Dolomite	
1	19	15	7	1	E	45N	Shale	
2	17	12	6	6	E	30.1	Shale	I
3	11	8	5	6	N SW	30N	Quartzite	
4	13	9	4	1	NGOW	0	Limestone	1
5	12	10	7	5	NJOC	30E	Basalt	
6	18	15	9	5	N85E	U	Basalt	
7	12	ઇ	5	6	N	0	Dolomite	1
8]	16	10	7	6	1409	2511W	Dolomite	1
9	19	13	11	7	N40W	0	Chert	
0	16	11	5	1	N4 5W	50.SW	Siltstone	W
1]	13	10	3	1	NSUW	90	Shale	LAV
2	23	17	8].	N4 5W	65NE	Dolomite	1
3	14	10	6	6	NSOW	0	Dolomite	1
4	28	21	11	1	NGSE	305W	Dolomite	
5	20	12	10	6	N20W	90	Dolomite	LAV
6	10	8	7	3	N85W	104	Dolomite	1
7	20	17	13	5	NBOE	101	Dolomite	T
8	13	8	4	6	N	90	Limestone	1-27
9]	19	15	10	3	E	90	Dolomite	LAV
o I	16	15	7	2	N1 5W	60.1	Dolomite	1
1	26	16	16	5	N15E	Ú	1,imestone	T
2	23	20	10	1	NSOE	45:W	Dolomite	I
3	40	16	12	6	NJUE	35S	Dolomite	Γ
4	17	12	9	5	N25E	20: W	Limestone	1
5	9	6	5	9	NISE	155	Dolomite	
5	15	12	8	3	N4CW	0	Dolomite	L
7	15	13	5	1	N40L	30NE	Shale	Τ
3	15	9	5	6	N35E	105W	Dolomite	T
9	23	18	10	1	NOON	40E	Dolomite	
D	23	14	14	7	N85W	45W	Basalt	
II	latrix	Compos	ition	Data: Colo	r <u>Gray</u> brow	vn Munse Composit	ll Soil Color No. 2	.59 4/2
Srat	n-Size	Distr	ibutio	n Expand	able	10	hlorite Carbonates	5 Cts.75
Sar	d S	110]	Clay	C1a	у 1113	lte 📔	plus Calcite	Dolomi
70		34	311		->- 1		the second second second	

Figure 20 Data for analysis site B3

Analysis Site Number ______ Morphostratigraphic Unit ______Bloomington

Location SE SE SW Sec. 25, T.23N., R.2E. Quadrangle LeRoy

Elevation 830' Depth of Analysis 20' Type of Exposure Borrow Pit

Topography B, C, and D slopes on distal side of moraine

ο.	D	imensio	ons	Shape	Strike	Dip	Lit	thology	Misc
1	31	24	20	5	E	35E	Dolo	nite	W
2	17	9	7	6	1.201	. W	Dolo	mite	
3	15	$\frac{10}{32}$	8	3	REOL	1 20W	Shale	<u>e</u>	+
-	14	12	8		N20E	90	10101	mite	1 10V
2		<u> </u>			<u>N857</u>	1-25E		nii te	
			↓		11904	- 90			1-1-1VA
.	-13	<u> </u>		<u>_</u>	NUL C	100	that	e	1.01/
÷ i	12			<u> </u>			- india		1.44
5			$-\frac{2}{1}$	0	11/04	CLN	Ch 1	<u></u>	+
ii I	8	6		<u> </u>	N751	40	0010	mite	LAV
2	18	12				40	Coll		1.AV
13	11	10	7	ς	11504	1- <u>70</u>	UDIO	nite	1
1	15	$\frac{-1}{12}$	10	7	NZOE	90	Chert	t	LAV
15	10	8	5	6	1.705	101:	lolo	nite	
16	20	19	12	2	N859	451	Gran	ite	1
17	16	15	9	2	NUSE	90	10101	nite	LIAV
8	15	13	7	2	N70E	20	Lise	tone	LAV
9	8	4	3	£	NUDE	1 20E	Lolor	mite	I
20	10	7	7	5	1406	4000	Gatb	ro	
21	14	10	10	5	NGOD	0	11210:	nite	
22	10	6	4	6	NSSE	<u>105</u>	Line	stone	1
23	9	7	5	7	E	55N	Chert	t	
4	8	5	3	6	NEOE	35E	Quar	tzite	
25	12	8		<u> </u>	NSE	45N	Line	stone	L
(t	9	6	4	<u></u>	NGO:	90	Limes	stone	t
		4		\	N85E	+	- Shall	e	t
(B	10		<u></u>		NUOE	+	110101	nice	·
2	- 14	- <u>15</u>			NOTE NOTE	- DUNA		nite	↓
				<u>+</u>	NUM	1 200	. <u> </u>	nice	+
2	┝╬┥	- ĉ			N745	+ <u>-</u>		HICE TO TO	L.W
12						1 100	Chor		+
	24	-12-1	-12		NADA				+
5		- <u>-</u>			N 201	1-40-	1.110	stone	1-120
6	24	21			NHOE	1 90		mite	TAV
57	11			4	N75W	1 0	Lime	stone	+
38	13				E	400	Dolor	nite	1
9	-īī -		- č t	5	Ē	50N	Tolor	tite	1
0	13	<u>-11</u>	4	2	NBOR	55N	Dolor	mite	<u>† </u>
1	19	14	7	1	NEOE	90	Cher	t	LAV
2	10	7	4	6	N85E	0	Cher	t	T
3	19	16	9	1	E	55N	Poler	mite	<u> </u>
4	14	10	9	5	NESW	90	Polor	nite	LAV
5	14	13	3	1	NUSE	90	Polor	ni te	LAV
6	12	8	71	7	NEOE	50NE	lolor	nite	
17	34	27	15	1	NSOE	0	Shill	e	W
18	48	26	16	6	NEOE	150m	Basa.	1t	
9	31	24	14	2	N 70W	90	Quar	tzite	LAV
50	14	13	7	5	NERM	35W	Gabbi	0	I
<u>11</u> -1	Matrix	Compos	sition 1	Data: Col	or <u>Violet</u> ay Mineral	Mun Compos	sell Soil Co	olor No2	.5Y 4/2
Gra	In-Size	e Distr	ibution	1 Expan	dable	T	Chlorite	Carbonates	s Cts./5
0.0	and I Silt Clay					1	te plus Calcite		
Sa	nd S	2112	Clay		ay [11]	lite	pius	Calcite	Lio10m1

Figure 21 Data for analysis site B4

 Analysis Site Number B5
 Norphostratigraphic Unit Bloomington

 Location NE NE NW Sec. 24, T.24N., R.1W.
 Quadrangle Danvers

 Elevation 820' Depth of Analysis 9' Type of Exposure Roadcut (new)

 Topography B and C slopes near top center of moraine

fill Pebble Data: Lithology Shape Strike Dip Kisc. No Dimensions 80 Quartzite 45 3 N20E 0 Т 53 N45E TUSW DOLOMICE 10 4 6 6 15E Shale 16 11 7 6 N85W 26 21 14 5 N80E Ō Colomite 45NW 2] 14 10 NGSW Limestone 6 2 N654 25SE Limestone 6 ננ 8 3 70CM 15 7 Limestone 24 11 N30E 25E Gabbro 8 24 8 1 NZOH 18 251 Dolomite G 35 N45L 8 5 6 10 13 N70.4 25SE Chert 11 8 7 255 Dolomite 11 18 NJ OW 11 9 12 NSSW 10!M 4 55 47 Basalt 94 W NGSE 25E $\frac{13}{14}$ Dolomite 50 33 20 6 1012 Diorite 19 12 12 4 N/UN Dolomite 20 12 10 6 H45E 15NE IUNE 16 21 13 9 6 NEOE Limestone 17 18 16 NYOL 501: Shale 3 2 7 N401 30:1E Linestone 17 13 19 25 7 0 Diorite Ŵ 18 15 E 20 20F Chert 39 6 N80E 84 45 Linestone NSON 25N 21 14 6 24 17 22 Dolonite N85W Ó 9 22 11 6 105W 23 2 Dolomite 17 14 7 N205 105 24 Dolomite N 35V 15 12 3 1 25 90 LAV Dolomite <u>21</u> 13 $\frac{15}{14}$ \$ NION WISE 20 155W Gabbro W 9 7 21 4 5 N156 υ Limestone W 28 28 16 13 6 35S N Dolomite 29 18 NBOE 40S 12 5 1 Shale 30 24 20 11 22 255W Dolomite N50E 31 10 19 18 N45W 0 Dolomite 32 14 9 105 6 N204 5 <u>Dolomi</u>te 33 43 26 13 N25E <u>405</u> Dolomite 19 10 24 34 W 10 9 6 N30D 0 Limestone 35 N30E 10NE 6 6 Limestone 36 18 16 5 N15E 0 Granite W 37 12 11 8 5 E 10£ Limestone 38 39 13 9 7 6 N20E 20NE Shale 12 7 6 6 N85E Dolomite 0 40 45 32 18 25N 20NE 1 N5E Limestone 41 16 10 8 N45E 6 Limestone 42 9 20SE 18 N70N 13 Shile 15 43 9 9 5 N1 SW <u>30:</u> <u>Polomite</u> 44 15 31 11 7 6 N70W 100E Chert 45 18 28 5 N35E 25NC Granite 46 20 19 16 8 1185W 10E Limestone 2 47 5 14 1 N 30N 20SE Shale 48 40S 15 13 9 NGOE Gabbro 19 28 12 18 6 NSW 30N Limestone 50 26 23 8 N 0 Shale Till Matrix Composition Data: Color Peach Munsell Soil Color No. 10YR 5/3 Clay Mineral Composition Grain-Size Distribution Expandable Chlorite Carbonates Cts./Sec Clay Calcite | Colomite Sand Silt Clay Illite plus 27 37 36 24 34 Minerals Kaolinite

