**Sample Selection**

Samples were selected based on MgO wt.% content (MgO wt.% > 8.5) to identify those that were initially closest to primary compositions. Trace element diagrams of these samples were then identified that match Type III magmas from the EARS (Rooney, 2017) to be used for further analysis. Phenocryst assemblages of this subset of samples were identified according to those reported in (Chiasera et al., 2018) to determine which mineral phases were fractionating and thus which needed to be corrected for to determine primary compositions. Samples selected using this processes and their chemistry are listed in Table A1. Sample chemistries for analysis from Akaki (Rooney et al., 2014) were chosen based on similar criteria as those from Galema; however, phenocryst assemblages were not determined and it is assumed that only olivine has fractionated from these samples. Akaki sample chemistries chosen for this work are listed in Table A1.

**Table A1.**

|  |  |  |  |  |  |  |  |  |  |  |
| --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- |
| Sample | **3030** | **3031** | **3042** | **3061** | **3063** | **3077** | **2001** | **2004** | **2006** | **2029** |
| SiO2 | 45.84 | 45.74 | 45.41 | 45.79 | 44.81 | 46.54 | 44.98 | 43.98 | 44.34 | 44.89 |
| TiO2 | 2.48 | 2.53 | 2.78 | 2.52 | 3.04 | 1.98 | 2.11 | 2.26 | 2.18 | 1.58 |
| Al2O3 | 15.57 | 16.03 | 14.63 | 15.88 | 14.24 | 13.47 | 13.63 | 12.68 | 13.3 | 12.84 |
| Fe2O3 | 1.77 | 1.77 | 2.36 | 1.74 | 3.27 | 1.61 | 11.86 | 12.00 | 12.00 | 11.29 |
| FeO | 10.56 | 10.74 | 9.98 | 10.44 | 9.44 | 9.39 | 10.67 | 10.80 | 10.80 | 10.16 |
| MnO | 0.19 | 0.19 | 0.18 | 0.18 | 0.20 | 0.18 | 0.17 | 0.17 | 0.17 | 0.17 |
| MgO | 8.25 | 8.07 | 8.61 | 7.94 | 8.87 | 11.03 | 12.11 | 13.61 | 12.41 | 13.76 |
| CaO | 10.17 | 9.79 | 10.34 | 9.68 | 10.12 | 10.25 | 11.32 | 10.59 | 10.88 | 10.96 |
| Na2O | 2.93 | 2.98 | 3.04 | 3.10 | 2.88 | 2.51 | 2.39 | 2.24 | 2.14 | 2.10 |
| K2O | 0.65 | 0.67 | 0.44 | 0.70 | 0.82 | 0.62 | 0.61 | 0.44 | 0.57 | 0.59 |
| P2O5 | 0.27 | 0.29 | 0.33 | 0.30 | 0.48 | 0.29 | 0.36 | 0.35 | 0.42 | 0.30 |
| Cr2O3 | 0.02 | 0.02 | 0.04 | 0.02 | 0.04 | 0.11 | 0.10 | 0.13 | 0.11 | 0.12 |
| Sc | 29.27 | 27.77 | 29.48 | 28.20 | 30.61 | 31.09 | N/A | N/A | N/A | N/A |
| V | 324.75 | 315.56 | 315.48 | 317.90 | 305.03 | 264.48 | 269.00 | 289.00 | 258.00 | 213.00 |
