A FIELD AND PETROGRAPHIC STUDY OF THE SANDSTONES AND CONGLOMERATES OF THE PORCUPINE MOUNTAINS, ONTONTAGON COUNTY, MICHIGAN

by<br>HATEM HUSSEIN EL-KHALIDI<br>$=$

A THESSIS

Submitted to the School of Graduate Studies of Michigan State College of Agriculture and Applied Science
in partial fulfillment of the requirements
for the degree of
MASTER OF SCIENCE

Department of Geology and Geography

All rights reserved
INFORMATION TO ALL USERS
The quality of this reproduction is dependent upon the quality of the copy submitted.
In the unlikely event that the author did not send a complete manuscript and there are missing pages, these will be noted. Also, if material had to be removed, a note will indicate the deletion.


ProQuest 10008739
Published by ProQuest LLC (2016). Copyright of the Dissertation is held by the Author.
All rights reserved.
This work is protected against unauthorized copying under Title 17, United States Code Microform Edition © ProQuest LLC.

ProQuest LLC.
789 East Eisenhower Parkway
P.O. Box 1346

Ann Arbor, MI 48106-1346

## ACKNOWLEDGMENTS

The writer wishes to express his gratitude to Dr. W.A. Kelly for outlining the field procedure of this study, supplying aerial photographs, base and topographic maps, and giving valuable suggestions for the improvement of this paper.

The writer wishes to express his sincerest thanks to Dr. B.T. Sandefur who gave freely of his time and effort in supervising the laboratory procedure, and for his invaluable help in the writing of this paper.

To Drs. S.G. Bergquist and J. Trow, the writer wishes to convey his deep appreciation and thanks for their help in reading and editing this paper, and for their valuable suggestions on the style.

The field work of this paper would not have been possible without the active help of Iir. Joseph Elliot to whom the writer is very much indebted.

## TABLE OF CONTENTS

Page
INTRODUCTION ..... 1

1. Location ..... 1
2. General Topography ..... 1
3. General Geology. ..... 2
4. Purpose of Study ..... 14
PROCEDURE ..... 18
5. Field Procedure. ..... 18
6. Laboratory Procedure ..... 18
DESCRIPTION OF SECTIONS ..... 21
7. Section ..... 21
8. Section ..... 24
9. Section ..... III ..... 27
10. Section ..... IV ..... 33
11. Section ..... 36
12. Section ..... 41
13. General Statement ..... 43
IIINERALOGY AND PETROGRAPHY ..... 49
14. The Sandstones ..... 49
15. The Igneous Pebbles. ..... 57
16. General Statement ..... 60
CORRELATIONS ..... 63
CONCLUSION ..... 66
APPENDIX ..... 70
DESGRIPTIONS OF SAMPLES OF THE SECTIONS
17. Section ..... 70
18. Section ..... 74
19. Section ..... 78
20. Section IV ..... 87
21. Section ..... 90
22. Section ..... 104
BIBLIOGRAPHY ..... 109

IIST OF ILIUSTRATIONS
Figure Page
1 ..... 17 ..... 17
3 ..... 68
4 ..... 68
5 ..... 69
6 -••• ..... 69
IIST OF MAPS
Following page
1 ..... 1
IIST OF PLATES
Following page ..... 110

## INTRODUCTION

1. Location

The Porcupine Mountains area is located in Ontonagon and Gogebic Counties, between the Iron and the Presque Isle Rivers in the Upper Peninsula of Michigan (Map 1). The area is a State park, which is easily reached from Silver City, Michigan, on State Highway M-107.

## 2. General Topography

The Porcupine Mountains rise from the shore of Lake Superior in T. $51 \mathrm{~N} .$, Rs. 42,43 , and 44 W , as roughly concentric ridges parallel to the shore line.

The first ridge, rising from the shore of Lake Superior, reaches a height of 850 to 900 feet above the lake level, within a mile and a half to the south (Plate l). It then descends abruptly 400 feet into the valley of the Carp River. This river follows a very irregular course from its headwaters to the southeast, empties into the Lake of the Clouds, and then proceeds in a southwesterly direction and empties into Lake Superior at a location in the northeast corner of Gogebic County.