Figure 22 Data for analysis site B5

72

14

14

Analysis Site Number <u>B6</u> Morphostratigraphic Unit <u>Bloomington</u>

Location NE SW Sw Sec.20, T.23N., R.6E. Quadrangle Arrowsmith

Elevation 830' Depth of Analysis 5' Type of Exposure Fipeline Trench

Topography C and D slopes on proximal side of moraine

No.	D	imensio	ons	Shape	Stri	ke D:	ip	Li	thology	Misc.
1	36	29	20	3	NUSU	V 9	0	Silt	stone	W
7	14	9	7	<u>ل</u>	145.	V	0	Bas	110	1
	14	10	5	9	N454	V	511.1	Quin	rizite	
4	13	7	5	6	N404	<u> </u>	SGW	1,1:14	stone	-
~~~	13		-10		11451	;;	000		stone	LAV
- <u></u>	20	10	10	<u> </u>	<u>NSOP</u>			Cher	Nito.	
8	16	10	14		NG 5V	<u></u>	56	Dolr	mite	
ğ	16	10	6	7	N75	;	<u>us</u>	Dold	omite	+
10	16	13	8	7	140		0SW	Line	stone	
11	20	15	9	3	NEOM	v 3	EN	Silt	stone	
12	33	11	6	5	N 5 04	v 1	5NV	Shal	6	
13	10	7	7	3	NH 51	v . 1	0	Quar	rtzite	
14	14	7	5	6	N 351	: 3	5:5W	Lime	stone	
15	18	14	3	1	N\$5W	/1	5NW	Sha	le	
16	18	12	2	5	11414	/	0	1010	omite	
17	27	18	17	7	N504	1	<u>ose</u> .	1010	unite	W
18	-29-	-17	4		N504	<u></u>	0	Chal	<u>le</u>	
20					N/04	<u> </u>	w +		scone	- +
21					N 4 1	;	200			-+
22		-12-		6	<u> </u>	·	<u>~~</u> +			
23	13	10	6		NS 59	1 2	0.ml	Sha	A	
24	24	15	15	3	N502	$\frac{1}{1}$	ONN	Ling	stone	Ŵ
25	15	11	6	5	N	2	51: 1	Lim	stone	
26	15	10	7 1	6	N40V	v 1	0	QUIU	rtzite	
27	8	7	4	7	N 5 5V	/ 2	5E	Dolo	omite	T
28	20	19	5	2	NJ.55	1 5	SSW	Dol0	omite	
29	23	_15_		5	NGOL		SEWL	Dold	omite	
30	25	_16	15	7	<u>NSE</u>	1	OSE	0010	omite	W
31	<u> </u>	7			1130/	<u></u>	9↓	<u> </u>	le	
32	- 2		-4+		<u>N60</u> v	<u></u>	<u>e</u> +	<u>Şha</u> ,	le	
33		- 12			<u>         NSOR</u>	<u> </u>	2-+	0010		
- 15		- 12-			N8 54		20-	Dolo		
36		-10	<u> </u>		N25v		004	Choi	of the second se	
37	- <u>*</u>			6	N506	;	<u>~</u> +	Dolo	mite	+
38	11	6	6	6	N4 54	1 2	SW	Sha	le	
39	18	11	10	3	NGO	1 5	OSE	Dolo	omite	
40	15	12	8	7	N40v	1 2	ONV	Limo	stone	
41	18	12	8	6	N40E	5	0	Chei	rt	
42	15	10	8	7	N70F	- 4	0:W	Dolo	omite	W
43	18	16	11	5	N15E	33	OSE	10010	mite	
44	24	20	14	3	N40	/ 3	SSW _	Lime	estone	·
45	14	12	10	88	<u>N50</u>	!	0	Sha.	le	
46	10		2	2	<u>N20</u>	<u> </u>	ONEL	Shall	le	
10	- 9			<u> </u>	<u> </u>		55	<u>boi</u> d	mite	
10					<u> </u>	<u> </u>	<u>on</u> q_	Quar	ctzite	_ <b>_</b>
30	10						V.MW			· †
						l {	للبائعان	0101		
FII	Matrix	Compos	ition D	ata: Col	or Yell	low tan Mu	unsel	I Soll C	olor No.	2.5Y 5/4
				CT	ay Mine	ral Compo	ositi	on		
Gra	In-Size	e Distr	ibution	Expan	lable		Ch	lorite	Carbonate	s Cts./Sec
Sa	nd	SILE	Clay		зу	Illite		plus	Calcite	Dolomite
2	4	35	41	Mine:	rals		Kao	linite	?	21
				1 5	1	70		17 1		

Figure 23 Data for analysis site B6

Analysis Site Number <u>B9</u> Morphostratigraphic Unit <u>Bloomington</u> Location <u>SW SE SW Sec. 27, T.23N., R.4E.</u> Quadrangle <u>Arrowsmith</u> •••••

Elevation 890' Depth of Analysis 5' Type of Exposure Roadcut (new)

Topography C and D slopes near top center of moraine

o.	D	imensi	ons	Shape	Strike	Dip	Lithology	Misc.
1	24	19	7	1	N35E	90	Polomite	
Z	13	9	15	6	N502	15.7	Chert	
3	12	6	1.6	6	NASW	2000	Basalt	
4	15_	2	2		11453		Shule	LAV
<del>}</del>		<u> </u>	$10^{-10}$	<u> </u>	N-IQL	2024	Limestone	
<u><u> </u></u>	18	<u> </u>	1-10	6	NUTA		llint	
<u>~</u>	18	12	<u> </u>	<u> </u>	1624	1 - 00 -	Dolomite	
8	- 12-		<u> </u>	<u> </u>		1_1	Uolonite	
~			<del>                                     </del>				Linestone	
1	16	<u> </u>		<u> </u>	- Magni		Quartzite	
5	18		<b>├</b>	<u> </u>		1000	Unionite	-+
	12		<del></del>	6	HASE	10.1	basar c	
Í	- 22	17	12	<u> </u>		$\frac{10}{10}$	liolonite	
5	-13-		<del>τζ.</del>	6	N350	1500	Dolonite	
ř.	- 9	- 7	6	Š	NISTA	1005	Lipestone	
7	14	10	6		NUSE	4500	(martzite	
Ŕ	11	8	6	6	NGOW	10:14	Pasalt	
9	13	10	6	7	N65W	255	Dolonite	
0	47	35	19	7	1451	35 11		
1	14	12	7	3	NSOE	0	Dolomite	
2	20	15	10	5	N85E	15E	Chert	
3	14	10	7	3	1:40W	20	Bisalt	LAV
4	14	8	6	6	MSE	30%	Basalt	
5	33	10	5	6	1140W	15SE	Basalt	
6	9	8	5	7	N75W	251	Shale	
7	17	12	5	2	N65E	25:1	Linestone	
8	- 13	8	4	6	NSOW	205E	Limonite	
9	19	14	9	7	N25W	1500	Dolomite	
0	12	8	5	6	N40E	20NE	Granodiorite	
1	11	7	S	6	E	157	Chert	
2	11		4	1	N20E	253	Dolomite	
3	21	76	12	7	<u>N65W</u>	358	Dolomite	
4	16	10	9	4	NJOE	205	Dolomite	
5	13	9	4	1	N2CM	155	Dolomite	
6	20	13	12	6	NBOE	40.W	Dolonite	
7	11	7	4	6	N 301:	35NI	Polomite	
8	16	12	3	5	<u>1170a</u>	450	Dolomite	
9	12	10	7	5	<u>N50E</u>	155W	Dolomite	
0			6	6	NHSW	20.1	l'olomite	
1	10		7	3	<u>N351:</u>	<u></u> 20	Dolomite	
2	13	10	-2	1	<u>N450</u>	30.2	Shile	
3	13			<u> </u>	NSSE	1.	Limestone	
4	-18	13	10		<u>N 30C</u>	65	Diorite	
2	7	32	- <u>,</u>	8	N 758	4011	Dolomite	
윽니	- <u>2</u> 0		1/	<u>ь</u>	N /OE		Sandstone	_ <del></del> ₩
<u>/</u>		2			N/5W	+ 0 	<u>Polomite</u>	-+
<u> </u>		-10	<u> </u>		NBSE	1_20E	<u>Shale</u>	+
			<u> </u>		NOUL	1 <u>105</u> M-	Dolomite	+
<u> </u>	20	12		0	NZUE	N	Dolomite	
117	Matrix	Compos	sition	Data: Col	or Yellow t	an Muns	ell Soil Color No.	LSY 5/4
				[7]	ay Mineral	Composi	tion	
Gra	in-Siz	e Disti	ributio	n Expan	dable		Chlorite Carbonate	es Cts./Se
Sat	Sand Silt Clay				ay 🛛 Ill	lite	plus Calcite	Dolomit

Figure 24 Data for analysis site B9

Analysis Site Number <u>B12</u> Morphostratigraphic Unit <u>Bloomington</u> Location <u>NE NE SW Sec. 33, T.23N., R.5E.</u> Quadrangle <u>Arrowsmith</u>

Elevation 880' Depth of Analysis _ 7' Type of Exposure Roadcut (old)

Topography B and C slopes near top center of moraine

<u>No.</u> ]	D	imensi		<b>M</b>		1				
]			JII.3	Snape	Strike	D1	P	L	chology	Misc.
	127	88	51	6	NGOW	10	hW	1)0] 0	mite	S
2	20	15	9	7	N70L	- 25		1.010	mite	
	15	8			<u>N75E</u>	30	<u>4</u>	Quar	tzite	
		10-				1-20	<u>u</u>		<u>e</u>	
6	27	21		1	N20F	60	<u>р</u>	<u> </u>		
7	21	15	7	1	N705	30	C I	Shal	e	
8	19	16	13	8	NSUM	20	22	Lime	stone	1
9	16	12	7	3	NZON	1.15	N	1010	nite	
10	20	15_	5	1	NIOC	90		Shal	c	
$\frac{11}{10}$	11	- 25 -	- 22	6	NIOE	90		<u>l,imo</u>	nite	
12	12	8	<u> </u>	<u> </u>	<u>N15C</u>		<del></del>	<u> </u>	mite	-+
-12-	-22-	-12-	<u> </u>		NIOC	1 12	N		slone	
15	10	11			<u> </u>	1-10		1,010	ni te	
16	14	10	n	ς	N100	20		1:010	mite	
17	20	15	10	7	1259	20	Na	Dolo	mite	-1
78	22	14	12	7	N259	0		Polo	mite	
19	13	10	7	5	N2CW	30	N	1010	mite	
20	13	8	3	1	N15W			Basa	لت	
21	45	28	25	6	<u>N20C</u>	50	<u>s</u> _	<u> </u>	ite	W
-22	30	16	10	G	<u>N304</u>			Lime	stone	
-23	- 37	- 22	-17-	6		+0		Lime	stone	
-25	<u></u>	- 24-		<u> </u>		+ <u></u>	<u>i</u>		scone	
26	16	12	10	B	HCOU	1			F =	
27	30	18	12		N201	1 0	-+-		mite	<b></b>
28	15	14	5		N25E	- 30	NE	Shal	e	
29	35	25	21	5	N25W	25	SET	Quar	tzite	
30	32	- 25	5	2	N4QE	ة ت	NE L	Shal	c	
-31	10	8		8	N5W	10	N_ I_	Lime	stone	·
- 32	18		10	6	N	.0		_1010	mire	h
-33	36					- 20	S <u>E.  </u>	<u></u>	nite	W
- 24	$-\frac{19}{10}$			<del>-</del>	N85W	190		010		
-36		- 10				1-10	<u></u>	Dion	<u>i</u> to	· •
37	12	- 10			N22W	1 10	NF -	Lime	stone	
38	25	17	15	3	N25W	30	···	Dolo	mite	LAV-W
39	11	7	6	4	N25E	l o	-	Lolo	mite	
40	16	15	4	2	N35E	L O		Shal	C	
41	14	9	88	6	2011	0		Silt	stone	W
42		6	5_	6	N 70E	10	w	Colc	mite	
43	10	7	<u> </u>	<u> </u>	N75W	<u> </u>	-1-3		mite	
44					<u> </u>	$-\frac{20}{30}$	N L		stone	-
-73-	- 20	-19+			<u>N805</u>	1-10	<u>~</u>	B,15,3		
47	15				10#		N. 1	<u> </u>	ito	
48	10	8	- 5	2	N75F				mite	-
49	23	17	7	i	NIOW	20	N	Shal	e	
50	10	7	5	6	NSOE	20	SW	Ouar	tzite	W
TIII	Matrix	Compos	ition	Data: Colo	or Yellow t	an Mu	nsell	Soll C	color No2	5Y 5/4
Gra	In-Size	Distr	Ibutio		iy Mineral Table I	Compo	Sitio	orite	Carbonato	s Cts.75ecl
Sa	nd	SILE	Clay		ıy [™]   111	lite		lus	Calcite	l'olomite
				-1	I ==		× 1	1 · · .		*
10	5	36 1	48	Minei	rais		1901	inite [	8	1 14

Figure 25 Data for analysis site B12

Analysis Site Number B13 Morphostratigraphic Unit Bloomington

Location _____ NW_ NE_Sec.14, T.23N., R.1E. Quadrangle _____ McLean

Elevation 790' Depth of Analysis 8' Type of Exposure Roadcut (old)

Topography C and D slopes on distal side of moraine

D.	0	imensi	ons	Shape	Strike	Di		Lithology	Misc
1	23	15	9	6	NIW	1 201	Siltsto	one	
2	27	20	14	S	NJ SW	60N	Basalt		
3	22	14	7	6	N4CW	301	Dolomit	:e	
4	23		13	3	N85W	10:1	Limeste	one	
5	_ 16	9	8	6	NEW	90	Limesto	ve	LAV
6	13	9	4	6	IISE .	90	Limesto	one	LAV
7	]4	11	4	1	N7CW	90	Limesto	one	LAV
8	19	16	12	7	NSSE	301	Lineste	one	W
9	20	]4	13	5	N75E	90	Siltste	one	W
0	12	10	9	3	E	90	Dolomit	:e	
<u>.</u>	45	35	25	5	NIOE	45N	Limeste	one	
2	19	17	6	1	NIOC	90	Shale		LAV
3	16	12	5	2	NGSW	90	Shale		LAV
4	20	12	9	6	NSOF	357	1 Dolomit	e	T
5	19	115	14	7	NSOE	2511	: [iltsto	ne	
6	16	12	1	6	N.LOM	4011	Cabbro		
7	12	10	5	2	NGSE	CON	Chert		1
8	31	27	22	7	HEOE	0	Polomit	:c	
9	30	20	16	5	N260	501	Chert		
0	57	31	4		111.5W	251	Limeste	one	
1		20	18	5	NISW	45N	Limesto	one	
2	_23	18	6	2	N208	20	Siltste	one	
3	_26	18	4	1	NZOE	1 40N	limestr	one	
4	33	18	10	1	NN	1.521	Shale		
5	28	11	3		N30E	_45N	Shile		
6	17	24	10	5	E	1 n	Siltete	ne	F
7	15	8	5	5.	NASE	_ 20	Polonit	ç	LAV
8	14	9	6	7	1155E	30:1	Uolomit	.c	
9	11	10	5	8	N705	4011	Dolomit		T
0	18	15	10	7	NHOE	0	Chert		
1	18	7	4	6	1150%	101	V Shale		
2	13	11	9	7	N85C	90	Basalt		LAV
3	37	23	16	6	NGOE	1 30C	Limesto	one	
4	22	14	4	1	N35E	90	Shale		1
5	21	15	13	3	N5E	35N	Dolomit	:e	<b></b>
6	18	10	_ 4	1	W10W	90	Limesto	one	LAV
7	24	18	10	7	N40N	45N	/ Dolomit	e	<b>—</b> ———
8	14	13	3	2	N85W	10E	Basalt		
9	25	19	3	1	N70E	90	Dolomit	.c	LAV
0	40	15	4	1	N85W	0	Shale		
1	49	19	3	1	N20E	101	Shale		I
2	20	16	9	5	NIOE	20N	Granite	Gneiss	TW
3	28	23	13	5	N85W	90	Dolomit	C	I JAV
4	15	10	10	5	11512	20N	Polomit	.e	T
5	24	15	10 I	4	N25E	101	Dolomit	c	· · · · · ·
6	20	13	11	5	N75E	201	l'olomit	e	<b>T</b>
7	25	20	14	7	N	90	Limonit	C	I I.AV
8	12		8	5	NEOW	90	Polomit	:c	T
9	14	13	5	1	N25E	0	Dolomit	e	r
0	17	15	7	1	N201:	15N	Dolomit	e	1
II I Grai	Matrix	Compos e Distr	ition D	ata: Col	or <u>Peach</u> ay Mineral dable 1	Mur Compos	ition Chlorite	lor No. 10	NR 5/4
Sar	nd	SILE I	Clav		av   111	ite	plus	Calcite	Doloni
2	0	46	34	Mine	rals		Kaolinite	6	32
			77						1 34

Figure 26 Data for analysis site B13
Analysis	Site Number	r <u>B14</u> Mo	orphostratig	raphic Unit .	Bloomington	
Location	NE NE NE	Sec.5, T.2	3N., R.1E.	Quadra	ngle <u>McLean</u>	
Elevation	800' De	pth of Anal	lysis <u>6'</u>	Type of Exp	osure Roadcut (old)	

Topography B and C slopes on distal side of moraine

-	l .	1 mar - 1 -		<b>Ch</b>	Conduc	n4 -	1.	(the) cm	
<u>10.</u>		1mens10	ons	Snape	Strike			1thology	Misc.
	13	12			112 Ju	2504	L1	nescone	_
		10		4		- 25		nestone	
		12-			<u> </u>	1 15:1			
		70					<u> </u>	110	
- 6	20-	24		<u> </u>	<u> </u>	25:1		lomite	
- 7	- 28	27	- 12-1		<u> </u>	- 2202	- <b>-</b>	nestone	
-	20	12				<u></u>		IORITE	+
ŏ		<u> </u>		<u>4</u>	1904	4212		116	+
10					N205	400			
11	12			<u>0</u>	NZUL	405		ses cone	LI TRU
12	22		10			- 20			
12	- 22-			<u> </u>			100		
14				<u> </u>	NL 05				
15	10				MG (M)		00	Lomice	
16		10			N755	90		artzite	
17	12	10			11751				-+
18	25				NIV ( AJ	1 30		10,110;	1 AV
19	12	12		5	N201			110	1.10
20				1	NGCH			nes cone	
21	10								
22	17	10			LICOL	00			
23	70	- 10	26		NOEU				-fhay
24	15	14				100			
25	28	12	6		11355	4555			
26	18	12	10						-
27	14	10	7	6	1700	20			
28	$\frac{1}{12}$			<del>č</del>	NIGU	1 105	-+	nor tono	-+
29	15	12	Ŕ		NISW	25N		intrito	
30	11	10	6		NISSW	- 40-		lonite	TAV
31	14	12		2	NSSE	90			
32	11	4	8		N40C	90-	Do	lomite	TAV.
33	11	8	7	5	N60	- 90-		lomite	- LAV
34	12	11	10	5	NZOE	90	Ch	Prt	LAV
35	11	9	5	2	N20E	50	Sh	ile	
36	13	3	5	2	N20W	90	Do	lomite	LAV
37	10	8	3	2	NZOE	205	Do	lonite	+
38	13	9	7	7	NBCN	0	Do.	lomite	
39	23	21	7	1	HSOW	1557	L'O	lonite	
40	11	10	5	7	NGON	1057	Do.	lomite	
41	12	8	8	3	N15W	90	Do	lomite	LAV
42	9	8	5	5	N70E	1055	Die	orite	
43	16	13	7	5	NIOW	15:	Lir	nestone	
44	35	13	4	i	NSOE	805	Do.	lomite	1
45	35	20	12	2	1120W	355	Li	nonite	1
46	11	/ 1	3	1	NSOE	90	Che	ert	LV-
47	17	10	5	1	N25E	90-	1. h	11e	LAV
48	18	14	- 5 1	3	N50W	30NM	Do	lomite	
49	20	14	9	7	N50W	0	1)0	lomite	-t
50	25	18	10	7	N4UW	2011	1 10	lomite	<b>+</b>
111	Matrix	Compos	ition 1	lata: Colo	or Salmon	Mun	sell Sofl	Color No. 7	-SYR 574
					av Mineral	"Ompos	ition	1	
Gra	In-517	e Distr	ibution	Expand	lable	- <u></u>	Chlorite	Carbonate	e Cte TCA
5.	Sand   Silt   Clay   Clay   Illite   plus   Calcite				1 10 001				
~ ~ ~	7	19			nals		Kaolinite		
	<u> </u>	- 20	<u></u>					1 1 10	1 47

Figure 27 Data for analysis site B14

Analysis Site Number <u>B15</u> Morphostratigraphic Unit <u>Bloomington</u> Location <u>SE NW NE Sec.19, T.24N., R.1W.</u> Quadrangle <u>Mackinaw</u> Elevation <u>790'</u> Depth of Analysis <u>7'</u> Type of Exposure <u>Roadcut (old)</u>

Topography C and D slopes on distal side of moraine

					T		.			1
No.	D	imensi	ons	Shape	Strike	D	1p	L:	thology	Misc.
	45	24	19	6	N	25	N	Dol	omite	S