| Co | 56.08 | 56.49 | 52.17 | 55.43 | 55.60 | 55.50 | 676.00 | 921.00 | 746.00 | 822.00 |
| Ni | 84.03 | 71.00 | 127.93 | 70.83 | 129.74 | 202.44 | 256.00 | 339.00 | 303.00 | 375.00 |
| Rb | 9.79 | 10.19 | 7.95 | 11.57 | 17.48 | 10.54 | 10.00 | 9.00 | 7.00 | 9.00 |
| Sr | 561.97 | 602.12 | 522.43 | 586.83 | 584.99 | 394.92 | 533.00 | 505.00 | 546.00 | 437.00 |
| Y | 22.96 | 22.65 | 25.39 | 23.09 | 30.12 | 22.98 | 23.20 | 20.90 | 25.00 | 20.10 |
| Zr | 135.87 | 141.45 | 175.14 | 147.32 | 290.70 | 127.86 | 99.00 | 93.00 | 94.00 | 89.00 |
| Nb | 21.58 | 22.51 | 28.23 | 23.58 | 39.08 | 19.66 | 32.90 | 30.20 | 30.60 | 23.40 |
| Cs | 0.13 | 0.11 | 0.60 | 0.13 | 0.21 | 0.06 | N/A | N/A | N/A | N/A |
| Ba | 212.11 | 218.16 | 269.58 | 227.44 | 387.25 | 187.42 | 247.00 | 193.00 | 287.00 | 290.00 |
| La | 18.09 | 18.50 | 24.00 | 19.42 | 34.04 | 15.12 | 22.20 | 19.50 | 23.20 | 16.70 |
| Ce | 38.57 | 39.91 | 51.67 | 42.07 | 70.53 | 33.85 | 41.80 | 38.30 | 42.20 | 35.70 |
| Pr | 4.99 | 5.14 | 6.39 | 5.54 | 9.05 | 4.47 | 5.46 | 4.98 | 5.81 | 4.35 |
| Nd | 21.84 | 22.41 | 27.05 | 23.45 | 37.42 | 19.56 | 22.80 | 20.90 | 24.60 | 18.40 |
| Sm | 5.03 | 5.13 | 6.04 | 5.27 | 8.17 | 4.62 | 4.86 | 4.45 | 5.17 | 4.01 |
| Eu | 1.77 | 1.75 | 1.96 | 1.80 | 2.62 | 1.57 | 1.61 | 1.49 | 1.68 | 1.37 |
| Gd | 5.10 | 5.08 | 5.92 | 5.23 | 7.75 | 4.90 | 4.62 | 4.38 | 4.96 | 3.84 |
| Tb | 0.74 | 0.74 | 0.85 | 0.77 | 1.10 | 0.74 | 0.71 | 0.64 | 0.76 | 0.62 |
| Dy | 4.40 | 4.44 | 4.96 | 4.55 | 6.28 | 4.50 | 4.10 | 3.77 | 4.34 | 3.60 |
| Ho | 0.83 | 0.84 | 0.92 | 0.86 | 1.15 | 0.84 | 0.81 | 0.73 | 0.86 | 0.68 |
| Er | 2.28 | 2.29 | 2.50 | 2.35 | 2.99 | 2.32 | 2.10 | 1.87 | 2.20 | 1.93 |
| Tm | 0.31 | 0.31 | 0.34 | 0.32 | 0.39 | 0.32 | N/A | N/A | N/A | N/A |
| Yb | 2.00 | 1.98 | 2.11 | 2.01 | 2.42 | 2.06 | 1.90 | 1.73 | 1.98 | 1.67 |
| Lu | 0.28 | 0.28 | 0.30 | 0.29 | 0.35 | 0.29 | 0.28 | 0.25 | 0.29 | 0.24 |
| Hf | 3.58 | 3.71 | 4.44 | 3.78 | 6.89 | 3.28 | 3.19 | 2.64 | 3.17 | 2.46 |
| Ta | 1.34 | 1.41 | 1.70 | 1.49 | 2.52 | 1.23 | 1.96 | 1.89 | 1.91 | 1.57 |
| Pb | 1.75 | 1.78 | 2.45 | 2.16 | 2.72 | 1.43 | 1.18 | 0.98 | 1.17 | 3.06 |
| Th | 1.56 | 1.63 | 2.20 | 1.69 | 3.16 | 1.29 | 2.41 | 1.98 | 2.14 | 1.93 |
| U | 0.43 | 0.45 | 0.53 | 0.48 | 0.84 | 0.37 | 0.53 | 0.47 | 0.48 | 0.44 |