Beyond the Carp River to the south, there is a steep escarpment leading up to an eroded plateau and the valley of the Little Carp River. Knobs and ridges occur, but these are not as extensive or as steep as the first ridge. In this area tre highest elevations of the Porcupine Mountains occur. The highest point, as listed by the Michigan Department of Conservation, is Government Peak. It is 2,023 feet above sea

level, or l,421 feet above the level of Lake Superior; but as far back as 1908, A.C. Iane and others have indicated that this figure is too high. Lane believed it to be closer to 1,860 feet above sea level. More recent and accurate measurements based on field work and study of aerial photographs by Dr. W.A. Kelly of the Department of Geology, Michigan State College, show that another peak to the southeast of Government Peak is the highest. Dr. Kelly assigns an elevation of 1,900 to 1,950 feet above sea level to this peak. The Little Carp River empties into lirror Lake in the heart of the Porcupine Mountains, then into Lily Pond to the southwest, and finally into Lake Superior at a location approximately one mile southwest of the mouth of the Carp River.

Rainfall in the Porcupine ilountains is relatively heavy, and the elimination of the meteoric water by drainage and evaporation is small; thus, large swamp areas are present. Some areas around the valley of the Carp River are swamps. Others are located about a mile southeast of Government Peak.

## 3. General Geology

The rocks that outcrop in the Porcupine Mountains belong to the series known as the Keweenawan. A description of the general character of the Keweenawan rocks is essential to the understanding of the rocks of the Porcupine Mountains, because of the very close relationship between them and the other Keweenawan rocks that outcrop elsewhere, especially in the Keweenaw Peninsula and along the Black River to the southwest of the Porcupine Mountains.

## Distribution of the Keweenawan Rocks: The area in

 which these rocks are exposed lies in the southern segment of the pre-Cambrian shield of North America and forms a part of the Lake Superior basin. Along the south shore of the lake, the area extends from Keweenaw Point in Michigan southwestward through northern Wisconsin and into Minnesota. The Keweenawan rocks border the north shore of Lake Superior in Minnesota to the Canadian boundary, crop out in Isle Royale, and in Canada appear in the Black Bay and Thunder Bay districts, and extend to the area around Lake Nippigon.Character of the Rocks: The Keweenawan series are comprised of coarse clastic sediments together with intrusive and extrusive igneous rocks. Both the base and the upper part of the series are composed of sediments. The basal sediments, with a maximum thickness of 1,500 feet, are overlain by great thicknesses of basaltic lava flows interspersed with relatively thin sediments and acidic lava flows. The upper part of the Keweenawan series is composed wholly of thick conglomerate overlain by thin shale which, in turn, is overlain by thick sandstone.

Age of the Rocks: In the Gogebic Iron Range of Michigan, the Keweenawan is underlain by Penokee-Gogebic iron series which are upper-Huronian in age. In the Black River section there is a space of nearly half a mile between the lowest exposures of the Keweenawan, and the highest exposures of the Penokee-Gogebic; thus, no direct evidence of an
unconformity exists. Irving* suggested that a disconformity

```
* Irving, R.D., "The Copper Bearing Rocks of Lake Superior", United States Geological Survey, Honograph 5, p. 24 (1883).
```

occurs between the Keweenawan and Huronian. A complete discussion of a possible conformity or unconformity between the Keweenawan and the Huronian is given by Irving and Van Hise in the United States Geological Survey, Monograph 19, on the Penokee-Gogebic Range.

Overlying the upper Keweenawan without apparent unconformity or abrupt change in character, is the Lake Superior sandstone which is considered by the United States Geological Survey to be upper-Cambrian in age.

The United States Geological Survey groups the Keweenawan as upper-Proterozoic, of pre-Cambrian age. Some geologists disagree with this classification. A.C. Lane* argues that

[^0]since the apparently major unconformity lies below the Keweenawan series, rather than above it, the major division between the Cambrian and the pre-Cambrian should be placed at the base of the Keweenawan series. Other geologists believe that at least the upper Keweenawan should be classified as Cambrian.

The main arguments for the Cambrian age of the upper Keweenawen sediments are:

1. Both are red sandstones of the same lithologic character.
2. No basal conglomerate to the Cambrian against the upper Keweenawan has been discovered; whereas, there is a very thick one at the base of the upper Keweenawan containing a wide variety of pebbles of the intrusives as well as extrusives of the lower Keweenawan.

Some of the arguments against the idea of the Cambrian age of the upper Keweenawan are:

1. The unconformable overlap of the Potsdam sandstone of upper-Cambrian age on the upper Keweenawan.
2. The lack of fossils.
3. The presence of pebbles of Keweenawan in tine Cambrian along the Keweenawan fault shore cliff.

## General Stratigraphy:

In the writer's opinion, Irving's* division of the Kewee-

[^1]nawan series into upper and lover is more logical than the divisions suggested by various other geologists, because it is based on the fact that the upper Keweenawan is devoid of any evidences of igneous activity and is composed wholly of clastic sediments. The lower Keweenawan, on the otier hand,
is composed mostly of a succession of basic lava flows interstratified with sediments.