	24	+ + + +				-1-12		<u>- Sna</u>	le	
	10	1-12		<u> </u>	HOLE		<u></u>	<u>- 501</u>		
5	- <u>10</u> -			2	<u>10056</u>		<del>~</del>	563		
6	33	17	8	2	1150	30		Cho	rt	
7	14	1	7		11.55	- 20		Lin	nctiona	
8	13	10		2	1104	25	1	(.h.)		
9	13	10	7	5	11306	20	NE I	Dol	omite	
10	15	13	6	1	NGOW	15	1.1	Dol	Unite	
11	59	. 50	- 33	3	NZOL	60	N	Lin	estone	
12	9	6	5	3	N 5W	70	!!	Che	rt	
13	11	9	9	5	NSOE	35	112	Dol	omite	
14	24	15	7	2	NZON	20	NA	Base	nlt	S
15	2]	17	14	7	NSC	90		hoj	onite	
16	18	15	8	5	11204	50	11	<u>Sha</u>	le	
1/	24-	+ 19	-15	/	<u> </u>	-+	<u>∂</u>	<u></u>	et tone	
10	13	10	<u> </u>	<u> </u>	N15L	+	<b>∺ </b>	1001	omite	
- 19	-12	20		<u>}</u>	11101			1.0.1	DELLE	
21	<u>+ : : : : : : : : : : : : : : </u>		<u> </u>	<u> </u>	N' 4		14.4	<u></u>	ouito	
22	12		6.		141534	+				
23	14	4		4	NSC	+ 30	;; <del> </del>		uni te	
24	13	12	6		N				omate	
25	21	15	12	6	N15E	- 25	N	Pol	onite	-
25	11	10	5	5	1120.1	20		Qui	ntgita -	- T.
27	27	17	16	7	H2OE	10	ME T	001	omite	
28	14	6	4	4	N15E	115	11	Pol	Gnite	
29	24	18	11	5	N2ON	50	//	001	omite	
30	62	30	17	6	N20W	$\pm 25$	NW ]	Do.)	omite	
31	30	10	16	6	N 30W	15	NW		omite	
	37	25	15	6	1175W	30	<u>N</u>	Linu	onite	
33	50	35	- 25	5	114 SW	0			omite	
-74	- 25	18			N40W		NW	511	cscone	
35	22	10	$-\frac{14}{10}$		NEUL		NE	0101	omite	
-17	- 19	-18	-10		N	90			rtzite	
-14-	20	14			114 (14	+-;;;	CE	Sha	0	
39	20	- 11-	- 4		NSE	1-15	N +	$-\frac{3\pi a}{Sha}$		
40	15				N20E	4	N	Lim	estone	• • • • • • • • • • • • • • • • • • • •
41	17	15	10		NSE	50	N T	- Sil	tstone	+
42	20	16	10	5	NSOE	0		Sano	lstone	w
43	16	14	13	7	N5W	0		Dol	omite	
44	28	20	17	9	NTOE	10	3	Qua	rtzite	1
45	32	19	15	6	N40L	45	NE	Lim	onite	1
46	34	23	5	1	NION	10	<u>S</u>	5h.1	le	1
47	10	8	6	5	N705	10	5	Sil	tstone	
48	11	6	5	6	N15W	20	N	Sha.	le	
49	11	8	8	5	N 30W	10	ŞE	101	omite	+
50	_24_	16	13	5	N65W	0		lim	estone	_L
TIII	Matrix	Compos	ition I	Data: Colo	or <u>Salmon</u> ay Mineral	Mu	unsell psition	son (	olor No. 7	.5YR 5/4
Gra	in-Size	e Distr	ibution	Expan	dable [		Chlo	rite	Carbonate	s Cts./Sec]
Sat	nd 3	511t ]	Clay		ay   11	lite	pl	us	Calcite	l'olomite
<u> </u>	6	37 ]			rals		Kaoli	nite	15	26
				<u> </u>	0	66	14	·]		

Figure 28 Data for analysis site B15

Analysis Site Number <u>C1</u> Morphostratigraphic Unit <u>Champaign</u>

Location NW NE SW Sec.9, T.20N., R.7E. Quadrangle Mahomet

Elevation 780' Depth of Analysis 9' Type of Exposure Roadcut (new)

Topography B and C slopes on distal side of moraine

	 _	
	 <b>n</b> - 1	<b>.</b>
PDF		

No.		Dimensi	ons	Shape	St	rike	D	ip	1	1thology	Misc.
T	20	113	6	8	NS	0E	15	NW	Li	nestone	
7	21	1 15	14	5	1-115	0.1	1-0		<b>R</b> •	CALC	
3	29	22	16	3	I NT	50	20	ND.	50	lomite	1
4	16	T 11	9	5	115	5.4	25	SE	Li	nestone	1
5	20	1 12	10	6	1 112	010	10	C.I.	lio	lonite	5
6	28	20	5	j	115	50	15	: .1	5h	ile	1
7	1 17	13	3	1	NS.	511	20	5	('h	ort	1
8	] ] 4	11	8	8	116	5.1	30	5.4	Do	lo:ni te	1
9	18	13	8	5	N 5	OL:	55	9	Ga	puro	W
10	28	1 16	15	6	N2	1.11	15	NW	Li	testone	1
11	17	1 12	7	7	ID:	01	20		Do	lomite	
12	20	14	11	6	NE	51:	25		10	lomite	1
13	13	1 10	E	3	1:4	5E	30	N	Do	loaite	
14	34	22	16	5	N5	F:	60	<u>c</u>	Do	lonile	1
15	1 15	1 10	9	6	1 113	512	0		ijo	onite	
16	1 15-	113	8	₹	1.6	(1)	10	N	: 11	rt	+
17	24	20	12	ς	1 10	SE.	40	<del>;  </del>	('h	ort	1
18	110	114	10	8	NX NX	50	311	1		lomite	1
19	1 21	14		<u> </u>	112	 !	30		1,0	lomite	+
20	20	1 10	11		HIC NIC	(+,)				orite	
21	1 72	1 20						34	11.1		+
22	1 10			t		0.1	;;;	<del>au</del> t			· · · · · · · · · · · · · · · · · · ·
23		25					-20	;"- <b>-</b> -		lonite	4
24				·····			-10	<u></u>		1001Ce	·
25					<u>↓</u>	24	- 22	*			
- 25	- 22 -			<u> </u>	<u> </u>	20	-30	26-1		lomite	- <del>  </del>
20	<u> </u>	12	10	<u> </u>	12		- 20	<u> </u>	<u></u>	irtzite	
	-24-	20	<del>;</del>	2		5W		<u></u> _	Do	ert	
	-2/				!`			3	100		
- 29	22	20		8	N4	<u>0w</u>	-15	NE		Lomite	
- 30	14			4	<u> </u>	UW			Uo.	Lomite	
- 31	15	- 9	8	6	N.2	SW		IIW 1	<u> </u>	ert	
	17	14	7	8	<u>N7</u>	50			Qu	irtzite	
- 33	19	14	13	5	hit.	51.	30	SW 1	Ch	ert	
	12	8	6	6	14	55	0		Po		
35		13	7	5		50		SM	DO	lomite	
36	9	7	5	6	N4	30	35	<u>SW</u>	Qu	irtzite	
	26	19	14	5	11	5E	35	SW	Do	lomite	
	13	10	4	1	NE	512	0		Gr	anite	
39	21	16	14	5	112	5Ľ	20	SW	Do.	lomite	
40	14	10	5	5	N4	5W	- 30	SET	Ch	ert	1
41	42	18	17	4	NS	<i>il</i> 1	20	Ś	Do	lomite	
42	13	10	2	2		N T	0		Ch	ert	T
43	18	13	9	6	N5:	N	0		Ch	ert	T
44	21	12	8	6	N 3	5E	25	5W	Ba	salt	1
45	14	10	8	6		E I	0		to	lomite	1
46	14	9	6	7		E T	0		Ch	ort	1
47	13	9	6	5	N7	5 <del>.</del> 1	-10	E-T	Do	lorite	1
48	27	26	ŝ	2	N6	5 <u>C</u>	40	SET	Sh	ale	t
49	26	21		2	NG	58 1	40	set	Sh	ale	t
50	22	13		6	NG	OE I	30	NE.	Ch	ert	t
			<b>_</b>	×							
TII	Matrix	Compos	Ition 1	Data: Col	or Gr	een gr eral (	ay Hu	insel ositi	1 Soll	Color No. 2.1	5Y 4/2
Gra	in-Siz	e Distr	Ibution	Expon	dable			Ch	lorite	Carbonates	S Cts./Secl
Sai	nd	SILE T	Clay		Clay Illite plus			Calcite	Polonite		
26	5	38	36	Mine	rals			Kao	linite	11	23
					5	79	, , , , ,		-18	<b></b>	المستقسسا

Figure 29 Data for analysis site Cl

Analysis Site Number <u>C2</u> Morphostratigraphic Unit Champaign

Location NE NW NW Sec. 26, T.21N., R.6E. Quadrangle Gibson City

# Elevation ______ Depth of Analysis _____ Type of Exposure Railroadcut (old)

Topography B and C slopes near top center of moraine

<u> 111 P</u>	EDDIE	Daid:		<del>r</del>	T				
No		imensi	ons	Shane	Strike	n	10	Lithology	Mac.
+ <del></del>	23	1 17	7 7		NHOE		····	Volomite	+
<u>⊢</u>	20	+ 12-	1 12		1950	-		'cloute	+
3	22	13	1 B	4	NOSW	10	N	Limestone	
4	13	9	7	6	NGUV	10	S T	Dolonite	
5	17	15	10	5	N2CE	0		Polomite	
6	16	8	7	4	N70E	305	SW W	Colomite	1
7	16	15	7	5	NBOE	25	NE	Bisalt	1
8	14	10	5	5	N	35	N	Limestone	1
9	15	9	7	4	N75W	0ڌ	N I	Chert	
10	29	20	13	6	NSOE	40	ia -	Dolomite	S
11	37	26	23	5	114 CW	10	·:: [	Dolomite	
12	28	2.5	10	7	140E	503	EN	Basalt	
13	25	23	13	8	11552	1 10	je l	Dolonite	
14	20	12	10	6		Ú		Quartzite	
15	27	18	4	1	NSE	35:	5	chile	S
16	26	15	5		11750	20!		Chule	
17	]4	1)	3	1	NION	40:	1	Chule	
18	16	12			1355	30.	1	liorite	_
19	28	22	3	5	N/OC	151		Linestone	
20	10	1	4	8	N25W	0		untrite	
21	22		9	5	N20E	251	4	Lisonite	<u> </u>
22	19_	16	7	5	NESE		ــــــــــــــــــــــــــــــــــــــ	5hule	
23	17_		8	6	NGOE	- 251		Granite	
24	19_	6		5	1754			loionite	
25	19	15	14	1	N45W	20	1.M	Gibbro	
76		_24	12	6	N30M		1	lolorite	
		17			1601	- 20		Chert	
20	35	23			NGOE	- 50	23-1		
29	- 28 -	-19	<u> </u>		1166		· · · · · · · · · · · · · · · · · · ·	Shale	
- 30	-25	11	- <del>c</del>	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	14032				
31		10			NEDJ			hant	+
- 32					NESE	+-25	·····	lomite	
1 20	- 15	26	- 17-	<u> </u>	N950	+ 10		Gubbro	
- 75-	- 21	20	12	<del>- č</del> -	N755-	- 30	NE	Polomire	
36	30	20	10	6	N250	309	SE	Siltstone	
17	30	-57-			NZGE			BASATE	+
38	29	21-	-15	Ē	N75E	10	39	Polomite	
39	80	65	43	3	N	358	5	Dolomite	
40	36	31	13	1	NSW	20:	<u> </u>	Delomite	<u>s</u> -
41	16	10	6	4	1851	-1-15	2	Limestone	
42	16	11	10	ь	150.1	35	SE 1	Greenstone	
43	13	9	7	5	11500	303	59- <b>1</b>	Colomite	
44	26	17	15	5	L C			Lolomite	
45	16	15	10	7	N85W	20	₹ <u>1</u>	'olomite	1
46	41	- 26	25	6	N40W	0		Granite	1
47	36	23	22	6	1455	40	5	Basalt	T
48	]2	11	5	7	NGSE	10	NE	Chert	T
49	20	10	10	6	NSOE	0		Dolomite	
50	11	8	5	8	NSSE	0		Dolomite	I
TIII	Matrix	Compos	sition	Data: col	or Gray	Mu	insell So	11 Color No.	2.5Y 5/2
1						-			
					ay Minera.	1 Compo	osition		
Gra	in-Siz	e Disti	ributio	n Expan	dable _		Chlori	ce Carbonate	s Cts./Sec
	nd	5110	CLAY		ay I.	11126	plus	Calcite	Dolomite
	8	59	43	Mine				15	
1				L4	<u>ــــــــــــــــــــــــــــــــــــ</u>	82	15	l	

Figure 30 Data for analysis site C2

Analysis Site Number C3_ Morphostratigraphic Unit Champaign Location ______ NW NE NW Sec.9, T.22N., R.6E. Quadrangle Arrowsmith

Elevation 820' Depth of Analysis 5' Type of Exposure Roadcut (old)

Topography B and C slopes near top center of moraine

H11 P	HII Pebble Data:											
No		imenal	00.	Shane	Sta		1.		thology	Mino		
	17	1 27	7	Silape	- SUP.	70	*+		1 chorogy	ALSC.		
1-2	15	+ + + + +	<del></del>		1-75							
3	24	10	8	4	N 75	10		5h.	10			
4	10	7	5	5	11%E	0		1)0]	onite			
5	8	7	5	6	N20	1 0		Pol	omite			
6	20	12	9	4	1130	N 0		Has	alt			
7	27	17	3	1	N55	E 40	S	Sha	le			
8	17	10	7	6	115E	25	SE	(i1	tstone	W		
- 10-	17	111-	2	6	1170	1 25	SE	1001	omite			
- 10		+			1130	20	M.J	Sha	10			
12	20	110			1.65			1001	omite			
1	14	$+\frac{12}{11}$	10		- H45	<u></u>	<u></u>	15.15	110			
14	62	46	25	<u>_</u>	NEC		N	0.10	omite			
15	12	6	<u>1</u>	6	NAS		<u></u>	1 1 1 1	ostone			
16	14	9	7	4	NAO.	<del>j n</del>		101	mite	+		
17	17	12	8	6	1 165	1 35	NE	101	onite	+		
18	35	25	20	5	N85	1 55	1.	Lim	estone	1		
19	20	10	5	6	I NSU	10	NE	Lim	estone	1		
20	26	20	10	7	NIU	N 15	5	Lim	estone	1		
21	35	23	22	Ę	N		S	Gra	nite			
22	14	8	4	<u>1</u>	N65	/ 0		Sha	1e			
23	14	<u> </u>	10	8	N	0		Uol	omite			
24	16	8		4	NSW	10	<u>s</u> _	Lin	estone			
- 22-	11	- 6			N45	25	WW -	<u>Dol</u>	omite	·		
27	-26-	<u> </u>		<u>    4                                </u>	1180	<u>v 20</u>	····	<u> </u>	rtzite			
28	10	2		<u> </u>	N80	<u>, , , , , , , , , , , , , , , , , , , </u>		001	omite			
29	13			6	N85		NT -	Cne	alt	+		
30	31	19	8	ī	NSS	25	NC -	001	omite			
31	10	8	4	8	N 35	v to		Oua	rtzite			
32	10	9	4	8	N20	N 0		Dol	omite			
33	22	16	14	6	N45	N O		Che	rt			
34	15	12	6	7	E	10	W	Dol	omite			
35	15	10	2	1	NEO	N 15	s	Sha	Je			
36	_17	11	3	1	NSM	35	SE	Sha	1e			
3/	$\frac{16}{10}$	15	- 6	7	N704	<u>N 50</u>	<u>s</u>	Sil	tstone			
20			;+	<u> </u>	N60	0		Qua	rtzite			
		10			NBU	<u>v 50</u>	N	Sha	10			
1 <u>4</u> 1 1	-10				N40	<u></u>	<del>+-</del>	L1m	estone	+		
42	- 21	17	A	2	N15			<u> </u>	tstone	+		
43	20				NASI		¦a- <b> −</b>	100	omite	+		
44	18	14	2		N.55	30		Sha	le	+		
45	17	12	8	7	NSS	N 30	SE 1	Dol	omite	· • • • • • • • • • • • • • • • • • • •		
46	14	7	6	6	N	0		Che	rt	4		
47	11	7	4	6	E	25	HAV T	Lim	estone	1		
48	22	10	7	6	N10	N O		Bus	alt	S		
49	49	38	11	7	N75	N O		Che	rt			
50	61	36	24	4	N50	<b>4</b> 0	NE	Dol	omite	W		
TILL	Till Matrix Composition Data: Color Olive Munsell Soil Color No. 2.59 5/4											
Gra	In-Size	e Distr	ibution	T Expan	dable I		Chi	orite	Carbonate	s Cts. /Secl		
Sar	d Silt Clay Clay Illi		Illite	D	lus	Calcite	Dolonite					
11		33 56 Minerals Kaoli			inite	-	17					
	2 84 14											

Figure 31 Data for analysis site C3

Figure 32 Data for analysis site C4

8		26	19	6	<u>N4</u>	00	201	IE .	Lı	mestone	
9	14	12	2	2	NS	5E	<b>30</b> 3	;		ale	· ·
10	14	30	6	6	N2	01:	101		Ba	salt	
11	19	12	10	6	Nl	00	30N		Gr	anite	
12	17	]4	8	8	NC	0.1	0		110	lomite	
13	26	14	1	1	NS	5.1	651	IE.	Sh	lle	
14	11	H	6	6	NG	01:	150	,	Li	nestione	†
15	12	11	6	<u> </u>	115	J		·+	100	lonite	•
16		10	10	¢	- NO		200		10	louite	+
17					NO	<u>.</u>		<u> </u>	0	10mrce	<b>{</b>
	10	-+			<u> </u>	24	20	+		11Scone	
+ + + + + + + + + + + + + + + + + + + +	<u> </u>	14	<u> </u>		<u>N3</u>	<u>01.</u>	<u> </u>		<u>Cn</u>	ort	<u></u>
19	57	18	12	66	<u> </u>	3 <u>C</u>	10	W_	<u>no</u>	lomite	<u>8</u>
20	14		8	5	<u>N2</u>	5E	100	E	Vo	lomite	
21	14		6	5	NG	5E	0		<u>Po</u>	lomite	
22	17	12	8	8	NG	SW	103		<u>Do</u>	lomi <u>te</u>	
23	13		4	6		N	701	[]	I.i	aestone	
24	14	11	6	7	I N1	04	159		زيل	nestone	
25	15	12	6	7	N4	0E	15%	IE T	Do	lonite	11
26	17	7	3	1	NG	51	300		Sh	1]6	1
27	19	9	7	1	NE	<del>.</del>	- 4n	t			T LAV 1
28	16	12	6	8		nu	- 10	-+	<u></u>	mestone	LAV
29	18	14	11	5	N4		90	+	Do	lomite	
30	44	- 15	18	<del></del>	N4	01:	503		Ch	ort	
1-11-	19		-11-		MS	<u>nu</u>	- 60-	+	Cr	anito	
1-12-1			- 21-		- NO		-200			louite	
	- 37		- 10				705	<u>-</u> +			
- 33	-23	10	12		N4		302		<u>Cn</u>	ert	
- 34	- 20	12		4	N/	WC	305	·	11	nescone	
- 35		- 21			<u> </u>	UL			Qu	artzite	
56		20	18		NS	5W	- 10		D1	orite	W
57	16	10	10		N5	UW	101	W	00	lomite	
38	15	11		5	NI	SE	305		L1	mestone	
39	11	7	5	6	NE	SW	305	5	Qu	artzite	
40	15	8	6	4	<u>N5</u>	OE	201	IE	Ch	ert	
41	14	10	7	5	N8	5E	10.		Li	mestone	
42	20	16	8	5	N2	56	<u> </u>	: T	Do	lomite	S
43	22	13	8	6		N	0		Do	lomite	
44	11	8	5	5	N2	64	0		Do	lomite	
45	13	12	6	7	N5	5E 1	555		Bu	salt	5
46	12	10	8	7	NG	01:	100	ET	Do	lomite	
47	10	6	- 5	6	NG	SE I	0	t-	1.1	nestone	
48	19	10	<u> </u>	4	N2	58	300	w t	0	lomite	<u>s</u>
1 49	15	10	-10		NS	54	405		0	ert	
50	16		- <u> </u>		NB	<del>~</del>	-201		<u> </u>	lomite	
	10		<u> </u>		10		201	·		I OMI CG	
TIII	Till Matrix Composition Data: Color Gray Munsell Soil Color No. 2.57 5/2 Clay Mineral Composition										
Gra	in-Siz	e Distr	ibutio	n Expan	dable			Ch	lorite	Carbonates	Cts./Sec]
Sai	nd	Silt	Clay	ст	ay	111:	ite		plus	Calcite	<b>Polomite</b>
10	18 40 42 Minerals				Kao	linite	9	20			
				2		79	9		19		
1	•										
	****										