**Conditions for running Petrolog v. 3.1.1.3. software**

Initial conditions for back-correcting samples to primary compositions using Petrolog v.3.1.1.3 software (Danyushevsky and Plechov, 2011) follow those detailed in El-Rus and Rooney (2017) except where indicated.

Partition coefficient values for trace elements in olivine, plagioclase and clinopyroxene were taken from El-Rus and Rooney (2017) except for Ni in olivine which utilized the value determined by Kinzler et al. (1990), as indicated in Table A2.

The reverse crystallization program in Petrolog was run using an initial pressure set to 7 kbar and oxidation state set to constant Fe2+/Fe3+ = 0.16 (QFM). These values were determined using previous models of magma evolution via Excel MELTS (Chiasera et al., 2018). Petrolog samples were back-corrected to MgO = 15.5 wt.%. Olivine was allowed to fractionate through the entire run and clinopyroxene fractionation was stopped at MgO = 10 wt.%. Clinopyroxene phenocrysts are not observed in Galema samples with MgO > 10 wt.% (Chiasera et al., 2018). Spinel was not included in the reverse crystallization calculations; however, primary conditions (Cr>750ppm, Ni>320ppm, Ni/MgO>25 (El-Rus and Rooney, 2017) ) were estimated to be achieved with <1% spinel addition through manual calculations. This manual addition of spinel had little to no effect on the trace element concentrations in the corrected chemistries.

As the Petrolog software does not determine when a calculation has reached primary composition, several of the calculated chemistries from end of each run were entered into the model of Lee et al. (2009) to determine which were primary. The chemistry that yielded the fewest possible necessary back-correction iterations in Lee model was chosen to be primary for each sample that was back-corrected in the Petrolog software. Akaki samples were not utilized in the Petrolog software and were back-corrected to primary compositions using the Lee et al. (2009) model. Utilization of the model of Lee et al. (2009) is detailed in a later section. Results of these back-correction calculations are in Table A3.

**Table A2.**

|  |  |  |  |
| --- | --- | --- | --- |
| **Element** | **Olivine** | **Plagioclase** | **Clinopyroxene** |
| **Sc** | 0.25000 | 0.01600 | 2.28000 |
| **V** | 0.09000 | 0.03500 | 0.60000 |
| **Co** | 5.21000 | 0.04200 | 0.95000 |
| **Ni** | (Kinzler et al., 1990)**\*** | 0.08900 | 2.84000 |
| **Rb** | 0.00032 | 0.02070 | 0.00470 |
| **Sr** | 0.00050 | (Blundy and Wood, 1991) | 0.12830 |
| **Y** | 0.01310 | 0.00800 | 0.49200 |
| **Zr** | 0.00068 | 0.01080 | 0.12300 |
| **Nb** | 0.00005 | 0.00140 | 0.05000 |
| **Cs** | 0.00032 | 0.00390 | 0.00100 |
| **Ba** | 0.00032 | (Blundy and Wood, 1991) | 0.00068 |
| **La** | 0.00003 | 0.05170 | 0.05360 |
| **Ce** | 0.00010 | 0.04000 | 0.08580 |
| **Pr** | 0.00027 | 0.03845 | 0.13655 |
| **Nd** | 0.00042 | 0.03690 | 0.18730 |
| **Sm** | 0.00110 | 0.01760 | 0.29100 |
| **Eu** | 0.00075 | 0.96060 | 0.32880 |
| **Gd** | 0.00120 | 0.01340 | 0.36650 |
| **Tb** | 0.00130 | 0.01130 | 0.40430 |
| **Dy** | 0.00140 | 0.00900 | 0.44200 |
| **Ho** | 0.00720 | 0.00710 | 0.41500 |
| **Er** | 0.01300 | 0.00500 | 0.38700 |
| **Tm** | 0.02150 | 0.00360 | 0.40850 |
| **Yb** | 0.03000 | 0.00220 | 0.43000 |
| **Lu** | 0.03900 | 0.00080 | 0.42300 |
| **Hf** | 0.00110 | 0.00090 | 0.25600 |
| **Ta** | 0.00005 | 0.00080 | 0.01900 |
| **Pb** | 0.00010 | 0.09440 | 0.07200 |
| **Th** | 0.00005 | 0.00020 | 0.01200 |
| **U** | 0.00002 | 0.00010 | 0.01030 |