Below is a generalized stratigraphic column of the Keweenawan rocks (after A.C. Lane and A.E. Seaman)*:


Sedimentary Rocks: The sedimentary rocks are subordinate in the lower Keweenawan series; but they constitute the entire thickness of the upper series where they are composed of conglomerate, sandstone, and shale.

The sediments are predominantly red in color, feldspathic, and poorly sorted. They exhibit ripple marks, crossbedding, and mud cracks; and so far as known, are devoid of fossils. According to Lane*, their characteristics indicate

[^2]that they were deposited largely under terrestrial conditions.
The conglomerates are composed mainly of felsite clasts, and range in thickness from a few inches to 3,500 feet. Some zones have been traced by geologists many miles along the strike. The conglomerates form a very small proportion of the lower Keweenawan and a large proportion of the upper Keweenawan.

In general, the felsite conglomerates are of silicious composition, the most abundant rocks in them being felsite and auartz porphyry, although basic pebbles occur in minor quantities. Nost of the pebbles are well rounded. The conglomerate beds are generally lenticular, thinning or thickening rather rapidly along the strike. The coarser material, with boulders rarely exceeding a foot in diameter, is usually found in rather thick beas.

At the top of the Keweenawan series occur the Nonesuch
shale and the Freda sandstone formations. The Nonesuch formation consists of black and red shales interstratified with layers of red sandstone. The Freda sandstone which represents the uppermost member of the Keweenawan series is mainly a red sandstone with some felsite and basic igneous rock debris.

Source of the Sediments: The composition of material of the felsite conglomerate and the sandstones is similar to that of several masses of felsite and acidic porphyries that outcrop in the area. These and other old masses that have been eroded probably acted as a source of the sediments.
A.C. Lane* and others have suggested that certain exten-

Lane, A.C., "The Keweenawan Series of Michigan," Michigan Geological Survey, Publication 6 (Geol. Series 4), 2 vol., $983 \mathrm{pp} .(1911)$.
sive acid flows imnediately succumbed to erosion, forming a nearly contemporaneous conglomerate and sandstone. Butler and Burbank* suggested that the conglomerate pebbles were not

[^3]of immediate local derivations, but came from long-lived uplands composed mainly of felsite and quartz porphyry. They infer that the debris from these uplands, together with a small proportion of other material from different sources, was spread on the adjacent plains occupied by the basic lava
flows. The Porcupine Mountains area is an example of such an upland.

The Igneous Rocks: The igneous rocks are divided primarily into two classes; namely, the basic flows and the acid porphyries.

The basic flows make up the greater part of the Keweenawan series. These are fine grained, basaltic, crystalline rocks. They consist of a lower compact portion that grades upwards into vesicular and amygdaloidal parts. The flows are interstratified with layers of unaltered red sandstone and conglomerate. The basic flows are true eruptives and constitute successive flows from fissures. This suggestion was introduced by Irving* and later accepted by various

[^4]other geologists.
The acid rocks like the basic, are true eruptives, mainly porphyritic, and occur between layers of basic rocks.

Structure: The Keweenawan series in Michigan occur on the southern rim of the Lake Superior syncline or basin. Butler and Burbank* suggest that this basin was probably formed

[^5]during Keweenawan time. The lower Keweenawan rocks dip steeply, and the upper ones dip progressively less steeply northwestwardly towards the center of the Lake Superior basin. The Keweenawan rocks are faulted. The greatest fault in the region is the Keweenaw fault which bounds the Keweenawan series on the south from Keweenaw Point to Lake Gogebic. It is a thrust fault with a northwesterly dip, along which the basaltic series have been shoved over tine Lake Superior sandstone of upper Cambrian age. There is very little discordance between the dip of the fault and the dip of the beds of the flows. Many branch faults and fissures are associated with the Keweenaw fault. Butler and Burbank*

[^6]suggested that the movement along the Keweenaw fault probably did not begin until late Keweenawan time, much of it occurring after the Lake Superior sandstone was deposited.

Ore Deposits in the Keweenawan Rocks: In Michigan, the area of exposed Keweenawan rocks was one of the greatest copper districts of the world and outstanding in the deposits of native copper. This area forms a belt in the Keweenaw Peninsula two to four miles wide and 100 miles long, 26 miles of which has been highly productive. This district is now almost exhausted.