Topography B and C slopes on distal edge of moraine

Shape

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Fill Pebble Data:

25

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Dimensions

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No.

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6

7

Analysis Site Number ______ Morphostratigraphic Unit Champaign Location SW SW NW Sec.32, T.22N., R.GE. Quadrangle Arrowsmith Elevation 760' Depth of Analysis 15' Type of Exposure Railroadcut (old)

Strike

N5E

N55W

N75W

NG5W N50L

N75E

N50L

Dip

550 1017

555 3017**N** 30N

20A

0

Lithology

Dolomite

Polonite

Limestone

Dolomite

Flint

Shale

Greenstone

Misc.

W

Analysis Site Number <u>C6</u> Morphostratigraphic Unit <u>Champaign</u>

Location SE NE NW Sec.33, T.23N., R.6E. Quadrangle Arrowsmith

Elevation 800' Depth of Analysis 4' Type of Exposure Roadcut (old)

Topography B and C slopes on distal edge of moraine

<b>F111</b> P	ebble	Data:		r	T				T
No.	D	imensi	ons	Shape	Str	ike Di		ithology	Misc.
I	24	22	13	5	11851	E 59	I Do.	lomite	
2	19	17	10	5	N75	30	l Da	lanite	1
3	17	10	9	4	116.5	N 153	u Do	lomite	I
4	10	8	3	2	162	1 2	Chu	ort	
5	22	21	<u> </u>	5	170	N 10:	E Che	ert	+
6	17	13	8	3	1185	<u> </u>		int:	
<b>⊢</b>	26	14	10	6	1 140	1.15	E Do	lomite	4
⊢ <del>°</del>	13-	- 10	4	6		x 10.		nestone	+
10	12	10	1 10	·					+
1	$\frac{17}{17}$	14	11		116.6				+
12	17	15	+ <u>+ </u>		117	1 10			+
13	15	17	+	5	140		2 DO		+
14	32	20	17-	6	N 30	7 3(-)	• / • /		1
15	14	10	9	5	1150	1 200			+
16	18	10	9	6	160	v 1 10°	Bar Bar		1
17	17	8	5	6	1185	V 100	Sh.		
18	22	12	9	6	NAS	/ 100	E Lis	estoue	
19	13	12	7	1	I 11152	1 205	Chr	ort	
20	18	13	11	6	11751	E 10.	/ Gat	obro	
21	17	11	10	E	N8 51	v 45.	Lin	nestone	
22	22	10	7	6	N80.	<u>v 40:</u>	Sh.	. <u>le</u>	
23	22	14	11	6	N 755	1 5.	l no	lomite	
24	_27	20	8	1	N65	<u>v 10:</u>	ta Sha	1]6	
25	13	1.11	9	5	<u>N65</u>	<u>v 115</u>	E Bas	<u>alt</u>	
25	30	22	15	· · · · · ·	<u>N 30</u>	301	<u>E</u> Sha	<u>1e</u>	
	15-	10	5	5	1/5			0.120	
28	15	10		6	E		51	tstone	
29	20	-15	6	<u> </u>				tstone	·
- 30	-1/	- 10	4		N40	2 100	C Sne	110	+
- 12-	12			<u> </u>				tetono	
1-32	17	- 1/		6	M 3 5/	4 - 4 5 -			+
34		-11	5	6	175	$\frac{1}{10}$		tstone	
1 25	11	á	à		N601	10			<b>+</b>
36	14	- Á	8	7	N .01	$\frac{10}{10}$	W Che	rt	
37	14	10	8	5	N25	. 0	10	louite	+
38	36	32	15	2	1150	1 30:	E Do	lomite	<b>1</b>
39	15	12	10	5	N70	1 0	Do.	lomite	1
40	10	7	4	6	N75	N 35/	i Do	lomite	
41	25	13	10	6	N30	E 25:	V Gra	inite	W
42	16	]2	7	5	N65	N 10/	Do.	Conite	L
43	41	18	10	5	N75	N 304	/i	lestone	
44	16	12	10	6	N 75	/ 0	Qu.	irtzite	
45	16	12	2	)	N15	10:	sh.	le	
46	10	7	6	6	N80	N0	1)0.	lonite	
47	20	18	12	5	1160	N 201	W DO.	lonite	·
48	36	28	18	6	<u>N70</u>	$\frac{1}{10}$	<u> </u>	irtzite	+
49	13	11	<u> </u>	7	N70	L 15	Cho Cho	oru	<b> </b>
50	13	11	5	2	LH5C	150	) DO	LOMITE	1
TIII	Matrix	Compo	sition	Data: Col	or Gra	<u>y tan</u> Mu	insell Soil	Color No. 2	.5Y 5/4
1								<b>1</b>	
	7				ay Min	eral Compo	sition	C	- AL- 14
Gra	10-51z	e Dist	ributio		ande	7114	Chiorite	Carbonate	S CTS./Sec
l ⊢ sa		3110			ay	111116	prus	Calcite	DOTOWIL
1	<u> </u>		33		1.972		Addinite		
[				L		62	<u>_</u>	1	

Figure 33 Data for analysis site C6

Analysis Site Number El Norphostratigraphic Unit Eureka

Location NW NW Sec. 15, T.23N., R.5E. Quadrangle Arrowsmith

Elevation 910' Depth of Analysis 6' Type of Exposure Roadcut (old)

Topography B and C slopes near top center of moraine

<b>.</b>	D	imensic	ons	Shape	Strike	Dip	Lithe	ology	Misc.
1	60	27	26	6	N	105	Poloni	te	
2	29	11.	17	9	NIOE	400	Bicalt		
3	52	32	25	5	N30E	2053	Siltst	one	W
4	23	20	14	7	N25W	_205E_	Dolomi	uc	
5	25	18	10			2006	Siltst	one	L
6	15	24	15	9	NISM	1011	Limest	0116	W
7	24	13	13	6	N40E	20:11	<u>Pisalt</u>		
8	26	31	10	1	N.35W	2056	Siltst	0n <b>c</b>	+
9	16		8	<u>`</u>	N350	2050	Limest	ong	<b></b>
10	15	-11			<u>N458</u>	25.1	Chert		+
11	- 40		-10-	<u>0</u>			101011	<u>te</u>	+
4			- 15	<u> </u>	HJ5L		Lines C	one	+
	32				N451.		101001	10	+
-	- 44	29			N401.	15.11	101001	<u>re</u>	+
13	-24	- 20	- 22	<u> </u>	NTON N		holoui	50	+
3	- 24	$-\frac{32}{16}$	- 29	<u> </u>	1765		holomi		+
8	10	$-\frac{10}{12}$			NASU	30110	holomi	<u>te</u>	+
Ğ	24				11751	2011	Cuanit	<u></u>	<del>†</del>
6	62	- 30 -	30	<b>A</b>	NASE	5511	Dolomi	te	+
21	27	-ip-			NASE	0	Dolori	te	<u>+</u>
2	27	19		i	1605	40N	Shale	**	+
23	23	20	- F		N159	255	Chale		+
4	19	16	10	8	N201	355	Dolomi	te	
15	30	20	17	3	N SW	305	Dolomi	te	
(6	27	23	12		NZSW	0	Dolomi	te	
7	32	17	10	1	N	10N	Dolomi	te	1
26	14	10	5	5	NSON	2015	Doloni	te	<u> </u>
9	22	15	12	5	N20.	0	Dolomi	te	1 W
0	27	19	16	3	N 30.4	15N	Limest	one	•
1	45	28	25	9	NSSE	400W	Dolomi	te	T
52	38	35	20.	7	N 30W	355W	Dolomi	te	
33	44	25	21	6	N SW	205	Dolomi	CLE	
34	20	14	9	6	N	105	Siltst	one	
5	28	25	18	S	NGSE	0	Dolomi	te	<u> </u>
36	19	17	11	5	NGSE	90	Dolomi	te	LAV
37	17	12	- 4	2	NGOW	705	Dolomi	te	I
38	18	15	12	5	NION	_205	Dolomi	te	I
19	15	17	11	B	MIZM	105	<u>Uolomi</u>	tc	L
10		13	12	7	NI 0W	20N		one	
11	15	-11	5	2	NZOC	GON		one	
2	26	17			NGSE	40NN	Shule_		
3	_20_		12	3	N10W	40N	<u><u><u> </u></u></u>		L
4	38		21	3	<u></u>	<u>15N</u>		one	
15	50		26	3	N104	255	Dolomi	te	<b> </b>
ь	28	<u>1</u> 0		<u> </u>	NSW	305	Leloni	te	<b></b>
4	$\frac{17}{20}$	15	5	2	<u>N5W</u>	<u>505</u>	Shule		<b>+</b>
8	- 38	- 51			NIOE	<u>50S</u>	Dolomi	te	<b>+</b>
3	25			<u> </u>	<u>N15E</u>	455	Dolomi	.re	+
0	23	12	10	b	N/VW	0	Basalt		
11	Matrix	Compos	ition I	ata: Col	or <u>Tan</u>	Munse	11 Soll Cold	or No. 2	5Y 5/4
	7 AT		mar		ay Mineral	Composit	1on		
Gra	Frain-Size Distribution Expandable				Dable	4 m 1 0	Iniorite (	arponates	1 UTS./S
Sa	nd S		CLAY	- 1 ^{c1}	ay   111	176	pius	Calcite	Poloni
		7 L I	10	M100					

Figure 34 Data for analysis site El

Analysis Site Number E3____ Morphostratigraphic Unit _____Eureka___

Location NE NW NW Sec.28, T.25N., R.1E. Quadrangle Danvers

# Elevation 830' Depth of Analysis 5' Type of Exposure Roadcut (new)

.

Topography B and C slopes near top center of moraine

o.	D1	mensio	ns	Sha De	Strike	Di		ithology	Misc.
Ť I	151	11	5	1	Ň	51	Dolc	mite	
				5	7.55W	- 255	E DOIG	mite	1
3	28	25	8	1	N2OE	15:	I Limo	stone	
4	11	9		1	1:75W	501	Dolu	mite	W
Ś	16	- 11		6	1/201	15	W Dolo	nite	
6	10	17	- 121	5	USP	154	Dolo	mite	
				4	N SOUL	15,	SV Dolc	mite	
Ŕ	10	10		4	NRSE	10	Live	stone	
å	24	- <u>,</u>	6	4	NSE		f Shal	Ċ	
10	19	16	6	1	211 014	25	Dolc	znite	
<b>n</b>	11	7	4	7	12 11	1 0	Cher	rt	
12	16	14	7	5	N4 5W	20'	Lime	stone	
17	20	16	15	3	NSON	15	W Cher	·r	
11	18	14		3	N'90W	10	I Gran	ite	
15	12	- 7		6	NISE	15	M Chul	<u>с</u>	T
16	17	- 15		5	N	250	Biot	ite Gneiss	W
17	20			2	<u>r</u>	250	<u> </u>	stone	1
18	17	-12	10		11254	251		voite	1
19	15			7	NS 5'J		0010	mite	LAV
20	- 22			2	NASE	1 0		0	1
21	21				1/451	200		ani to	Ψ
22	16	-16-1		6	NIOF	-1-20	Linc	stone	
52	14				N25E	205	W Cher	rt	1
24	10	- 6			N35E	201	U Ou Ir	tzite	
25	28	-18			NISU	250	E Dolo	unite	
26	- 22	-12-1		4	NSE		I Gran	ite	
27	10			6	NGOW		Lirc	stone	
21	+++++++++++++++++++++++++++++++++++++++				11001			wite	
20	1 22	- 16			N75F	250		mite	
23	- 16				NIOF				
30		- 15 1			NLOC	- 200		stone	
31					NAOJ			stone	+
34					NU()E				
33	- 22	- 12 +			NACIJ	-1		stone	
34		- 57			140W			stone	
35					11 E.C.			mite	
36	+2				N75E			stone	
3/					N2 31			scone	-+
38					1140.		<u>n</u> 1010	trite	
39					.1356				
40					11000	<u>+</u> ;		mito	-+
41				<u>،                                     </u>				mite	-+
42	20	18			MI W	_ <del></del>		unite	- <del> </del> ^w
45					- t.	1.5%		2	
44	73			0	NC DW		<u>- 500.</u>	0	+
45	- 32		18	0	WUCH	-1-251	W DOIC	aite	+
46					NOC		E 1010	nice	+
47		14			NEUL	201	snal	.e	-+
48	22	16	<u> </u>		NSSW		E Lime	stone	
49	10	┝─┼┼			NE SW		<u>c 1010</u>	rai te	+
IIF	Matrix	Compos	ition	Data: Col	or <u>Tan</u> ay Minera	Hu 1 Compo	nsell Soil sition	Color No. 2	SY 5/4
Gra	in-Siz	e Distr	ibutio	n Expan	dable		Chlorite	Carbonate	s Cts./S
54	Sand Silt Clay Clay			ay I	llite	plus	Calcite	Dolomi	
	20	40	40	Mine	rals		Kaolinite	7	1 17
	_		_				and the second se		

Figure 35 Data for analysis site E3

Analysis Site Number E4 Morphostratigraphic Unit Eureka

Location SE SW SW Sec. 34, T.25N., R.1E. Quadrangle Danvers

Elevation 840' Depth of Analysis 6' Type of Exposure Roadcut (old)

Topography B and C slopes on distal edge of moraine

501	Pebb	e	Jata	:

No.	Dimensions		Shape	Str	ike	Di		Lithology		Misc.		
	46	57	20	5	N 35	E	90	<u> </u>	Gree	nstone	1	
2	26	<u>† 15 –</u>	1 15	3	170	<del></del>	0		TII		+	
3	15	1 13	8	7	140	W	351	:	Uplo	aite		
4	61	44	32	7	3150	W	40	·	1:01:01	nite	1.59	
5	10	5	3	4	14.0		10:	11	' 010'	nite (?)	<u> </u>	
6	12	11	5	7	16.6	N	3.1	2	Cher		1	
7	12	4	7	5	1:41	W	3(1)	1	l'in!		1	
8	11	1-7	6	6	111.0		0	· · ·	L'her		1	
9	17	15	7	1	1175	E.		51	Line	stone		
10	12	н	7		181		0		Pele	nite	1	
11	12	8	6	G	Mpr	2	401	:	1111	tone	-	
12	18	12	19	4	N41	IE I	1:01		1010	nite	+	
13	13	12	7	8	365	0	105	1	1010	nite		
14	15	13	<i>'</i> 9	5	1.70	12	10	iC 1	Uner		+	
15	22	14	12	L	11.5	1	451	VE.	1 1' 11	11		
16	18	14	4	]	N65	E	90		5211	•	1	
17	]4	9	9	5	15	i l			1.010	aite	1	
18	6	E E	3	3	111	w l	Q		1.25.09	stone	1	
19	18	14	8	7	151	;	25	<del>; 1</del>	ther		1	
20	9	8	3	1	145	E	0	+	Fed :	hile	1	
21	32	27	5	1	1155	E	0		Chel	) )	i.V	
22	15	11	5	5			20	0	Cher			
23	30	24	3	1	165	E I	451	N.	Shile	3		
24	21	12	10	6	N35	C	0		1.010	nite	1	
25	11	7	4	2	NIC		109		Gran	ite		
26	16	12	11	5	715	<u>.</u>	10.	,	0101	aite		
27	12	8	7	<u>(</u> ,	1.1	E T	15:	N	Dolo	nite		
28	12	6	(	6			7.19	V I	1.7-0	tone	1	
29	13	9	ć	6	N 30	E.	10	VET	Dolo	nite	1	
30	11	S	5	6	ŀ		- 0	-	Cher			
31	16	]4	13	5	N20	2	25		Dolor	nite	1	
32	12	8	4	6	145	i:	303	VE I	Line	tone		
33	10	E	3	4	N85	1	301		lines	tone	· · · · · · · · · · · · · · · · · · ·	
34	16	12	10	7	1.		55		Polo	nite		
35	36	35	19	7	15		301	: 1	Polo	hite		
36	15	11	10	5	11.15		0	-	Gran	ite		
37	16	11	7	(	:11	·	255		10101	nite		
38	12	11	7	7	17		0		Cher		1	
39	11	7	- 6	5	N 75	C I	15:	W	1:010	ai te	1	
40	10	8	2	5	N		20.	: 1	1010	nite		
41	15	10	- 8	6	E		1.1		Flin			
42	13	11	5	7	N30	6	90		Cher		I	
43	16	11	6	6	1185	£	35/	V I	[\o] or	nite		
44	44	35	10	1	N4 5	E	301	VE T	Dolo-	nite	I	
45	11	6	5 ]	4	N55	E	10	VE T	Dolo	nite		
46	8	7	6	5 1	N 75	E	35!	NE I	Cher		I	
47	11	6	5	6	N4C		0		Polc	rite		
48	20	15	8	3	NIC	E	408		Dolo	nite	Γ	
49	11	2	6	6	N 30	DE I	40	NE I	1010	nite		
50	22	15	14	5	N70	30	15:	IE I	Geod	2		
TILL	Matrix	Compos	ition 1	Data: Colo	or Ta	in	Mu	nsel	1 5011	Color No. 2.	SY 574	
	1. 01-	Diet	dhutd		ay mine	eral C	ompo	5171	on			
	111-51Z		100110		uadre	7114	- 1	ú	LOTITE	Carbonate	5 UTS./Sec	
		25 1			ay	1111	te	v	pius	Calcite	rolomite	
			70		rdis			<u>kac</u>	T			
				L"	·	0.5		<u> </u>	10	l		

Figure 36 Data for analysis site E4

Analysis Site NumberE5 Morphostratigraphi	c Unit <u>Eure</u> l	<a< th=""></a<>
Location SE NW NW Sec.10, T.23N., R.6E.	Quadrangle	Arrowsmith
Elevation 900' Depth of Analysis 5' Type	of Exposure	Pipeline Trench
Topography B, C, and D slopes near top center	of moraine	

0.	D	imensio	ons	Shape	Strike	Dip	Lithology	Misc
1	15	13	ð	3	1765E	0	Limestone	
2	11	6	6	6	11201	30;;	10100126	
3	13	11	4	2	1:45:1	25/12	Polonite	
4	11	Ŀ	6	ს	1120E	155.8	Limestone	
5	]4	<u>1</u>	2	1	- 180M	100	Chale	
6	19	13	11	5	1+30E	1 15. 1	Dolomite	
7	18	- 12	7	7	145E	35. N	linestone	
8	15		3	1	11E 5W	21.5	Enale	
9	13	33	4	2	1.4 OE	35:00	Dolomite	
10	11	9	5	5	::5.1	10.5	Dolomite	
11	36	11	5 1	2	N 35W	1 10:1	Solomite	
12	35	12	11	5	:1755	0	Limestone	
13	9	5	4	4	::40E	10.77	Lisestone	
4	11	8	5	5	1.401	-90-	Cuirtzite	
5	12	10	- a		11101	30.	Cabbro	
6	-		- i	- Á	1.705	200	Quirtzite	
17	- 52	17						
6							La courte	
3							holowite	
2	-::-		28	~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~	100W	t	horomite	
<u></u>	<u>_</u>			<u> </u>			POTONICE	
1	58		25	ь	1356		Conglonerate	
12	- 22	15	8	6	1135W	<u>55%</u>	Dolomite	
23	22	14		6	!:	5.51	Dolomite	_
24	12	5	5	1	332	20.	Limestone	
25	20	17	10	7	14^^E	45:15	Siltstone	
26	_15	<u> </u>	4	6	:.43E	12:34	thale	
1	19	38	14	7	3308	55.7E	Dolomite	1
28	]4	13	1	7	:13 );;	10:1:	Limestone	
29	11	7	6	4	N250	0	Shale	
30	15	13	9	7	1.SUN	10.	Chert	
31	21	10	5	1	11351	10:22	Shale	
32	12		4		1.500		Sitstone	
13	15	10			151	200	Chort	
34	16	10			1254	10.5	Dolonite	
15	16	$-\frac{1}{10}$			110561		Chalo	
2	10				10.00		Chale	
20						7.51.	- Chaite	
26	- 24					335-	DOTOULLE	
0							Quartaite	
22	3/		16	<u> </u>	<u> 445E</u>	40	1.1mestone	LAV
<u></u>	15	-13-	10	<u>_</u>	W		Quartaite	
1	52	17	16	.,	N7.W	401	Liolomite	
2	15	10	5	1	Nasw	45::	Dolomite	
3	21	18	11	2	NEON	40:40	Giltstone	
4	21	30	1	4	1:70W	20:T	Jolomite	
5	13	12	2	1	N20W	30	Shale	T