**Table A3.**

|  |  |  |  |  |  |  |  |  |  |  |
| --- | --- | --- | --- | --- | --- | --- | --- | --- | --- | --- |
| **Sample** | **3042** | **3063** | **3061** | **3031** | **3030** | **3077** | **2001\*** | **2004\*** | **2006\*** | **2029\*** |
| **Hamms temp trace** | 1430 | 1420 | 1440 | 1440 | 1450 | 1418 | 1435 | 1430 | 1430 | 1438 |
| **Hamms pressure trace** | 3.1 | 2.95 | 3.15 | 3.15 | 3.2 | 2.9 | 3.1 | 3.1 | 3 | 3.05 |
| **Hamms cont%** | 1 | 1 | 0.5 | 0.5 | 0.8 | 0.6 | 1 | 0.8 | 1.5 | 1.5 |
| **Hamms temp majors** | 1450 | 1474 | 1439 | 1449 | 1442 | 1430 | 1400 | 1427 | 1401 | 1434 |
| **Hamms pressure majors** | 2.4 | 2.6 | 2.3 | 2.4 | 2.4 | 2.2 | 2.3 | 2.4 | 2.3 | 2.5 |
| **Hamms f% majors** | 9 | 9 | 11 | 10 | 9 | 11 | 4 | 5 | 5 | 6 |
| **Hamms H2O%** | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 |
| **Hamms depletion %** | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 | 0 |
| **Lee temp** | 1452.308 | 1471.641 | 1439.281 | 1452.828 | 1446.588 | 1435.159 | 1445.802 | 1450.249 | 1452.414 | 1415.981 |
| **Lee Pressure** | 2.283879 | 2.600453 | 2.155021 | 2.286314 | 2.207761 | 2.127618 | 2.490639 | 2.52891 | 2.504108 | 2.122986 |
| **SiO2** | 45.86 | 45.10 | 46.35 | 46.01 | 46.17 | 46.38 | 44.98 | 43.98 | 44.34 | 44.89 |
| **TiO2** | 2.18 | 2.39 | 1.94 | 1.93 | 1.92 | 1.82 | 2.11 | 2.26 | 2.18 | 1.58 |
| **Al2O3** | 11.31 | 11.09 | 12.05 | 12.06 | 11.88 | 12.36 | 13.63 | 12.68 | 13.30 | 12.84 |
| **Fe2O3** | 10.67 | 11.02 | 10.50 | 10.74 | 10.61 | 10.35 | N/A | N/A | N/A | N/A |
| **FeO** | 1.54 | 1.59 | 1.51 | 1.55 | 1.53 | 1.49 | N/A | N/A | N/A | N/A |
| **feO tot** | 11.14 | 11.50 | 10.96 | 11.21 | 11.07 | 10.80 | 10.67 | 10.80 | 10.80 | 10.16 |
| **MnO** | 0.14 | 0.15 | 0.13 | 0.14 | 0.14 | 0.17 | 0.17 | 0.17 | 0.17 | 0.17 |
| **MgO** | 15.24 | 15.87 | 14.86 | 15.25 | 14.84 | 14.56 | 12.11 | 13.61 | 12.41 | 13.76 |
| **CaO** | 9.91 | 9.28 | 9.36 | 9.19 | 9.79 | 9.40 | 11.32 | 10.59 | 10.88 | 10.96 |
| **Na2O** | 2.32 | 2.22 | 2.31 | 2.21 | 2.20 | 2.30 | 2.39 | 2.24 | 2.14 | 2.10 |
| **K2O** | 0.33 | 0.63 | 0.51 | 0.49 | 0.48 | 0.57 | 0.61 | 0.44 | 0.57 | 0.59 |
| **P2O5** | 0.25 | 0.37 | 0.22 | 0.21 | 0.20 | 0.27 | 0.36 | 0.35 | 0.42 | 0.30 |
| **Cr2O3** | 0.03 | 0.03 | 0.02 | 0.01 | 0.02 | 0.11 | N/A | N/A | N/A | N/A |
| **Sc** | 29.27 | 27.77 | 29.48 | 28.20 | 30.61 | 31.09 | N/A | N/A | N/A | N/A |
| **V** | 324.75 | 315.56 | 315.48 | 317.90 | 305.03 | 264.48 | N/A | N/A | N/A | N/A |