The copper lodes are of three types: (I) amygdaloidal,
which occur in the upper permeable parts of flows. Native copper occupies the vesicles. The lodes average 13 feet in thickness and contain from 0.6 to 1.5 per cent copper, (2) the conglomeratic lodes, in which native copper fills the interstices, (3) fissure veins, in which the native copper occurs in fissures and joints.

Butler, Burbank and other geologists consider that the copper deposits were formed by hydrothermal solutions given off from underlying basic intrusives. They believe that the structure is pre-ore in origin, and that the relations of copper beneath impervious cappings indicate rising rather than descending solutions. The solutions rich in copper, sulphur, and arsenic were guided upward by the permeability of the conglomerates and amygdaloidal tops. The hematite of the lava tops yielded oxygen which combined with the sulphur and arsenic, causing reduction of copper to native metal.

General Structure of the Porcupine Piountains: In the Porcupine Mountains the succession of rocks and their structure do not conform exactly with the structure of the Keweenawan rocks to the northeast and the southwest. In their geologic map of the Porcupine Mountains and vicinity, F.E. Wright and A.C. Lane* indicated the existence of a large

[^7]fold with a general northeasterly trend. The general structure is, moreover, complicated by the existence of numerous transverse and cross faults. The most prominent fault is the Main Porcupine Fault that borders the southern edge of the main felsite (Red Rock) area which forms the central portion of the Porcupine lountains. Here the Red Rock is in direct contact with the Nonesuch shale. The Main Porcupine Fault is a transverse fault with a general northeasterly trend. According to Wright and Lane*, it extends from sec. I,

[^8]「. 50 N., R. 42 W., westwards to sec. 7, T. 50 N., R. 42 W., and thence southwestwards to sec. 18, T. 49 N., R. 43 W. Any relation between this fault and the Keweenaw Fault has yet to be established.

The detailed geology of the Porcupine lountains is still in the most part unknown. More extensive field work is needed to throw more light on the problem. Loreover, attempts at the correlation of the numerous core drilling data which exist will lead to a much clearer picture of the Porcupine Mountains problem.

Stratigraphy of the Porcupine Mountains: Following is a generalized stratigraphic column of the Porcupine Mountains area. The information given here was taken mostly from Irving*, with some modifications by the writer.

[^9]| Name of formation | Description of formation <br> Nonesuch Shale | Thickness <br> in feet |
| :--- | :--- | :--- |
| Upper Keweenawan  <br> Porcupine Moun- <br> tains Sandstone shale and dark <br> sandstone Sandstone with thin <br> conglomerate zones | 1,500 |  |

## Lower Keweenawan

| Basaltic flows, | 300 to |
| :--- | :--- | :--- |
| amygdaloidal at top | 500 |

Sandstone and
1,900 conglomerate

Diabase, melaphyre, 400 amygdaloids, diabase porphyry, one bed of porphyry-conglomerate

Red Rock(?)
Felsites and quartziferous porphyries, banded rhyolite

All the above formations, with the exception of the Red Rock, show a general northeasterly trend and northwesterly dip, except in an area along the Iron River about one mile south of the Lake Superior shore. Here the outcropping

Nonesuch shale exhibits marked anomalies in strike and dip indicating the probability of the existence of the "nose" of the main fold in that area. On the other hand, these anomalies in strike and dip might be as a result of the faulting. The Red Rock, according to Irving*, shows no apparent

[^10]trends; but Thaden*, who spent the summer season of 1949

Thaden, R.E., Personal communication
studying the Red Rock in the field, disagrees with Irving's conclusions. Thaden indicated that the Red Rock exhibits definite trends. He also observed that the exposed contacts of the Red Rock with the basaltic lava flows are faulted. The reader, however, has to be content with this statement unsupplemented by evidence until the full text of Thaden's paper on the Red Rock is completed.
4. Purpose of Study

The purpose of this paper is to report the structure and petrology of exposed sandstones and conglomerates in the northern part of the Porcupine liountains along assigned traverses, and to investigate the possibilities of correlations of these rocks between the traverses.

The traverses along which the writer studied and sampled these rocks were confined to that sedimentary formation
stratigraphically above the Lake Shore iraps formation (Map 2). On Plate 14 accompanying the United States Geological Survey Professional Paper 144, the name assigned to the formation is the Outer conglomerate. This name was first given to the formation by Irving*. He suggested that

[^11]since it lies directly above the Lake Shore Traps, it was correlated with the Outer Copper Harbor conglomerate that outcrops in the Keweenaw Peninsula, and which also lies directly above the Lake Shore Traps. The Outer conglomerate in the Keweenaw Peninsula is composed essentially of coarse conglomerate with thin lenses of sandstone, and the name seems descriptively appropriate. Gordon* described the Outer conglom-