6	15	1	6	2	N75W	20.:	Shile	
7	14	11	7	5	N26E	<u> </u>	Polomite	
8	16	10	10	7	NEOW	30:11	Shale	
9	18	15	$-\frac{1}{10}$		1:05W	4	hert	
0	24	18	$-\frac{1}{12}$		NAOE	SUNE	Chert	+
11	Matrix	Compos	Ition D	ata: Colo	or <u>Tan</u>	Muns Composi	ell Soil Color No. 2	2.59 574
Gra	In-Size	Distr	ibution	Expand	lable		Culorite Carbonate	es Cts.75
Sa	nd	STE T	Clav		111	ite I	plus Calcula	
20		10 1	- 34	H Mine		···   v	alinite	1 10
<u>د ر</u>		1	J.7	1 1 1141101	i terteren i	1 1		1 17

Figure 37 Data for analysis site E5

Analysis Site Number _E6 ____ Horphostratigraphic Unit _____ Eureka

Location NE GW SW Sec. 35, T.24N., R.6E. Quadrangle Gibson City

Elevation 880' Depth of Analysis 5' Type of Exposure Pipeline Trench

Topo	Topography B and C slopes near top center of moraine												
<b>F111</b> P	ebble Lata:												
No	Dimensione	Chane	Strike	nto	Lithology	Mac.							

No.	Dimensions		Shape	Str	ike	Dij	P	Lithology	Misc.			
	14	в	7	6	N5E		0	De	olomite			
2	31	2.4	1 13	2	1115	E.	100	- Pe	olomite			
3	30	20	10	1	1125	30	154	E (1	hert			
4	19_	9	6	<u>د</u>	N151	H,	15.7	A De	olomite			
L S	16	11	1	2	11/5	E	401	<u>a – 5</u>	hale			
6	19	14	2	1	1/81	in		S	1410			
7	15	12	7	6	Ner	<u>M</u>	- 90	C)	hert	LVA		
8	14	10	7	6	115%		0	D;	iorite			
9	37	27	21	5	NSS	W.	30::	:/D	olomite			
10	<u> </u>	9	8	5	N60	1d	25		iltstone			
11	21	18	15	5	1150	W	401	a	hert			
12		$+ \mathbf{n}_{-}$	3	6	1 <u>NEC</u>	14	0	C	hert			
13	12	- 9	1	<u> </u>	NAC	2d	355	<u>i</u>	olomite			
	14	1 12	8	<u>S</u>	<u>N65</u>	<u> </u>	0		ime: tone			
15		<u> </u>	6	6	NZ5	E	200	d L:	inestone			
16	10_	1_1_	4	1	1125	ül		E	olonite			
1/	16_	1 11	10	6	نيكال	i	50	P	olonite			
18	13	1 10	<u> </u>			w	<u>_30N</u>	_	ilt_tone			
19	19_	1 13 -	5	1		<u> </u>	<u>-400</u>	d£	iltations			
	17	1-15-	12	3	- 115		255	E				
	16	1-10-	<u> </u>	<u> </u>		<u>+</u> +-	<u> </u>		inestone			
	-22-	1 <u></u>	6	11		<u> </u>	0	· · · · · · · · · · · · · · · · · · ·	14e			
	_ير_	1	<u>, ( –</u>		<u>- N25</u>	<u>k</u>	تنكي	<u>u - 1</u>				
	-32-	+ 2.	18	<u> </u>	<u> </u>	<u></u>	- 90		oromre — —			
26	-13	+	- <u></u>	<u> </u>	<u> </u>	<u>u</u>	-450	ai	olouite			
27	-20-	+ +++++++++++++++++++++++++++++++++++++				<u></u>	. 1044		Jule			
-36-		1 <u>1</u>		— <u>;</u> ——		<u> </u>	-767	<del></del>				
24	12	t - t					90	a	lint			
1		1-14	E.	·;	- 117	<u>;</u>	-443	-+	hale			
1	30	1 11	- 23	ξ	1175		-400		ihhro			
10	16	1-11-		6			101		alouite			
1-11-	16	1 13	12	3	MIS		-255	;;	1)tstone			
34	16		4	7	1195	E I	100		Domite			
1 35	24	16	-5-		N35	r -	-505	V S	ltstone			
36	16		3	1	280	DE T	251		nale			
1 37	12	8		2	215	E	805		hale			
38	12	4 4		6	N60	W	2510	J 1	imestone			
39	13		- 9	7	N50	5	90	De	olomite	LAV		
40	13	9	8	6	NSE		155	В	isalt			
41	32	23	9	1	NIC	W	0		olomite	- <u>†</u>		
42	16	12	8	6	N20	ra 🕇	<u> </u>	Be	nsalt			
43	43	40	16	7	N	1	505	0	nert	1		
44	66	36	28	6	N20	1.1	20M	V G	111 20	S		
45	33	23	13	3	NSS	W T	30:	V T.	in stone			
46	9	77	4	2	1151	W	105	E De	picmite	1		
47	12	10	7	5	1	·	10.5	. 10	alomite	-T		
48	28	20	14	6	N55	W	55M	N DO	olomite	T		
49	23	18	12	7	NSW		158	10	olomite			
50	12	9	9	5	N15	E	25N	GI	ranite	I		
TILL	Till Matrix Composition Data: Color Tan Munsell Soil Color No. 2.57 5/4											
							_					
1				CI	ay Min	eral Co	ompos	ition	]			
Gra	in-Siz	e Dist	ributio	n Expan	dable		T	Chlorite	Carbonat	es Cts./Sec		
Sa	nd	Silt	C1	Clay III			plus	Calcite	Dolomite			
2	4	39	Mine	inerals			Kaolinite		15			
1					5	81		16				
1												

Figure 38 Data for analysis site E6

### Figure 39 Data for analysis site E7

1 1	21	1 12	8	6	NSE	51	8.15	110				
2	26	17	10	6	1.115	V		tstone				
3	13	e ع	6	4	I N201	: 0	101	omite				
4	20	16	9	7	1110/	V 55:	i Che	rt				
5	14	11	- 11	5	1130	1 10	W Bit	ult.				
6	12	4	6	5	11201	200	I Che	rt				
7	17	12	10	3	N 354	35	9 Dol	omite				
Ŕ	17	12	10	6	125		tio!	ouite				
- ö	5.6	50	17		110			throne	LAV			
10		1.			N 400	- 20	E DOI	onite				
			11		N/0	<u> </u>		1:1000				
+++++++++++++++++++++++++++++++++++++++	14		7		1.5	· · · · · · · · ·		()mito				
		14			ND 50			conite				
			14		1 10 50			Carte				
19	21	13				10						
12	20	18	15	9	1170:		· · · · · · · · · · · · · · · · · · ·	ies cone				
16		<u></u>	4	6	1.1	10	101	onte				
17	18	15	10		1	101	101	Unite				
18	16	12	8	<u> </u>	1301	. 0	1.17	restone	-+			
19	36	22	9	1	1156	701	<u> </u>	omite				
20	20	]4	10	3	N451	<u> </u>	N DO!	chite	- <b> </b>			
21	33	10	5	1	11158	: 45.	I Dol	onite	-			
22	30	6	3	6	1170v	v 20.	l Dol	onite				
23	19	16	9	5	NSW	90	001	omite	1.AV			
24	17	10	6	6	1 :41 0	0	Sha	le				
25	11	7	3	6	N20	0	1/10	ristone				
26	11	8	5	6	1451	: 102	E Lin	nestone				
27	ìQ	ú	4	6	11201	40.	t. 101	omite				
28	20	13	12	5	N2 04	1 15:	V Dol	Chite	W			
29	5.0	40	- <u></u>	1 1	N550	1 151	Sha					
30	17	-14	10		1 110							
	22	$\frac{1}{10}$	12		NAO	2 200						
	15		<u> </u>		11.0	201	<u></u>	onito				
12	-10-		- <u>c</u>		120				-+			
22	10		4		1.60			escone				
- 34	10	-11-	10	·	NEW		·	<u></u>				
- 33	74	1/			NZ3			Q"11.				
36	18	12	10	<u> </u>	1257	-255	E	icr.tione				
51	14	10	<u>ر</u>	2	N/U	<u>v 10</u> ;	·	rt				
38	12	11	<u> </u>	77	N60		<u>Chr</u>	<u>rt</u>	- 4			
39	12	10	6	5	NS01	35	W Lin	instone				
40	36	12	8	7	<u>N45</u>		Che	rt				
41	22	15	7	2	<u>Ni25E</u>		E Che	<u>rt:</u>				
42	12	10	4	1	N10:	3.1	i Chr	·rt				
43	10	4	2	4	1:70	<u> </u>	<u></u>	<u></u>	·			
44	30	20	16	3	N25V		B.15	, ilt	<u> </u>			
45	16	7	6	4	N 759	V 25V	l l'ol	omile				
46	21	17	11	3	1:20	V U	201	onite				
47	21	12	9	6	N	155	I Dol	omite				
48	13	9	6	6	N15	201	I Sha	le				
49	18	14	E E	6	N10	V 45N	l Sha	le				
50	36	50	22	8	N/5	351	.in	estone	S			
	<u> </u>				• • • • • •			CALL & MALE				
TIL	Matrix	Compos	sition	Data: Co.	lor <u>Ta</u>	an Mu eral Compo	insell Soil	Color No. 2	.5Y 5/4			
Cra	In-SI?	e Dist	ributic	Expu	ndable T		Chlorite	Carbonat	es Cts./Secl			
	nd T	SILE	Clay		lav	Illite	plus	Calcite	Dolomite			
1		10	- 24	- Min	arale		Kaolinite	2				
ــــــــــــــــــــــــــــــــــــــ	<u> </u>	20	<u> </u>		1 972							
ł				L	·	t	12	1				

Topography B, C, and D slopes on proximal side of moraine

Shape

HII Pebble Data:

Dimensions

No.

Analysis Site Number E7 Morphostratigraphic Unit Eureka

Elevation 810' Depth of Analysis _____6' Type of Exposure Pipeline Trench

Strike

Dip

Lithology

Misc.

Location SW SE SW Sec. 24, T.24N., R.6E. Quadrangle Sibley

Analysis Site Number E8 Morphostratigraphic Unit Eureka

Location NW NW NW Sec. 3, T.23N., R.3E. Quadrangle LeRoy

Elevation 840' Depth of Analysis 5' Type of Exposure Roadcut (old)

Topography B and C slopes on distal side of moraine

fill Pebble Data:											
No		imenei	<b>ADe</b>	Shane	C++	-110			thology	Miac	
	1 10		TA	A	- NI		511	L'OLOT	ite	-	
	1 12-	10	+		-+	<del>77</del>	n:	Dolon	ite	+	
3	13	7	6	4	NAC	7.0 5	51112	1:15 all		1	
4	8	5	4	5	116	N I	0.7	liolen	ite		
5	11	н	4	5	115'	W 1	0:12	Shale	(spores)	1	
6	25	22	12	7	N //	JN J	OF E	Polon	ite	1	
7	10	7	5	6	N10	N 1	<u>े ।</u>	Dotom	ite		
8	10	7	4	6	:14(	M 2	(m) 2.9	1.2 mes	Lone		
9	37	23	12	7	N 30	M 1	14. <u> </u> ]	Chert			
10	25	74	12	7	1	: ] ]	(1.1	10100	ite		
11	13	8	5	6	!	1 <u> </u>	58	Finent.		1	
12	27	17	14	5	N60	W 4	ONE	Doloa	ite	W	
13	10	8	4	4	1180	<u>):</u> 3	011	Chert			
14	17	10	8	7	111	W 2	55	Chort			
15	15	10	7	5	N79	1.1 1	54	Chert			
16	10		5	6	112(	20 2	0.4	10100	i''C		
+ 1/	17	14	11	+	<u>NS/</u>	<u></u>	0	- Shille			
18	13		10	6	- 114	<u> </u>		Grani	ce	+	
19	51	27	1 12	<u> </u>		<u>w -                                   </u>	0	51105	Lone	LAV	
	20	<u> </u>		<b>↓</b> − − <u>↓</u> − − −		<u>sw</u> <u>s</u>	0HW	inert		+	
	+ + 5		10	<u>↓</u>	- 112	<u>- 9</u>	0	1.010m		+	
	18	1/	<u> </u>		_ <u></u>	<u>.</u>	<u>5W</u>	1010m	ite		
	16	- <u>,'</u>	<u> </u>	4	- <u>N</u> 2		0	<u>naic</u>			
	16		<u> </u>	8	121		<u>0:w</u> _	1.010m	120		
1-22-		17	<u> </u>			<u></u>	<u></u>	<u></u>		·	
	- 4/	- 10				<u></u>	CN		cone		
	-27		14			<u>~</u>	SNW .	1:010m			
24	21	- 20				N 2	<u> 211W</u>	<u> </u>		+	
- 23							int-	nile		<b></b>	
	- 21	10		<u></u>			2017				
-6-							0011	Chilo		+	
- 11							0	lialon		+	
34			6		- N'2		0	1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1 1			
35	25	20			NS(	<u> </u>	50	Lines	ione		
36	17		10	6	NA	NJ G	500	Chert		1	
57	- 56	19	15			<u></u> 5	0w	Chert		+	
38	14	11	10	7	- 11	W 1	51	Chert			
39	15	13	6	Ŕ	N 50	w 1 5	000	Lines	cone	+	
40	59	18	5	<u> </u>	NB	W 1	0M	Shale		+	
41	28	15	12	6	NRO	WII	ſW	Limes	tone	1	
42	19	12	5	2	N 79	W 1	ON	Shale		1	
43	13	10	6	7	NBO	W O	0	Lines	tone	t	
44	19	17	15	5	N4	5 5	5N.V	l'olom	ite	1	
45	28	24	10	2	N50	W D	0	Shale		LAV	
46	26	17	8	7	N79	DE 2	5E	volom	ite	1	
47	22	15	10	6	N50	W 9	0	Quart	zite	1	
48	67	49	32	5	N 51	V 4	55	Polom	ite	<b></b>	
49	25	22	10	2	N 39	9 W	0	Limes	one	1	
50	25	15	12	3	N70	W 3	0.V	Silts	cone		
Till Matrix Composition Data: Color <u>Tan</u> Munsell Soil Color No. 2.5Y 5/4 Clay Mineral Composition Grain-Size Distribution Expandable [ Chlorite Carbonates Cise/Sec]											
Sa	nd	511:	Clay	·	Пау	Illite	1	plus	Calcite	ivlomite	
ني ا	8	43	29	Mir	nerals		Ka	olinite	6	21	
				L	3	81	1	16			

Figure 40 Data for analysis site E8

Analysis Site Number <u>E9</u> Morphostratigraphic Unit <u>Eureka</u>

Location NE SW SW Sec. 10. T.24N., R.2E. Quadrangle Normal

Elevation 840' Depth of Analysis 5' Type of Exposure Roadcut (new)

Topography C and D slopes near top center of moraine

No.	D		ons	Shape	Strike	Dir	a	Lithology	Misc.
<u> </u>	42	28	23	6	1185W	2.5W	Dc	olomite	5
2	46	38	21	5	E	1 गणा	1	DI ONT CG	<u>+</u>
3	20	14	10	7	N65W	25.1	<u>;</u>	ert	
4	1-22-	11	111	3/	N 35C	1 15		lomite	T
<u>~</u>	42	1 52	<u>↓ </u>		NEOE	5111	<u> </u>	lomite	<b></b>
<u> </u>	25	20_	<del>↓+</del>	<u>1</u> /	11808	30.	<u> </u>	ule	<u> </u>
- <del>'</del> -'	1 1 1	+-10	╉╌┽┽	jl	11805	+		lomite	
		<u>↓ 4</u>	<u>↓</u>	<u></u> I	N	+		lomite	<b></b>
-16-1	1-16-	1 20	+ <u>+ <u>+</u> <u>+</u></u>	<u>`</u> {	N/01	$\frac{1}{100}$		lomite	
11 1	1-15-	1	┝╌╦╌┤		MAGW	+		/ICMILE	-+
12	1-19-1	+ + - + - + - + - + - + - + - + - + - +	1-10 t			+ 200		) 1 CM LCE	· <b> </b> · · · · · · · · · · · · · · · · · · ·
13	1 18	+++++++++++++++++++++++++++++++++++++++			111.00	1 500		Jonito	+
14	1 25	1 20 1			NUSA	+		lomite	+
15	28	1 25	20 1		r <u> </u>	+ - <u>40</u> -	-+	Monite	+
16	28_	20_1	14 I	6	NESE	104		1) omite	-+
17	20_	1 13_1		4	NASE	1 35:00	-t	Nomite	
18	<u>11</u>	5_7	41	6	N45L	+ <u>5</u>	<b></b> 1.1	mestone	
19_1	15	14	121	5	HEUN	1001	10	lanite	+
20	13	10	71	<u> </u>	NSSE	10!1	<u>i t - c</u> i	ort	+
21	24	22	12		N75#	50 -	it ne	lomite	+
22	26	15	10	6	1 1401:	90	Dc	olomite	LAV
23	15	12	10	5	N 707	<u>† 10w</u>	N De	olcinite	
74	55	57	24	6	N70E	10:	1.1	mestone	
25	<u></u>	25	]4	5	N45.V	30NE	aDc	bionite	+
26		16	2	5	11751	<u>104</u>	nc	louite	
1	[_26_]		10	5		101		licaite	1
-36	ر يە	1 32 1	24	<u> </u>	lieog	2011	De De	olomite	1
29	29			4	NEOE	158	TCr	nert	1
30	46	39	19		NIUW	45N	De	olomite	1
31		13	-11		N75W	90	TDc	olomite	1
32	45 ]	44	20	3	N40W	0	1_13	nes tone	T
33	22	18	14	<u></u>	N15E	45.1	J DC	olomite	1
34	<u>10</u>	8	4	8	E	25:.	<u>Qv</u>	artzite	1
35	ر در آ			3	N20W	103	1:0	olomite	1
36	20	15	10		NIOW	20N	De	lomite	I
37	16	10		<u> </u>	N50E	25:4		mestone	Τ
38	21	11	10	<u> </u>	<u> </u>	10W		inrt	<u> </u>
39	25	$\frac{21}{21}$	$\frac{18}{18}$		NESW	251	T De	Jomite	T
40 1	-22	-15-		<u>-</u>	<u>N5//</u>	40N	<u>L</u> 1	mestone	1
41	$-\frac{12}{2}$		4	<u> </u>	1150W	0	<u>pe</u>	lomite	T
44		-14		<del></del>	<u>N15w</u>	10N	<u> </u>	mestone	
	-16-	┍═╧╡═╋			NOUE I	1291		ude	
44	18	<u>⊢_16</u> +			<u>N 700</u>	<u> <u> </u></u>	<u></u>	iort	I
42-1	12	<u>→</u>			<u>N50w</u>	350L	<u>1 11</u>	mestone	
40-1	-30+	-12+	<del>}+</del>	<u></u>	<u>N700</u>	400	10	Jonite	1
**					<u>N 50E</u>	JUNE	Do	<u>isalt</u>	⊥,
40		-14+			<u>N501</u>	35Nr	<u>}</u>	lomite	<u> </u>
**		<u>/</u> +		<u> </u>	N	407		lomite	+ <u>,</u>
<u> </u>	<u>ــــــــــــــــــــــــــــــــــــ</u>	<u> </u>		<u> </u>	NBOW	JUW	<u> </u>	inds tone	W
411	Vatrix	Tompos	TETON D			Mun		Totos No. 2.1	
*** .	10.000	Compose	1000 -	dia	<u></u>		Sell Jun	(010F NU	<u>)Y 5/4</u>
					iy Mineral	<u>Compos</u> :	ition	]	
Gra	in-Size	2 DISLT	ibution	Expanu	lable	. 1	Chlorite	Carbonate	s Cts./Se
541	<u>id</u>	SIJE +	Clay		іў — ттт	ite	plus	Calcite	Polomit
1.4		40	311	]   Miner	cals (	1,	Kaolinite		27

Figure 41 Data for analysis site E9

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Analysis Site Number E10 Morphostratigraphic Unit Eureka

Location <u>NE SE Sec. 3, T.25N., R.1W.</u> Quadrangle Danvers

Elevation 770' Depth of Analysis 6' Type of Exposure Roadcut (new)

Topography C, D, and E slopes on distal side of moraine