| **Co** | 56.08 | 56.49 | 52.17 | 55.43 | 55.60 | 55.50 | N/A | N/A | N/A | N/A |
| **Ni** | 84.03 | 71.00 | 127.93 | 70.83 | 129.74 | 202.44 | N/A | N/A | N/A | N/A |
| **Rb** | 9.79 | 10.19 | 7.95 | 11.57 | 17.48 | 10.54 | 10.00 | 9.00 | 7.00 | 9.00 |
| **Sr** | 561.97 | 602.12 | 522.43 | 586.83 | 584.99 | 394.92 | 533.00 | 505.00 | 546.00 | 437.00 |
| **Y** | 22.96 | 22.65 | 25.39 | 23.09 | 30.12 | 22.98 | 23.20 | 20.90 | 25.00 | 20.10 |
| **Zr** | 135.87 | 141.45 | 175.14 | 147.32 | 290.70 | 127.86 | 99.00 | 93.00 | 94.00 | 89.00 |
| **Nb** | 21.58 | 22.51 | 28.23 | 23.58 | 39.08 | 19.66 | 32.90 | 30.20 | 30.60 | 23.40 |
| **Cs** | 0.13 | 0.11 | 0.60 | 0.13 | 0.21 | 0.06 | N/A | N/A | N/A | N/A |
| **Ba** | 212.11 | 218.16 | 269.58 | 227.44 | 387.25 | 187.42 | 247.00 | 193.00 | 287.00 | 290.00 |
| **La** | 18.09 | 18.50 | 24.00 | 19.42 | 34.04 | 15.12 | 22.20 | 19.50 | 23.20 | 16.70 |
| **Ce** | 38.57 | 39.91 | 51.67 | 42.07 | 70.53 | 33.85 | 41.80 | 38.30 | 42.20 | 35.70 |
| **Pr** | 4.99 | 5.14 | 6.39 | 5.54 | 9.05 | 4.47 | 5.46 | 4.98 | 5.81 | 4.35 |
| **Nd** | 21.84 | 22.41 | 27.05 | 23.45 | 37.42 | 19.56 | 22.80 | 20.90 | 24.60 | 18.40 |
| **Sm** | 5.03 | 5.13 | 6.04 | 5.27 | 8.17 | 4.62 | 4.86 | 4.45 | 5.17 | 4.01 |
| **Eu** | 1.77 | 1.75 | 1.96 | 1.80 | 2.62 | 1.57 | 1.61 | 1.49 | 1.68 | 1.37 |
| **Gd** | 5.10 | 5.08 | 5.92 | 5.23 | 7.75 | 4.90 | 4.62 | 4.38 | 4.96 | 3.84 |
| **Tb** | 0.74 | 0.74 | 0.85 | 0.77 | 1.10 | 0.74 | 0.71 | 0.64 | 0.76 | 0.62 |
| **Dy** | 4.40 | 4.44 | 4.96 | 4.55 | 6.28 | 4.50 | 4.10 | 3.77 | 4.34 | 3.60 |
| **Ho** | 0.83 | 0.84 | 0.92 | 0.86 | 1.15 | 0.84 | 0.81 | 0.73 | 0.86 | 0.68 |
| **Er** | 2.28 | 2.29 | 2.50 | 2.35 | 2.99 | 2.32 | 2.10 | 1.87 | 2.20 | 1.93 |
| **Tm** | 0.31 | 0.31 | 0.34 | 0.32 | 0.39 | 0.32 | N/A | N/A | N/A | N/A |
| **Yb** | 2.00 | 1.98 | 2.11 | 2.01 | 2.42 | 2.06 | 1.90 | 1.73 | 1.98 | 1.67 |
| **Lu** | 0.28 | 0.28 | 0.30 | 0.29 | 0.35 | 0.29 | 0.28 | 0.25 | 0.29 | 0.24 |
| **Hf** | 3.58 | 3.71 | 4.44 | 3.78 | 6.89 | 3.28 | 3.19 | 2.64 | 3.17 | 2.46 |
| **Ta** | 1.34 | 1.41 | 1.70 | 1.49 | 2.52 | 1.23 | 1.96 | 1.89 | 1.91 | 1.57 |
| **Pb** | 1.75 | 1.78 | 2.45 | 2.16 | 2.72 | 1.43 | 1.18 | 0.98 | 1.17 | 3.06 |
| **Th** | 1.56 | 1.63 | 2.20 | 1.69 | 3.16 | 1.29 | 2.41 | 1.98 | 2.14 | 1.93 |
| **U** | 0.43 | 0.45 | 0.53 | 0.48 | 0.84 | 0.37 | 0.53 | 0.47 | 0.48 | 0.44 |
|  |  |  |  |  |  |  | \*Akaki not back-corrected before Lee thermobarometer | | | |