III Pebble Data: No. Dimensions Shape Strike Dip Lithology Misc. Т 12 N70C 90 15 Dolomite LAV B 2 13 9 6 1135! JUIL Linestone W 12 7 9 1145% 6 5500 Chert 4 45 23 NSCE Fink Dolomite 13 ZONE 14 10 6 7 NASW 35114 Chert 6 12 4 20 8 NICA SUN uolchite thile 12 8 2 5 1706 105; 8 13 £ 5 6 N25E 150W Polonite 9 16 15 5 18.94 20.1 2 Linestone 10 13 N300 8 f, 6 250 Limestone 11 100W 15 10 7 140E 5 Polomite 12 14 11 8 ٢, N55E 0 Dolonite 13 9 7 3 11 4CN Dolonite 14 52 32 22 NZOE 304 Dolcmite E 15 13 10 6 N20E LCSW Doloaite 5 16 NSSW 21 17 15 8 90 LAV Linestone 17 15 12 ۲, 104 E Shile 18 16 14 11 112.04 Limestone 255 19 17 13 10 NIOE 305 Limestone 20 15 35.W 12 10 7 N2SE Quartzite 21 12 14 300 115E .9 Quirtzite 22 18 14 13 NISA 401 Dolomite 450E 24 NASE Limestone 20 10 24 J ONE 10 N45C 14 9 Nolomite 25 115:1 16 38 23 15NL Dolomite 26 17 105 19 10 N755 Basalt 27 28 20 15 15 3 ND OM 0 Lisestone Lines Lune 10 25. N35E 29 30 10 5 5 20M Chert 1: 5 N85W 65E Dolomite - 9 7 6 NICH 31 10 13 8 20N Red Jasper 32 38 N2OL 30N 14 Dolomite 6 33 14 N105 3011 21 13 6 Chert 15N 20NW 34 10 Dolomite 17 NSW ક 35 15 **J**0 5 N55W Chert 36 13 10 2 N45E 13NC Shale 1 37 14 10 8 NGSE Dolomite Ω <u>38</u> 39 50W 7 2 N85E 12 Snale 26 14 15NW 15 N55W Chert 40 12 7 5 NBOM 30.1 Dolomite 6 41 15 7 N255 205W 6 6 Shale. 42 14 10SE 9 8 Dolomite 6 NSUN 43 13 12 4 NBOW 204 thule 44 9 Chert 13 12 N35E 105W 45 12 9 5 S N35E 10NE Polonite 46 14 11 7 N45E 6 5:1 Chert 47 17 11 7 1000 6 N254 Dolomite 48 19 12 9 N2OE 25SW Dolomite 6 49 10N 40 N5E 20 6 Sh.ile s 50 10 NTON 12 15CE Dolomite 6 3 Till Matrix Composition Data: Color Builtan Munsell Soil Color No. 2.57 5/4 Clay Mineral Composition Grain-Size Distribution Expandable Chlorite Carbonates Cts./Sec SIL Illite Sand Clay Clay plus Calcite | Dolomite 42 20 38 Minerals Kaolinite 13 5 84 11

Figure 42 Data for analysis site E10

Analysis Site Number <u>Ell</u> Morphostratigraphic Unit <u>Eureka</u>

Location _SW NE SE Sec. 34, T.26N., R.1W. Quadrangle Danvers

#### Elevation 700' Depth of Analysis 6' Type of Exposure Roadcut (new)

Topography C, D, and E slopes on proximal side of moraine

<b>F1</b> 11	Pebble	Data:	

No.	D	Dimensions		Shape	Sti	rike	D	ip	Lithology		Misc.	
	10	7	5	4	T N	1 OE	0	)	Flint		1	
7	15	12	1 7	1 3	1 11	inw -		)	Chert			
3	16	11	2	6	1.	5512	1'	Ŵ	1.1115	tone		
4	12	10	1	5	ni Ni	301;		, C,	1:010:1	ute	1	
<b></b>		115	6	↓	<u>       N</u>	<u>50</u> W	2(	):1.1	Lices	tone		
	20	12	<u></u>	6	<u> </u>	<u>858</u>	21		Lines	tone		
	<u> </u>	7	<u> </u>	6	is	55W	· · ·	)	<u>_Grate</u>			
	1		4		<u> </u>	<u>7 F.</u>	(	!	<u>G (bbr</u>	0		
10	<u> </u>	12_			<u> </u>	251	1		liolor	nte		
- 10	-13-		<u> </u>		<u> </u>	<u></u>	1 11	111:	1.1:205	tone	∔	
12				<u>b</u>	<u> </u>			1114	1.11.6.5	tore	<b>}</b>	
1				<u> </u>			+	<u></u> _	10.01	.1.00	<b></b>	
111	- 12-	10		<u>├</u>		<u> </u>	<u> </u>	· · · · · ·	1,164	Cone	+	
15	<u> </u>	- 10	12		<u>[1</u>	<u>101</u>			- POTC:		<b></b>	
16	20	- 20		<u> </u>		<u>21.</u> A 1.11	<del>1 - ;</del> ;		1.1.1.1	Long	+	
17	11	12	51		;;	SUP	1-20			tone	+	
18	16		4	1		2.84		11.5			·····	
19	39	- 25	22	<u> </u>		150		<u>,</u>	10107	nite	1.1.1	
20	11	E	7	7	;;	3012	20	101	10107	lite	· · · · · · · · · · · · · · · · · · ·	
21	17	E	7	6		251	1		1'010	ni te		
22	15	12	10	н	Ņ	305	21	ELE 1	[olo:	ite		
25	18	13	<i>'</i> }	5	N	5714	41	11.1	1) <b>0</b> ,10ir	ui te	†	
24	13	11	7	5	N	054	24	555	Unort	v Polotite		
25	35	12	9	7	N	206	10	I.W	Lines	Cone		
26	_11_	7	5	6	N	154	1	.I.A	Dolor	ite		
27	] 1	13	H	5	И	SIN	10	1.1.1	Dolo	nite	1	
26	11	9	7	8	N	4 MW	4!	. E ]	Quart	zite		
29	27	_17	9	4		15W	9	5C	Limes	tone		
30	14	12	8	7	N	5 OW	) 9	IW ]	_Dolor	ite		
31	12	7	6	6	<u>N</u>	5.1	10	DN	00107	ນ່ ::e		
32	12	7	5	66	N	F5C		21	Chert			
- 33		7	6	5	N	F.OW		E		uite		
34	43	24	15	1		<u> </u>	20	<u> </u>	Shale			
1-32-1		- 18	<u>_15</u>		<u>N</u>	1:5W	90			iite	LAV	
- 20-	- 52				<u>N</u>	<u>75%</u>		<u>-</u> +	Lolor	ute		
				6		<u>r SW</u>		<u>/</u>	00107	<u>te</u>		
10						WWW	<u> </u>	<u>-</u>		cone		
1 40 1	- 10 1					3 <u>-</u>		+ ·	<u></u>		·	
41	- 38		- 42-	,		<u>258</u>		1.1.1				
42		- 12				71.2	$-\frac{2}{10}$	344				
1 43			- 20	6		N			Time	1.000		
44	- <u></u>	-14	10		N	CIRJ 1			(35.10)	te	• · •	
45	- 23	19	11	6	<u> </u>	6114		╤╧╋				
46	35	-211		1	N	751	<u>-</u>		11.10			
47	16	12		ī		601			hild			
48	33	22	14	5	1	20W	1		Lolon	ite		
49	11	<u> </u>	5	i i		TW	-i(	set.	CULIE	zite		
50	24	11	ť	6	N	40W	10	0.00	thert			
											······································	
	Till Matrix Composition Data: Color l'an Munsell Soil Color No. 2.59 5/4											
1												
				<u> </u>	iy Min	eral (	Compo	ositi	on			
Grai	n-Size	Distr	ibutio	n Expand	iable			- Ch	lorite	Carbonates	Cts./Sec]	
Sar			Clay		y I	Illite			plus	Calcite	Dolomite	
	<u></u>	22	29	Miner	Minerals			Kao	linite	10	20	
I				<u> </u>		8	4		ذل			
L												

Figure 43 Data for analysis site Ell

 Analysis Site Number N1
 Morphostratigraphic Unit
 Normal

 Location NW SE NV Sec. 18, T.23N., R.6E.
 Quadrangle
 Arrowsmith

 Elevation 830'
 Depth of Analysis 5'
 Type of Exposure Roadcut (old)

No.	D	imensi	ons	s	hape	Str	ike	DI	P	L1	thology	Misc.
<u> </u>	73	51	35		5	1110	)E	15	N	:'olomite		
2	17	12	1 10		6	1:1	·ř	35.	12	Bacalt		, N
3	26	36	12		7	N50	DF:	0		.'or 1 10	ncil	
4	32	27	20		ы С	I N1º	ΓA	25		Granite		
5	26	22	12		8	NS(	JF;	- 20	1	lolomite		
6	]4	11	6		5	192	;t,	1.10	11	Lineston	¢	
7	34	20	18		6	1		0		<u>bolomite</u>		
8	13	9	5		1	12.	<u>, 1</u>	- 25	· ii	Lineston	<u>c</u>	
9	]4	12	1		5	16	<u> </u>	30	<u>.</u>	Dolomite		
10	37	21	13		5	14	5.1	0		Lineston	<u>0</u>	
11	30	14	17		<u> </u>	111	<u>5</u> 1.	12	( ·	Guirtait	<u>0</u>	
12	12	10	12		<u> </u>	111	5/1	20	<u>ii</u>	(hort		
13	20	15	11		<u>5</u>	1144	4.1	<u> </u>		<u>on artzit</u>	<u>e</u>	
14	24	15	12			1.11	<u></u>	25		lolente		
15	14	10	/		Ь		);; <b>[</b>	10	<u></u> .	lines or	e	
16	16	1-15-	6			1	N.	20	<u></u>	<u></u>		
17	14	<u>12</u>	4		<u> </u>	1 <u>- N44</u>	<u></u>	13		thert		
18	14		1		<u> </u>	N7	<u></u>	<b>⊢</b> <u>†</u>		1.11.11		
- 19	12		<u></u>			<b>├</b> ─_ !!!!	<u></u>	<b></b> 15	<u>.</u>	Quirtzit	<u>e</u>	
	24	<u> </u>	14		<u>b</u>	<u> </u>	<u>, , , , , , , , , , , , , , , , , , , </u>		<u>.</u>	1.5.1.C		
	40	- 44	-25		<u></u>				<del>.</del>	00100100		
	<u> </u>	<u> </u>	<u> </u>		÷		<u>)[</u>		<u></u>	Flint		
	<u> </u>		<u> </u>		-	<u>N/</u>	<u></u>	10	<u>х</u>	Limeston	<u>c</u>	
24	- 13-	<u> </u>	<u> </u>		<u>-4</u>	<u> </u>	<u></u>		<u>×</u>	SILCECO	<u>e</u>	
							<u></u>	- 22	2	Chert		
	- 35-	- 50	-18		<u> </u>		<u></u>	- 15	<u> </u>	Lines tor	<u>e</u>	·
	- 22-		<u> </u>		<u> </u>					11100000	2	·
29	14	- 14	- 9		<u> </u>	NA S	JE	10	C.F.	Light Cr	vstalline(2)	
	22	- 26	19		<u></u>	NR:	18	-=="	26	Unlogite	ystarrine(.)	
-11	14		- <u>+</u> +		6	- N1 (	11			315 11		
-12	14				<u> </u>		1	0		Basalt		
11	15				6	1020	15	15		Giltston	<u> </u>	
34	13	12			7	10	W	-10	- <u>1</u> -	Granite	· · · · · · · · · · · · · · · · · · ·	
	17		7		6	1.50	71	15	й [—] '	Lineston		
36		- <u>-</u>			1	N 50		15	1	Shile		
17	15	$-\tilde{13}$	-5-		<u> </u>		DE.	20	1	Lisonite		
34	13	11			5	115	5.1	25	512	Chaie		
33					Í	N7	5E	25	í. – I	1.1 monton	ie	
40	10	ÿ	6		9	NEC		45	sēT	Quartzit	e	· · · ·
41	16	14	9		7	NJC NJ	DE.	10	<u></u>	There		
42	75	-22-	13		5	NIC	Ŵ	40	W	Basalt		
43	20	18	12		5	N60	DF:	20	5777	Volumite		
44	15	12	7		8	NU	5E	15	N.	Dolomite	··· •• •• •• •• •• •• •• ••	
45	14		6		7	- N70	55	- 30	51	Limiston	ю;	
46	14	11	7		7	N70	DE	0		Dolomite		
47	13		5		4	N4(	W	20	<u>,</u>	polomite		
48	19		b		6	36	5W	Ö		Eiltston	ie	
49	18	10	B		6	112	w	15	5	Bisalt		
50	13	10	9		5	:170	)E	15	S	Chert		
Till Matrix Composition Data: Color Tan Munsell Soil Color No. 2.57 5/4 Clay Mineral Composition												
Gra	1n-51z	e Dist	ributio	n	Expan	dable		. 1	1 2	hlorite	Carbonates	Cts./Sec
Sa	Sand Silt Clay			Clay Illite			plus	Calcite	Polomite			
L_19	9	48	ذذ	<b></b>	Mine	rals			Ku	olinite	6	
			-	6			8 7		16			

Figure 44 Data for analysis site Nl

138

Topography <u>B and C slopes on distal edge of moraine</u>

## Figure 45 Data for analysis site N2

	11	1 15	i		L _ 1150		1 20!		1:010:01	.e	1	
5	12	1 10	8	5	1:7	:::/	90		Cherr		LAV	
6	11	H	6	ζ		•.;	90		Lalonit	-0	LAV	
7	22	10	6			<u></u>	10	- Ir		<u>.</u>		
à		<u>├</u>			<u></u>	<u></u>		11.	1010411			
<u>⊢ ⊸</u>		<del>  <u></u> <u></u></del>		<u> </u>	بلج تلب	<u></u>	- 20'	·	( nore			
	-1/_	1 12	- 2 - 1	8	lisi	·	50.		1:01 OIL 1 C	.c		
10		1-14		7	اذ ا	ie .	4, 5	1	1.1.0			
11	13	10	10	5		:	453	,	Polonit	.e		
12	18	1 16	8	5	1	1	40	1	LOI WHET	e e		
13	33	8	6 1	6	1.85	514	20.		l'olomit	C C	1	
14	30	25	5	1	1.6.1	JC	40	1	(hile			
15	21	19	17	3	P.A.	1.7	0		Latonit			
16	H	6		4	NI/I	w	0			0	- <del>4</del>	
17	10						10					
16						<u>.</u>	<u></u>		1.1.111	one		
16	- 42-			<u> </u>			- 90		1010010	е	TIAN	
- 13							159	<u>v</u>	1:01 cm3 t	.e		
	18			5	<u>N75</u>	» <u>r.                                    </u>		⊻	1 ino locit	.c		
21	11-	1-14-1		8	131	W	_10:	<u> </u>	lo] omit	.c		
22		<u>18 I</u>	1	J	NSC	1:1	20!		Shale			
23	16	14	9	B	N 31	W	1.5	W	Lineste	one		
24	12	8	7	5	151	:	.0	;	l'elogit			
25	56	46	20	5	1125	E I	55	;	Cillete	one		
26	16				NT (			·	Limorte	000		
777-	10-					×			Delonit			
-36-		<u>├;</u> ;+	<del></del>			<u>"</u>	-;		O	<u> </u>		
- 20-	18			<u>4</u>	N/:	<u>N</u>	<u> </u>	<u>~</u>	OOm10	e		
	<u> </u>					W	_90		fhole			
- 30	_27	161	15	5	<u>N7(</u>	₩		<u>v</u>	_1 <u>10</u> 1 om i t	<u> </u>		
51	16	9	8	4	1 <u>: N</u>	): <u> </u>	0	1	1010 i t	CC		
32	18	9	9	6	N 30	)E: [	15	E 1	lolomit	ce	1	
33	42	31	23	5	N40	DE	Ū	- 7	Polomit	te in the second	- <u>-</u>	
34	14	10	9	5	NoC	DE:	15	N - I	Limeste	one		
35	31	20	13	6	NAC	1.1	-25	- T	Dolomit	-0		
36	30	16	15	2	N10		20	-	Silter			
37							-20	<u>-</u> +				
14						21/	<u> </u>	×+	- increy	L'OIQUILLE		
- 10					<u> </u>	<u>~</u>						
37	-21-				N?	₩ <u> </u>	<b></b> 6	N	_ <u>Silista</u>	pn <u>e</u>		
40	_ 32			5	NC;	W	<u>;0</u> 1		<u>[[o]omit</u>	.e		
41	21	16	<u> </u>				10:		l'clomit			
42	16	13	10	7	N80	M I	90		Chert		LAV	
43	25	16	10	6	N20	DE T	40	N	Delomit	te		
44	27	25	13	5	N20	DE I	60	N T	Shile			
45	21	101	12	9	334	DE T	- 0		Polomit	.e		
46	45	31	24	3	1110	30	40	<del>, 1</del>	Dolomit	e		
47	10		<del></del>		1:2:		- 30	in t	Chartza	te		
16					1100	r.1	- 70					
					1140			12	Share			
		- 28			N 30	<u> </u>		-+	( ranite			
50	18	9	6	6	<u>N15</u>	W	251	W	Shale			
Till Matrix Composition Data: Color Tan Munsell Soil Color No. 2.59 5/4												
	Clay Mineral Composition											
Grat	in-Size	e Distr	ibution	Expand	lable			C	lorite	· Carbonate	es Cts./Sec]	
Sar	nd 13	Silt T	Clay	]   C1a	ıy İ	1113	te		plus	Calcite	IUlonite	
1	.7	41	42	Miner	als	Kaolinite II I			10			
					3 83 14					·		
l										1		

Topography <u>B and C slopes on distal edge of moraine</u>

Shape

2

HII Pebble Data:

25

No

Dimensions

19

10

8

20

Analysis Site Number N2 Morphostratigraphic Unit Normal Location SW NE NE Sec. 14, T.24N., R.1E. Quadrangle Danvers

Elevation 800' Depth of Analysis 10' Type of Exposure Roadcut (new)

Strike

NEON

1185W

MESH

Dip

Shale

Diorite

Limestone

10:

0

224

Lithology

Misc.

Analysis Site Number <u>N3</u> Morphostratigraphic Unit <u>Normal</u> Location <u>NE SE SE Sec. 8, T.23N., R.4E.</u> Quadrangle <u>LeRoy</u>

•

Elevation 870' Depth of Analysis 6' Type of Exposure Roadcut (old)

Topography <u>B and C slopes on distal edge of moraine</u>

D	imensi	ons	Shape	Strike	Dip	Lithology	Misc.
27	17	10	6	NSW	105	Linestone	
78	27	71	5	11255	0	Dolumite	
12		9	5	NISW	MEGE	Linestone	
18	14	6	2	NIZOW		<u>Shaje</u>	
12	<u>. b</u>	6	<u> </u>	<u> </u>	0	olomite	
15	10	8	5	<u>N_/DE</u>	1.2.1	Linestone	
	- 10		6	<u> </u>	10 10	1 Greenstone	
<del></del>	<u> </u>	<u> </u>		105.1	1 150	Dolemite	
11-		;	<u> </u>	1.1.04	0	Li ae: Cone	
-22-1			<u> </u>	<u>Neo:</u>	10.7		
<u>~</u>	- 20			WC2N	2 SNW		
+2-1				E.		Liner Cone	
12 1	10			<u></u>	0	iolomite	
<del>-10</del> -1		┝━━╄़─┼	<u> </u>	450.1	1984	Limentone	
15 1	10	<u> </u>		NJON	40%i:	Shile	
			6	3 . 1.1	0	loloaite	
<del></del> +			8		55111	Dolomite	
<del></del>	- 7			<u>N854</u>	10:1	hue	
-18-1	-10-1	- 0 +		1706	2011	Siltstone	
- <u>18</u> -1	-10-1		<u> </u>	NESW	3556	Limestone	
- 4 - 1	- 8	4		N70E	1100	Polomite	
-14-+	- <u></u>	8	<u> </u>	<u>N656</u>	305	liolomite	
<del></del> +				<u>N5:</u> _	1-25	l'oloina te	
			<u>-</u>		- 95	0010m1te	
<del>~~</del> +	-16-1		<u> </u>		1	Polomite	
13-1	<u> </u>		<u> </u>	<u></u>	<u> </u>	1.1 mestone	
49-4		<u>8</u> -+		Neur.	+	Linescone	
+		<del>~ 8</del> +	<u>8</u>	N_3\ <u>04</u>			
+ +			<u>8</u>	<u>[409</u>	1-1501	polomite	
+2-+				<u>N251.</u>	90	polomite	LAV-S
			<u> </u>	<u>II</u>	1-25:5-	Silestone	
<del>~2</del> +							
	- 4-		<u> </u>	<u>N55M</u>	1556	(abbro	
				N207	1-1000	1.0100376	
				<u> </u>	1 10	Limestone	
	-12		<u> </u>	NOUL	- JONL -	Limestone	
-19-+				N	+- <u>15;</u>	1:0100126	
		;÷+		11111		Limestone	· · · · · · · · · · · · · · · · · · ·
+++++++++++++++++++++++++++++++++++++++						Granotiorite	••••••••••••••••••••••••••••••••••••••
-				NUUL	- 355	Siltstene	
10 +	-14-1			N	<u> </u>	Shale	
	-10-1			NSUL	20:W	Limestone	
			<u> </u>	N/UW	10.50	Linescone	
- <del></del> +	-44-1				1056-		
++++				NOOW	- COSE	Loionice	
<del></del> +	<u>u</u>			N		Limesrone	
++++		<u>P</u>		<u>L</u>	100	1010m1Ce	
	- 8				<u></u>	Chert	
<del>~~</del> +					408L	Siltstone	
				<u>L</u>	106	Chert	
trix	Compos	ition D	ata: Colo	or <u>Gray</u> t	Composit	ell Soil Color No. 1	UYR 5/2
		<u></u>		aute 1		alue Carbonal	Ces Cts./Sec
-+- [°]		20				plus Caleite	e polomite
<b>I</b>	22	<u> </u>		a13			1 27
-Size	3	Distr Dr 53	Distribution It Clay 53 30	Distribution Expand Distribution Expand Clay Cla SS 30 Miner 7	Distribution Tt Clay S3 30 Tt Clay Minerals T	Distribution Clay Clay Illite 53 30 Minerals Ka 7 81	DistributionExpandableChloriteCarbonaItClayClayIlliteplusCalcite5330MineralsKaolinite2278112

Figure 46 Data for analysis site N3

Analysis Site Number <u>N4</u> Morphostratigraphic Unit <u>Normal</u>

Location <u>NE NW NW Sec. 1, T.23N., R.2E.</u> Quadrangle LeRoy

#### Elevation 860' Depth of Analysis 5' Type of Exposure Roadcut (new)

Topography <u>B and C slopes on distal side of moraine</u>

<b>.</b>	D	Imens 1	ons	Shape	Strike	D	P	Lithology	Misc.
1	18	7	5	6	NBOE	10	E Basalt		
2	31	77	18	3	11701	1 15	6 Southte	one	W
3	12	11	5	5	NEDW	0	thert		
<u>.</u>	14		4	2	112/511	203	<u>Shale</u>		
<u>&gt;</u>	15	10	6	6	N20W	- 10	Polomit	c	<u>i.nv</u>
<u>-</u>	<u> </u>			5	NRSE	10	: Linesto	ne	ļ
<del>.</del>	-12	10	4	<u> </u>	110.7	40	<u>3 Basalt</u>		<b></b>
8	-18	14	+++++++++++++++++++++++++++++++++++++++	<u>_</u>	N UV	- 0	Chert		l
<del>7</del>			16-		11/00	45	5 10010:01	с	
Ť	12				11409				<b></b>
2		14					Dol Chit	<u>e</u>	+
3	16	14	10		NYOU			<u>e</u>	
ž l		- 17-		<u>5</u>	N850	1-33	Culo	e	
5	13	11			1000	1 20	1 Unionit		t
6	15	11	6	ξ	LILLINI				t
7	18	15	10		N759	45	Liorite		+
8	<u> </u>	6	6	6	NEOE	1 0	Linerto	ne	t
9	23	38	13	5	NGCA	50	SE Lineste	one	1
0	17	13	30	8	NIOI:	5	5 Basalt		t
1	14	11	6	6	1170W	15	Chert		1
2	15	10	9	5	N55W	90	Dolonit	e	W
3	22	12	8	٤	1185W	40	1: Dolomit	e	
4	21	14	3	1	N101;	20.	3 Shale		
5	16	ົມ	6	5	N156	10	S Linesto	me	[
é _	15	8	3	6	N4SE	5	NE Granite		
7	49	34	25	6	N40F	5	W Granite		W
3	$-\mathbf{n}$		4	E	N75E	10	6 Shale		
<del>.</del> 9		_25_		5	N40E	10	WE Granite		
)	30	20	9	77	N85E	10	L Chert		
	12		6	55	NUO!;	10	CLinesto	ne	W
	_12_	6	2		N75W	50:	S _Shule_		
2		22		6	N 35W	25	NW Dolomit	e	
-	18	<u> </u>	8	6	N80W	<u></u>	Limesto	ne	
<u>}</u>		<u> </u>		6	<u>N10E</u>	20	U Dolomit	c	
<u>-</u>	- 11 +			8	<u>10E</u>	- 25	<u>M</u> Granite		L
		-64	5	6	NS5W	0	Limesto	ne	
<u> </u>	-20+	-15	-91		N70E	-20	W Dolomit	e	
-	- 20-+	-26-	-18	<u> </u>			Dolomit	c	
-	- 24 +	-19-1		<u>è</u> +	N20E	1 12	Basalt		W
5-1		-10-1		<u>-</u>	N8SE		<u> </u>		
-	-46+	-4	-24-1		E	45	<u>1010m1t</u>	<u>c</u>	
í-t					N20W	1 10	S B15410		
; - I					NOUN	$+\frac{20}{10}$		e	
		╶╌┽┥	-17-1		N	t 10		e	
5	- 25 +	121	$-\frac{10}{10}$		NA W		J Durk C.	vetalling (2)	
5	23	16	15		E	<del>1 7</del>	Lineste	Jacarrine (1)	
9 1	19	16	13		N75W	25	J Dolomit		
	20	19	6		NSOE	50	NE Dolomit	. <u>u</u> .o	
	atrix	Compos	Ition 1	Data: Cold	or Tan	Hu	nsell Soll	Color No2.	SY 5/4
					Marris			·	
	n. C1 ==	hiet-	Thirt		y mineral	Compo	sition		
5.0	-51Ze	11501	1DUCIO		apre .		Chiorite	Carbonates	Cts./Se
301	iu s					ute	plus	Calcite	Dolomit
	, ,	40	34		CALL AL		KAO1171T#	1 10 1	

Figure 47 Data for analysis site N4

### Figure 48 Data for analysis site BD5

o.	D	imensio	ons	Shape	Strik	e D1	D	Lithology	Hisc
1	H	6	5	6	1150	150	CI	iert	
2	19	10	$t_{11}$	5	1500		Te	olumite	1
3	10	7	6	C C	Neon	10:	M La	mestone	
4	8	6	5	C,	UC3N	U	11	plomite	
5	10	6	6	3	NZOC	20:	; (1	iert	
6	24	17	12	5	NUSW	10.0	tal Li	mestone	
7	24	19	3	].	NESE	60!	1 11	ule	
8	15	12	5	2	NGUM	255	E Ch	iert	
9	12	)]	6	8	N85W	15.	/ Li	mestone	
10	32	20	14	6	N257	101	E De	plomite	T
11	11	8	6	6	N4 59	40 [.]	U Li	incstone	
12	11	8	6	6	N554	401	ાં ૬૫	artzite	
13	15	13	10	7	N55./	30:	1.1 110	olomite	
4	8	6	2	1	N554	105	E Si	ltstone	1
15	55	33	14	1	NSW	35		nale	L W
E	32	22	3	1	NSOE	551	IN St	nale	T
17	20	13	11	6	11551	10:	A Po	olomite	1
8	9	4	1	6	16355	201	IE EF	1 ile	
19	10	6	5	6	N/UL	25:	M 100	olomite	
20	15	12	7	5	N651	555	W 31	nsult	1
21	30	1.8	14	9	N70.1	35:	M 1.i	imestone	
22	10	6	6	6	HSOF:	20:	ज त	iert	
23	12	7	5	6	NSCE	25:	W De	olomite	
24	7	5	4	5	N25W	0		untrite	
15	9	<u>{</u>	5	9	HUIM	311	W Qu	urtzite	
26	29	21	12	<u>ئ</u>	NESW	40		ltstone	
27	1 8	4	4	υ	hise	1 250	i yı	ancaite	1
28	10	9	В	5	N701;	10	: I'c	olomite	1
29	6	5	3	5	NSC	0	I.i	inestone	
30	14	9	8	6	N41.4	201	N I FI	lint	1
31	20	16	12	7	11251;	0	CI	iert	T
32	40	27	23	5	NIOW	301	l De	olomite	S
33	]4	13	2	2	N	10	1 51	nile	
34	12	8	5	9	NGOW	105	E l.i	mestone	
35	9	7	4	6	N25W	455	ELi	inestone	
36	20	19	14	5	N5W	35.	Do	lomite	
37	17	11	7	6	NSSE	355	W Cr	nert:	
38	7	5	3	5	NSOE	0	Di	orite	1
39	14	7	5	6	NION	451	L Do	Jonite	T
10	10	5	4	6	N40E	205	W Li	mestone	
1	11	8	5	6	NGOE	205		listone	
2	14	8	8	6	N45W	10:	N Ci	iltstone	W
13	15	12	8	7	N4 51.	201	M CF	nert	
14	30	7	7	6	NISE	150	i [,i	inestone	
15	18	13	Ŗ	6	N 50.1	51	M Lo	olomite	
6	12	10	8	5	NEOW	0	1.1	imestone	1
17	9	8	4	7	NSON	201	N St	ale	
8	20	12	8	2	N10C	101	[S]	nale	
9	15	12	4	1	HUBH	404	<u>دا ا</u>	Jale	W
50	72	50	37	3	NJOW	104	1 54	uristone	
m	Matrix	Compos	sition	Data: Col	or <u>Tan</u>	Hu	nsell Soil	Color No. 2	.5Y 5/4
C=-	10-01-	<u> </u>	athurt-		dable I		Chlorito	Carbonate	s Ctern
618	n-51Z				uante	7114		Carbonate	1 101 m
- 38		3110			ay I	111116	Kaoliaiee		1,010,00
<b>1</b>	<u> </u>	44	5/	I L mine					1 41

Topography B, C, and D slopes on interlobate moraine

 Analysis Site Number __BD5 __Horphostratigraphic Unit __Bloomington

 Location __SE_SE_SE_SE_C. 22, T.23N., R.6E. __Quadrangle __Gibson City

 Elevation __840' __Depth of Analysis _5' __Type of Exposure __Roadcut (old)

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Analysis Site Number <u>BD7</u> Morphostratigraphic Unit <u>Bloomington</u> Location <u>SW NW NW Sec.24, T.23N., R.6E.</u> Quadrangle <u>Gibson City</u> Elevation <u>B20'</u> Depth of Analysis <u>5'</u> Type of Exposure <u>Rondcut (old)</u> Topography <u>C and D slopes on interlobate moraine</u>

<b>.</b>	D	Imensio	ons	Shape	Strike	Dip	Lithology	Misc
Ì	44	36	32	5	N70V	5500	Granite	
2	23	15	114		14.77	KIOL I	Linestone	
3	25	13	12	6	NALIN	5.1	filtstone	
Ť	15		8	G	11551	10	Polemite	
Ś	24	1.	q	4	11417.4	SUE.	Limestone	- T
É	10	11	1		1:109	1 10:	Lipestone	
7	- <del></del>	111	10	<u></u> ζ	11549	90	Quartzite	
<del>。</del>	22				1201	$\frac{1}{0}$	Chile	
8	- 10				1150	2011	Linestone	
3		24			11.5:	1 1	liolomito	
10				<u>'</u>		6	Chalo	
	25	24	<u> </u>		12.01.	1 100	Chulo	
2	19	12-	<u> </u>	<u> </u>		10.1	balconite	
13	24	18	15	<u> </u>		401.4		
4	13	12	10		16531	4	10104110	
15	14	12	10		11.07	Sill	<u>90190110</u>	
16	15	10	1	· · · · ·		1-158	1010m100	
17	27	25	11	6		41.11	Polonite	
8	30	25	14	5	1130.1	<u>- 5936</u>	<u>Chile</u>	
9	10	2		<u>9</u>	;;45W	<u>1000</u>	Chert	
10	9	7	3	)	1145W	0	Linestone	
21	11	10	6	7	11304	4/11/	Limestone	
2	10	5	S	E:	11SE	<u> </u>	Linestone	
23	35	13	9	5	1155	L 200	liolomite	
24	23	15	5	1	NESE	1 204	Shue	
5	19	6	4	6	1124	0	Limestone	
15	17	10	3	7	115.0%	3522	Chala	
27	21	18	13	7	11201	401	Dolomite	
8	11	10	8	5	N	203	Chert	
14	13	12	9	5	Ē	104	Dolomite	
50			15		NPOW	1 0 -	Litestone	
<u>ii</u>		16	8	6	1450	20:	Chert	
12	14	$-\frac{10}{10}$	~~~~	6	14 5 34	1 10.0	Limertone	
11	- 12-	- 10			1354	10 1	Chart	
14	<u> </u>				11 5 251	t	(hont	
24		<u> </u>				<u>+</u> -	- <u>Cherc</u>	
<u>~</u>		-12-1	<u> </u>	<u>&gt;</u>	NSOL	1 10.	Uimestone	
20		<u> </u>	<u> </u>	6	<u></u>	101	- indie	• 🚽
57	13		6	5	<u>NSI</u>		Shale	
38	25	18	_15	<u>S</u>	W		Shale	
59	15	_11_	8	5	<u>180F</u>	• <u> </u>	Dolcmite	
10	_24			5	LISE	403_	Dolomite	
1	_15_	_14	9	5	<u>NSE</u>	10:	Polonite	
2		6	3	1	NSE	100	Shale	
13	2	8	6	6	N15E	L0	Dolomite	
14	9	8	5	<u>8</u>		<u> </u>	Quirtzite	
15	25	18	17	3	N874	0	Quartzite	
16	18	12	_ 9	9	M5CE	0	Linestone	
17	13	10	8	8	NGOL	355W	Dolomite	
8	11	7	4	2	N25E	30:11	Dolomite	
9	15	11	8	6	N3CW	-60%W	Polomite	
50	40	25	21	6	NISW	40E	Dolomite	
m	Matrix	Compos	ltion	Data: Col	or <u>Thin</u>	Muns	ell Soll Color No.	2.5Y 5/4
Gno	10-01-	Die+	othutto		dable		Chlorite Cyrbonst	OS CIE 7
Gra	-51Z		CIDUCIO	비ᅜᅇ	uapre .	400	chippine Carbonat	
		40	Clay Clay Illite plus Calcite D				1 10100	
	/ I	44.11						

Figure 49 Data for analysis site BD7

Analysis Site Number <u>BD7</u> Norphostratigraphic Unit <u>Bloomington</u> Location <u>SW NW NW Sec.24, T.23N., R.6E.</u> Quadrangle <u>Gibson City</u> Elevation <u>820'</u> Depth of Analysis <u>5'</u> Type of Exposure <u>Roadcut (old)</u>

Figure 49 Data for analysis site BD7

Kaolinite

Minerals

Topography C and D slopes on interlobate moraine

HII P	ebble	Data:						1		1		
No.		Dimensi	ons	s	hape	Str	ike	D	lp	L	thology	Misc.
1	11	8	3		6	N 31	W	10	11.1	Bas	ilt	
7	16	13	2		5	164		1 0		Loi	cmi t'e	
3	19	1 8	7		6		<u></u>	1- <u>20</u>		<u>i ha</u>	10	1.4V
4	22	10			<u></u>	Hot	<u>w</u>	<u>↓ 10</u>	<u> </u>	<u>tha</u>	10	
	30	- 25	13				·	42	<u> </u>		rt estono	
<u> </u>	45		1 <u>4</u> 8 -		<u> </u>		.w	- 10	1	7.10	ult	
		12	14		- 2	112	14	1-50	1		ouite	
- ô		14			5	115		10	1.1	Lin	onite	
10	28	22	10		<del>- ć</del>	115		90		Bas	alt	LAV
11	14	1 10	9		7	1185	50	25	Ē	Lin	estone	1
12	21	12	8		6	17	SN .	20	<u>.</u> ]	bol	omite	
13	18	5	5		6	.In	E .	10	Ω	Qua	rtzite	1
14	15	10	9		6	151	чГ. —	90		the	rt	LAV
15	11	7	6		4,	1170		40	11.4	Lita	entone	
16	15	12	8		<u> </u>		<u>.</u>	35	<u>.</u>	100	<u>onite</u>	
17	J4	1 11	6		5	1125	<u></u>	60	1	Lan	estone	
18	12	10	8		5	ha(	<u></u>	35	<u>11. –</u>	1:01	ciai te	
19	15	11			<u></u>		<u> </u>		<u>110</u>		o <u>mite</u>	-+
20	1-14	1 10			<u> </u>	N/	<u> :</u>		<u></u>	<u> </u>	escone	
-22-	<del>L U -</del>	1	<u> </u>		<u></u>		<u> </u>	- 912	N	<u> </u>	0:01 (C	
27		+ 10-	<u> </u>			- 1170	<u></u>	- 60	<u>مانم</u> ۱۱			
24	<u>::</u> ₽-	+ <u>- <u>-</u></u>			A	1.70	)! <u>'</u>					
25		1 10			1	170	1:1	90		Lina	16	
20		H	2		1	N.21	21.1	10	NE:	i sha	10	1
27	41	25	17		6	N30		20	NE.	Gra	nite	W
28	15	1 10	10		7	<u>81</u>	)C	0		Che	rt	
29	28	17	13		6	150	).:	10	NE	101	omite	
30	39	19	17		6	N30	<u>M</u>	90		C01	onite	LAV
31	57	44	37		5		<u>.</u>	10	<u>لا</u>	<u>ro1</u>	omite	<u> </u>
32	12	10			8	N40	<u>w</u>	10	EE_	Qua	rtzite	
33	10	10	7		_6	130	<u>ic</u>	وتيا	£W	<u>001</u>	onite	- +
34	17	+ <u>11</u>	20			<u>N40</u>	<u>w</u>	10	N./	Lim	estone	
35	13	$+ \frac{10}{10}$	<u> </u>		<u> </u>	<u> </u>		- 30	NE.	<u> </u>		
- 17		+			-4	<u>1150</u>	<u> </u>	- 40	20	<u></u>		+
	- 40	$+\frac{2}{12}$			5	N 30	1J	50	CP	Ras	alt	
39	-12	+	6		6	111	<u></u>	15	N	Che	rt	
40	12	1 10	9		7	1150	!	10	S	Che	rt	
41	11	8	5		9	N25	50	25	NE	Qua	rtzite	
42	17	14	12		5	11/0	פר	03	51	1:01	omite	
43	15	10	9		6	lip(	)Ľ	20	NEL		omite	
44	53	37	10		1	1122	35	10	<u>_</u> M_	the state	]c	
45	13	2	7		6	N51		5	Ν.	<u>'.o'</u>	omite	
46		+ $ +$ $+$ $+$ $+$ $+$ $+$ $+$ $+$ $+$ $+$	6	ļ	_2	N23	<u>,</u> Е	-25	24	Lin	<u>esi one</u>	·
47	22	1_17_			_3	1170	<u>)</u> [	90		101	cmite	LAY
48	14	2			<u> </u>	112	£	30	NL_	<u>Sha</u>	<u>le</u>	
	- 22	+ 15	<u> </u>		<u> </u>	<u> </u>	×	<u></u>	<u></u>	<u> </u>		- <del> </del>
<u> </u>	1 16	<u> </u>	<b>S</b>		<u> </u>	<u>N</u>	<u>.</u>	40	nı.		<u> </u>	
TIT	Matri	x Compo	sition	Data	: Col	or <u>T</u> ay Min	n eral	Mt Compo	unse osit	In Soil C	olor No. 2	.5 <u>Y</u> 5/4
	ind T	511+	Clav	-	C1 C1	AV	m	ite		plus	Calcite	Lolonite
	<u>10</u>	36	15 Minerals Kaolinite R In				10					
<b></b>					5		8	2		13	<u> </u>	
L												

Figure 50 Data for analysis site BD8

144

 Analysis Site Number
 BDB
 Morphostratigraphic Unit
 Bloomington

 Location
 SW SW NM Sec.8. T.23N., R.7E.
 Quadrangle
 Gibson City

 Elevation
 B00'
 Depth of Analysis
 B'
 Type of Exposure
 Roadcut (old)

Topography B and C slopes on interlobate moraine

•

Figure 51 Data for analysis site BP7

Торо	graphy	B and	C slo	pes near e	dge of int	erlobat <b>e</b>	moraine					
HII F	ebble	Data:					· · · · · · · · · · · · · · · · · · ·					
No.	D	imensi	ons	Shape	Strike	Uip	Lithology	Misc.				
	70	53	28	7	E	5511	Granite					
2	15	10	6	7	N705	257	Linkstone	+				