**Mantle Potential Temperature and Pressure**

To estimate the mantle potential temperature and pressure of the Galema range melt, the model of (Lee et al., 2009) was utilized. While this model does back-correct to primary compositions for olivine fractionation, it does not account for clinopyroxene and plagioclase fractionation. For this reason, Galema range samples which had displayed evidence of olivine, plagioclase and clinopyroxene fractionation as identified in (Chiasera et al., 2018) were back-corrected using Petrolog v.3.1.1.3 (Danyushevsky and Plechov, 2011) prior to inputting into the Lee model. The Lee model identifies if an inputted sample has the same forsterite content as what is determined to be primary and if so, will not attempt a back-correction. Initial conditions for this primary forsterite content (Fo of mantle source) were set at 0.90, after Lee et al. (2009). The initial conditions of *f*O2 molar = 0.16, as determined by Excel MELTS modeling (Chiasera et al., 2018) and the Petrolog software corrections, was also set for all input chemistries. All other initial conditions were left as set by Lee et al. (2009). The results of the thermobarometry calculations for mantle potential temperature and pressure of the Lee model for the samples of the Galema range are listed in Table A3.

Akaki samples chosen using the above sample selection criteria were not corrected to primary compositions using the Petrolog software as it is assumed that only olivine fractionation is present. Instead, these samples were back-corrected using the native calculations within the Lee model. The results of the thermobarometry calculations for mantle potential temperature and pressure for Akaki samples are also listed in table A3. The results of these calculations indicate the pressure and temperature where the melt last achieved equilibrium with the asthenosphere (Lee et al., 2009).

**Mantle Melting Conditions**

Samples from the Galema range, back-corrected to primary compositions (Table A3.), were then inputted into HAMMS1 software (Kimura and Kawabata, 2014) to determine the lithology and composition of their melt material. Akaki samples (Table A1.) were also inputted into HAMMS1, but were not back-corrected to primary compositions. The initial composition of the source mantle utilized in the HAMMS1 models was set to the source mantle composition determined in (Rooney et al., 2014). This composition, estimated to be composed of 20% enriched mantle and 80% depleted mantle is shown in Table A4. It has been previously determined that this source mantle composition also has mixed within it lithospheric material in the amount of ~1% (Rooney et al., 2014). To account for this addition of lithospheric material, a composition that represents the trace element abundances of metasomatized lithospheric mantle was added in varying degrees to the 80%/20% source within the HAMMS1 model to achieve a match with the observed Galema sample chemistries. This composition, determined through experimental melting analysis of clinopyroxene-hornblendite (AG7-4) is shown in Table A4 (Pilet et al., 2008).

Within HAMMS1, each of the Galema sample compositions were compared to inputs reflecting varying degrees of composition AG7 (0-2% Fcont/%) addition, variable pressures of final melting (2.5-3.5 GPa Pmt/GPa), varying mantle potential temperature of final melting (1400°-1450° C Tp/C), variable water content (0-1% H2O(i)/%) and variable source depletion (0-1% CsDep/%ext). The initial conditions that best match the Galema sample compositions are shown in Table A3 and the resultant normalized trace element variation diagrams are shown in Figure 4 of the main text.

In addition to this, pressures and temperatures of melting determined from MgO melt thermometry (Herzberg and O’hara, 2002; Kimura and Ariskin, 2014) and degree of melt fraction (f%) of a peridotite source within HAMMS1 (CMAS) are also reported in Table A3. The variance between these two pressures and temperatures within HAMMS1 reflect the difference between the conditions of primary source melting and where the melt last achieved equilibrium with the mantle (Kimura and Kawabata, 2014).

**Table A4**

|  |  |  |
| --- | --- | --- |
|  | 80%/20% Source | AG7-4 Composition |
| Rb | 0.167 | 11.84 |
| Ba | 1.8482 | 422.00 |
| Th | 0.02332 | 0.60 |
| U | 0.00676 | 0.19 |
| Nb | 0.2614 | 26.57 |
| Ta | N/A | N/A |
| K | N/A | N/A |
| La | 0.291 | 16.80 |
| Ce | 0.795 | 44.70 |
| Pb | 0.0286 | 2.80 |
| Pr | 0.1408 | 6.50 |
| Sr | 10.3512 | 693.00 |
| Nd | 0.7356 | 31.00 |
| Sm | 0.28 | 7.30 |
| Zr | 6.3056 | 128.00 |
| Hf | 0.1874 | 4.20 |
| Eu | 0.1104 | 2.40 |
| Gd | 0.4056 | 6.50 |
| Tb | 0.0776 | 0.87 |
| Dy | 0.5514 | 5.00 |
| Y | 3.5724 | 21.30 |
| Ho | 0.1248 | 0.85 |
| Er | 0.3744 | 2.03 |
| Tm | N/A | N/A |
| Yb | 0.3906 | 1.41 |
| Lu | 0.0612 | 0.20 |

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