3	28	25	3		1101	45N	Siltstone					
	17	<u>12</u>	12		<u> </u>	1 1500	Polonite	+				
	10				N65E	25.1	Chert					
7	17	10	6	<del>č</del>	214 (1.)	1010	Limestone					
8	16	11	7	4	16.00	200	lolomite					
9	22	9	9	4	ND 59	5511	Dolouite	1				
10	10	<u></u>	5	2	N70M	150	polomite					
11	19	10		1	11751:	30	Cuirtzite					
$\frac{12}{17}$	22	15_	- 7		NBOL		(h) (h) (0) (0)	+				
		-13-		<u> </u>	N454	3000	Gillero					
15	- 13-			5	N SOL	45NU	Dolomite					
16	24	13	10	6	N 70.1	0	Linestone	1				
17	17	16	10	1	NSSE	0	Bogonla					
18	20	17	12	5	NISE	20%	Doloaite					
19	_ 55	45	25	5	N2SE	0	Doloate					
20	20	12	- 12		1250		Lisistone	·				
	- 12			<u>-</u>	N35W	+- 90	Doloaice					
-23	- 22	$-\frac{69}{12}$	- 10	<u> </u>	N35E	0	Chert	<b>.</b>				
24	-25	10	14	6	11251.	0	Dolomite	<b></b>				
25	39	27	25	3	N204	S21M	Dolomite	L				
20	21		5	1	NEOW	0	Siltstone					
17	16	12	5	1	1125L	253		4				
26	- 25	18	-10		NESW	41/1	balcrite	+				
29		20	A	<u>+</u>	N70W		Granite	+				
- 30	-15-	10		<del>~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~~</del>	N40.	1 Sirw-	Siltstone	+				
32	14	12	- 9	8	1201	155	Dolomite					
33	28	15	13	5	NGOE	0	Quirtzite					
34	18	14	9	7	NESE	155W	Greenstone					
35	14	10	9	3	NGOW	5MV	Quartzite					
36	17	10 - 10		6	N55W		Chert					
$-\frac{57}{76}$		-19		<del>2</del>	940	+ + +	Tolsito					
	-12-			<u>6</u>	N304	1052	Fed Lamestone					
- 40-	- 20-	16		<u>5</u>	N350	305	Borite	· ·				
41	24	12	4	T	NEOW		Siltstone	<b>†</b>				
42	49	32	32	3	NIOL	205	lolcaite	1				
43	13	9	6	6	11856	0	lolonite	<u> </u>				
44	13	7		1	N 30E		Linestone					
45	10				N 30C	55W	10Joinite					
40				<del>-</del>		+		<b> </b>				
48	- 31	21		<u>6</u>	NSE	+- <u>5</u> š <u>s</u> -†	bolomite	<b>+</b>				
49	13	- 5			NSOW	1-200E	Linwstone					
50	19	10	2		NSE	TSN	Chale	+				
TIII	Till Matrix Composition Data: Color Yellow tan Munsell Soil Color No. 2.59 5/4											
Gra	In-5120	e Distr	ibutio	T Expan	dable	T in	lorite Carbonates	S Pts /Secl				
Sa	nd T 3	SIL	Clay		ay 11	lite	plus Calcite	Colomite				
24		36	40	Mine	rals	Ka	olinite 7	17				
					5 6	3	12					

 Analysis Site Number
 BP7
 Morphostratigraphic Unit
 Bloomington

 Location
 SE NE NE Sec.20, T.23N., R.6E.
 Quadrangle
 Arrowsmith

 Elevation
 830'
 Depth of Analysis
 5'
 Type of Exposure
 Pipeline Trench

Figure 52 Data for analysis site BP8

TP	ebble D	lata:						
	Di	mensio	ons	Shape	Strike	Dip	Lithology	Misc
Ĩ	27	20	15		<u> </u>	90	Siltstone	
2	1-19-1	13	21	1	1:4 W	4000	Chale	
3	1 9 1	5	4	G	N101:	0	Basalt	
4	10	6	5 1	6	N750	25W	Quartzite	
5	151	9	4	Ū.	N75E	0	Dolomite	T
6	151	91	7	Ē	NYUN	10.1	Polomite	
7	1 18	12	10	7	111 OW	1 0	Chert	
R I		17			1145.0	40:07	Dolomite	
ă	151	-12-1	$r = \frac{1}{6}$		N-101	10.11	polonite	-
<u></u>	1-15-1	-10			H15E	55	Shile	W
÷					1055	1 40	Siltstone	
	- 14 +		/ <del></del> +		NICE	1 60	Shile	-+
4	1-13-1		┝╾╶╪╼╋	<u> </u>	HISI'	100	Chalo	-+
3	1 14					+ 700-	Dolomita	
4	- 30	-25	<u> </u>	<u>`</u>	N75N	0000	POIONICE	
.5	7	4	3	<u> </u>	N55W	<u>↓                                     </u>	:)010m11e	
6	8	6	4	3	N755	0	Linestone	
7	8	7	5	É	N75E	0	Basalt	
8	13	11	2	3	NEOS	ENC.	Shale	
9	11	6	4	5	NSON	T 10N	Colonite	
10	12	11	4	2	N70E	90	(din)e	
5-1	121		7	5	NSIW	350	Dolonite	- T
5	1-16-1	101	9		N557	SOUL	Dolomite	W
5	1-14	-11 1			NSOE	TISW	Chert	
1	1-13-1				N701	1-55	Quartzite	
-	1					+-586	Dolouite	
<u></u>						1.250	Chant	
<u> </u>						+ 30.18		
<u>''</u>	1-32+			<u>}</u>		1-30 -	LICETU	- +
<u>,6</u>	18.1	-14-+			N551	<u>+-20</u>	1.1 mos cone	
19	1_21_1	_16	14	5J	N <u>50W</u>	1_9-	Dolomite	
<u>šo</u>	10	10	4		N85.V	- 20		
51	21	14	]2	7	N75W	40NN	Dolomite	
32	15]	10	5 1	2	NEOF	<b>]_9</b> 0	Shale	
33	161	91	6	6	N2OE	T JON _	lolomite	Ι
4	121	9	6	6	NSON	20111	Quartzite	- I
ŝ	591	17	6		NIOS	1-45N-	Shale	
2	1 5ú t	- 50 1	- ă		N20N	5N	Dolomite	-+
20-1	+ ::+	+			NODY	1 250	Limostone	+
*	+	<del></del>			N765	1 1 54	Limostono	· · • • • • · · ·
58	- 31	<del>+</del>	<u> </u>	<u>{</u> /	N/56	+-101 		· · • • •
23	121	-12				1 JUN	<u>Cherc</u>	<b></b>
10	12	10	<u> </u>		<u> </u>	20N	Chert	-+
11	16	10	5_+	6	N10!.	<u> </u>	Dolomice	
2	14	8	3		N 30W	0	Shale	
13	10	6	4 1	6	N25W	55 <u>5</u> E	Basalt	
14	18	15	6	2	N305	45:W	Dolomite	T
15	71	6	3	2	N2OE	35NE	Chert	Τ
16	12	9	6	5	N35E	10: W	Limestone	
7	14 1	8	4	<u> </u>	N4.5W	1 90 -	Basalt	-1
6	1-521		9		120F.	50SE	Shale	
i i	<del>  _ ; ;  </del>	- <del>15</del> +	<u> </u>		N151	1 90	Dolomite	
	1-12-1		<del>~ ~ †</del>		NA SU	1 1000	Chort	
50	1.5		L	0	114 514	1.500	VIII C	
m	Matrix	Compos	Ation I	Jata: Col	or Yellow	tan Muns	sell Soil Color No.	2.58 574
		- Diety	-Thut i or		dahla I		Chlorite Carbonat	ne 115./5
Gra	10-51ze	1 015 CL		4150		1400		
<u>5a</u>	nd i	5110	<u>L'Lay</u>		ay in	Lite		100100
	<i>.</i> .							

Analysis Site Number <u>BP8</u> Morphostratigraphic Unit <u>Bloomington</u> Location <u>SE SW SW Sec. 21, T.24N., R.7E.</u> Quadrangle <u>Sibley</u> Elevation <u>810'</u> Depth of Analysis <u>5'</u> Type of Exposure <u>Roadcut</u> (old) 

٥. '	D	imensic	ons	Shape	Strike	Dip	Lithology	Misc
Î	150	1357	76	7	NIOE	201	Dolomite	
7	172			7	1155E	1 10:34	Chert	
3	13	<u> </u>	6	6	NEOE	TIME	Polonire	
<u> </u>	<u> </u>			6	11251	100	Lolonite	
<u> </u>	19	1 14 1	10	3	REDE	<u> 6054</u>	Polomite	
<u>.</u>	19	1-14-1	10	6	<u>N</u>		<u> </u>	
<u>'</u> _'	1.18	<u>ب چر ا</u>	<u></u>	<u></u> '	11284	150	Polomite	
<u>8</u> _'	1 <u>11</u>	<u>↓₽</u> _	4-4-	5'	1150E	1 101	Limestone	
<u>_</u>	<u>↓_<u>↓</u></u>	1 21	1 <u>-</u> +	<u>\$</u> '	1250	+ 0	Shale	
<u>`</u> '	ᡶᡶᠧᢇᡃ	1-31	┢━━┾╾┥	'	1005	<u>+ 15; 4</u>	<u> </u>	
<u>,</u>	13	<u>↓</u>	<u>⊢ ∻</u> +	<u></u>		1-255		
<u>~</u> '	+ <u>-20</u>	1-14-1	<u>}</u> ;+	<u>لب</u>	14400	+	<u> </u>	
<u>}_</u> '	<u>↓ ;                                   </u>	<del>1.4</del>	<u>→</u>	¥	11251	بالإيب لم	Gaptero	
<u> </u>	175-	1-18-1	<u>⊢−<u></u><u></u>+</u>	<u> </u>	1-1620E	1-10-	<u> </u>	- <del> </del>
5_	-27	<u>↓</u>		<u> </u>	<u>:55F</u>	بنالب	Lindstone	N
<u>5</u> _'	<u></u>	1-20-1	<u> </u>	<u>+</u> /	1000	<u></u> ₩	Jule	
<u> </u>	سيدل	L B	-3-1		5700	1-90-	Siltstone	· _ <b></b>
<u>8</u> _'	ل_عد ا	1 74 1	<u> </u>	J	11651	90	thate	
<u>9</u>	ت_بد_ا	1 14	4_4	2	1154	1 20		
<u>o_</u>	سيديد بله	27	25	<u> </u>	ND04	20:	Dolomite	¥
<u>i_'</u>	<u></u> '	22	<u> </u>		<u></u>	1 201E	Ind Limestone	
<u>2</u> '	ب_ 2_ ]	لبت			N75:	20:		
3	16_/	10	8	3_1	NZ SW	<u>10:1</u>	Limestone	
4	22	<u>[]</u> 6	12	5	NSW	1451	Chert	<u> </u>
5_	<u>]]</u>	[]]0	<u></u> H	<u> </u>	<u> </u>	<u> </u>	Lolomite	<u> </u>
ē_'	19 /	[is]	4	·]	NGSE	1350	Shale	.Г <u>.</u>
ī_'	14	10	6	7	N75E	1 0	Polomite	1
ε_'	10 !	9	4	1	N250	105.		1
<u>9</u> _'	<u>' <u> </u></u>	7	5	6	N 70E	1 JONE	Limestone	····
0_7	15	13	6	5]	N204	1-401	Dolomite	
1_'	13	8	6	6	N	JON	Limestone	
2	1 10	6	4	6	N15E	30N	Quartzite	
3	17	12]	8	5	N55F	3054	Chert	· +
4	115	12	- 7 ]	6 ]	INSE	155	Polomite	-
5,	1-10	6	4	4	NISE	205	Shale	
<del>i</del>	1 36	26	171		N45E	355	Dolomite	-+
5-1	1-17-1	1121	-in 1		N 70E	205	Timestone	+'-···
<del>4</del>	+	1	6	- Á	NGSE.	TONL	Chort	-+
<u>6</u>	1-12-1			<u> </u>	NGOE	1-1SNE	Racalt	
<u>~</u> -1	+ 15-+				1000	1-172	- Pisour Ivolonito	
<u></u>	1	101			NAGU NAGU	100	Dolomito	
<u>+</u> ,	1 20 -	1			184 JW		Dolomito	
$\leftarrow$	1-5-1	► <del>**</del> +			N/30	167	10100100	-+
	1 2 1	+			N/SW NAGE	1-101	Limescone Delomite	
<del>~</del>	+++++++++++++++++++++++++++++++++++++++				NHUL MOGUI		POLONI CE	· +
ىرچ	+ + - + - +	+ <del>0</del> +			NZ SW	35.		
<b>ب</b>	1		+	<u></u> ,+	<u>h55L</u>	4UNE	Linestone	
<u>-</u>	+++++	1 <u>1</u>	+	+	NPOM	<b>↓</b> 0	Shile	_
<u>ف</u>	1-14-1	<u>+ 10</u> +		<u> </u>	<u>N</u>	258	Chert	
ي چ	┥╌╧╴┥	<u>↓</u>	<u> </u>	<u> </u>	N5L	305	Dolomite	
<u>o</u> ,	16	<u>111</u>	<u> </u>	3	NIOL	205	Polomite	
n 1	Matrix	Compos	Ition D	Jata: Colo	or Yellow t	an Muns	sell Soil Color No2	7.59 574
Gra	In-Siz	• Distr	Ibutior	Expan	dable 1	<u> </u>	Chlorite Carbonate	as CLS. /S
5,	ind T	the T	Clay	HITCI	10 I III	11+0	nius Calcite	The curi
	·***	<del>,</del>	38	-   Mine			Paulinita ?	+ 1000 mm
4.	. 4	<b>7</b> 0 .	•••	••••• ·	1072 1			<u> </u>

Figure 53 Data for analysis site BP10

Figure 54 Data for analysis site BP11

0.	D	Imensio	ns	Shape	Strike		Lithology	Misc
T	30	19	6	1	NSOW	SULE	Ciltstone	W
2	10	7	5	6	NSOL	0	1010/100	
3	נו	5	3	4	N20E	300	Dolomite	1
4	13	]]	- 9	7	NISE	20:1	Colomite	1
5	]]]	?	1	2	NSOW	2508	Chert	
6	10	9	7	7	N55W	350 <u>6</u>	Chert	
7	18	15	14	3	10.02	351:	Chert	
8	15	13	9	5	NGOW	1056	lolomite	
9	12	7	6	4	NIOS	103	Lolomite	
10	12	10	5	2	N4 0E	1085	Polcmite	
1	28	- 15	15	3	NLOW	90	Polomite	
2	28	4	9	2	N25M	15NW	Linestone	
3		7	3	3	N7GN	1000	Shale	
4	17	14	2	7	N4 05	15:14	Lolomite	
15	311	32	17	5	<u>111 S E</u>	30%	Linestone	
.6	14		8	6	NUCE	N. 10[	Folonite	
2	10	5	5	6	8306	3011:	l'oloni te	
8	12	ع	5	6	NCOL	0	Hisilt	
9		6	4	3	H	<u> </u>	Limistone	
0	32	16	14	E	NSOR	<u>100W</u>	Basilt	
1	16	10	5	1	N150	25N	Lolonite	W
2	22	16	_ 4	1	N1Gi:	255	Grinite	W
3	_23_	18	9	5	N45E	_ 305	: olonite	
4	9	6	5		N150	105	<u>Dolomite</u>	
2	12	9	4		N25U		Limestone	
0	2-4	-21	4	5	<u>N</u>	<u> </u>	Quirtzite	
<del>.</del> –					11.0.	10:11	<u>"halo</u>	
e	$-\frac{21}{3}$				<u>N650</u>		Chert	
2-	-20-		-10 +	<u>6</u>	N251:	15NC	Nolomite	
	- 20	-44	-2+		N SCW	1011	Dolomite	
1	8	<u>-</u>	- 5 +		N10.4		Red Limestone	
2	-10-1	<u></u>			N40L		L'CIOMI CE	-+
A	- 15				NAUE		1º01chilte	LAV
-						- <u></u>	Chert	
6	┝╼┽┽╺┽			<u>-</u> <u></u>	<u>N 30.</u>	JUNE	Colomite	
3	- + - +		<u> </u>		N/5W	405E	Doloaite	
6			- 2+		NICE		Chale	
9					N		Cuircaice	
<del>6</del>		- 14				- JONW	<u>1'010m1te</u>	
ř –				<u>P</u>			Chert	
2				<u> </u>	NA00			-+
1	17			<u>-</u>	N 7 C 2		Cilletone	
ž l	16				NGSW		Toletto	
5	16	-10	$-\frac{1}{2}$		NGSE		Thile	
6	15	10				- 30 -	Folomite	
7	10			<del></del> -+	NIOC	-iow+	Shale	-
8	12	10	- 1		NOSE	TONE	Shale	+
9	14		- 6	- 6	NGOD		Linestone	
0	- 13	- 9			N40W	25NW	Chert	+
11	Matrix	Compos	Ition D	ata: Colo	or Yellow ta	n Munsel	1 Soll Color No. 2.	5Y 5/4
Gra	in-Size	Distri	Ibution	Expand	lable	l Ch	lorite Carbonate	es Cts.7S
Sal	na S	110	Clay		y 111	te	plus Calcite	Poloni
20		42	38	I I Miner	als I	I K30	linite   12	1 19

Topography C and D slopes high on interlobate moraine

 Analysis Site Number BP11
 Morphostratigraphic Unit
 Bloomington

 Location
 NW NE NW Sec. 12, T.23N., R.GE.
 Quadrangle Gibson City

 Elevation
 850'
 Depth of Analysis
 10' Type of Exposure
 Rondcut (old)

Eleva	ition _	750'	Depth c	of Analysis	23' T	ype of E	xposure <u>R</u>	padcut (new	<u>ı)</u>
Topog	graphy	? (51	tratigra	aphically :	lower till	l unit, t	op till see	E10)	
511 P	ebble I	ata:							
No.	Di	Imensio	ons	Shape	Strike	Dip	Lit	hology	Misc.
1	13	8	7	6	W) (W	5N	1010	nice	
- 2	27	12-		2	NSJ.				LAV
- 4	13	9		6	N25W	30PM	Gran	itc	
5	13	7	5	4	1:35W	309E	Dolo	pite	
3	16	11	8	3	HZOW	250E	1;Q10;	110	
	- 12-	10	- 25-		N15W	455			
- ⁰	37	13	12	4	NI SON	20:11	1.1110	aite	
10	36	15	10	5	1750	45E	Line	tone	
11	22	36	10	5	N/2 C M	0	Line	stone	
12	13	11	10	3	NEW	0	<u>Silt</u>	tone	
13	11	10		<u> </u>	NEW	25:5	Gree	nstorie	W
19		16			N I I		Gabb		
16	23	14	10	5	N5:.	401	1010	nite	
17	24	- 23	11	8	11257	10.1/	Dolo	mite	
18	24	13	6	1	NICW	704	laine.	stone	
19	18	15	14	5	N5.1	455	Basa	15	
20	30				NECA		<u>101</u>	nite	
22	20	16	10	5	NASW	2011	1.010	mit:o	
23	21	17	12	7	NSUW		Cher	t	
14	19	13	9	7	N 301:	35NE	[)010	nițe	
25	13	10	6	6	N 35W	30CE	1)0101	mite	
26	- 22 -	$\frac{17}{16}$	12	5	N 30W	<u>35NM</u> _	Sand	stonc	
		18-		<del>[</del>	NAUW		Shale	°	
29	11	7	- Š	6	NZSW	0			
30	14	12	6	7	NZCA	200	Dolo	nite	1
31	20	13	9	6	N55W	15NW	1.122	1	
32	15	9	8	6	NIOW	101	pera	<u>lite</u>	-+
- 33		-23-			N S SW		<u> </u>	lt	-+
34	1 10			4	N75./				
36	12-	<u> </u>	6	5	NSSE	25NE	Shale		
37	12	10	6	5	N55W	15SE	Gran	ite	]
38	13	10	8	5	NASW	90	1)0101	nite	LAV
39	$\frac{1}{14} - \frac{1}{14} - \frac{1}{14}$	10			N 20.7	-1-12NM	+ pojo	nite	
40	110-	$\frac{11}{12}$	20		N 3CH		Chor		1. AV
42	115	11	8	6	NSCN	2551	Dol or	aite	
43	27	15	13	4	N104	30N	Polor	nite	
44	20	12	7	2	N35W	35SE	Dolor	nite	+
45	14	<u>↓ ⊥↓</u>	4	2	N405	- <u>305M</u>		nite	+
1 40	┨-┿╤	+				150	Licior Rico		
48	114	1-10-	8	Ŭ Ŭ	N 35.N	0	Gran	i t:e	1
49	27	22	8	1	N15W	90	Sands	stone	1
50	38	20	12	6	N 30.1	355W	Shale	2	
TIII	Matri	Compo	sition	Data: Col	or <u>Red b</u>	rown Muns	sell Soil C	olor No	7.5YR 5/4
Gr	ain-Siz	ce Dist	ributic	n Expan	dable .		(hlorite	Carbonato	es Cts./Sec
	and	41			ay 1 rals	····e   ,	(aolinite		24
1	·		JJ4		4	70 1	16	L66	
1				E	·····				

Figure 55 Data for analysis site EB2

Analysis Site Mumber <u>ER2</u> Morphostratigraphic Unit <u>Euroka</u> Location <u>NE SE NE Sec.3, T.25N., R.1W.</u> Quadrangle <u>Danvers</u> •