Gordon, Vi.C., and Lane, A.C., "A Geological Section From Bessemer Down Black River," Michigan Geological Survey Rept. for 1906, pp. 597-507, map (1907).
erate along the Black River section southwest of the Porcupine Mountains. Here the lower two-thirds of the formation is coarse conglomerate, and the upper one-third is sandstone. In the northern part of the Porcupine riountains, the writer found the majority of the outcrops to be sandstones with thin zones of conglomerate. It may be argued that the exposed rocks along the traverses comprise only a small percentage as compared to the unexposed rocks; thus, the unexposed rocks
might be conglomerates rather than sandstones. Fron his observations in the field, the writer noted that the conglomerate zones are more resistant to weathering than are the sandstone zones, and they occur in the form of knobs and prominent ridges (Figure I) as contrasted with the sandstone outcrops that lie adjacent to the surface of the ground (Figure 2). Thus, if the rocks in the area were mostly conglomerate, more knobs and ridges would have observed. It is, therefore, inferred by the writer that along the traverses the rock is mostly sandstone with subordinate zones of conglomerate.

Dr. W.A. Kelly* of the Department of Geology, Michigan

[^12]State College, pointed out that since the writer found more sandstone than conglomerate, the name Outer conglomerate seems very inappropriate and is misleading to the reader. IIoreover, even though the formation lies directly above the Lake Shore Traps, it might not be the same formation as the Outer Copper Harbor conglomerate. Also, the name Outer conglomerate is inappropriate since it signifies no geographical locality. Dr. Kelly suggested the name Porcupine Dountains Sandstone for that sandstone formation in the Porcupine llountains that lies stratigraphically above and in contact with the Lake Shore Traps. Thus the name: Porcupine Mountains Sandstone is introduced in this paper.


Fiqure l. A typical conglomerate knob.
liote the sandstone lens contained.

## 1. Field Procedure

The field work of this study was completed in the summer of 1949. Six traverse lines or sections, perpendicular to the general strike of the Porcupine Mountains Sandstone, were assigned and traversed (Map 2). Traversing was done mainly by pace and compass. Along Section $I$ where the brush was scanty, distances were measured with a steel tape. In other locations where the brush was very thick and accurate pacing or taping was impossible, the position of the outcrops was located by the use of aerial photographs.

Samples were collected from outcrops along the traverses. Where exposed rocks were continuous, samples were collected from the different lithologic zones. Strike and dip of bedding as well as prominent joints and other primary structures were observed, measured, and recorded.

## 2. Laboratory Procedure

In the laboratory, the samples were examined with the aid of a binocular microscope. Their descriptions are tabulated in the appendix accompanying this paper. Pebbles collected from the different conglomerate zones were classified and tabulated.

Exact mineral and petrographic determinations of the sandstones were made by examining selected thin sections. Also thin sections of type samples of the pebbles were examined and classified.

Sandstone grain size classifications in this paper follow the standards of the United States Bureau of Soils:

| Very fine grained | . $05-.10 \mathrm{~mm}$ |
| :---: | :---: |
| 2. Fine grained | . $10-.25 \mathrm{~mm}$ |
| 3. Miedium grained | .25-. 50 mm |
| 4. Coarse grained | . $50-1.0 \mathrm{~mm}$ |
| 5. Gravel | 1.0 mm and up |

Sphericity and roundness measurements were made on the pebbles and sand grains by the visual quantitative methods described by G. Rittenhouse* and W.C. Krumbein**.

[^13]> **Krumbein, W.C., Measurement and Geological Significance of Shape and Roundness of Sedimentary Particles," Journal of Sedimentary Petrology, Vol. 11, pp. 68-69, plate I, abstract (1941).

To determine the sphericity and roundness of the sand grains, a portion of each sandstone sample was crushed, and the loose grains examined under a binocular microscope. Ten grains were picked at random, and their sphericity and roundness were determined by comparing their two-dimensional images seen under the microscope with the images of standard grains and pebbles reproduced in the two charts of Rittenhouse and Krumbein. The calculated average sphericity and roundness of the ten grains for each sample, together with the sphericity and roundness of the pebbles, are tabulated in the appendix.

An attempt was made to correlate the exposed rocks along the traverses by means of:

1. The distinct Iithologic characteristics of the outcrops.
2. The sphericity and roundness of the sand grains.
DESCRIPTION OF SECTIONS

Section I
The location of Section $I$ is shown on Map 2 . This section follows the course of
Cuyahoga Creek from Location 1 on Highway M-lo7 northward to Location 2 on the shore
of Lake Superior. Taping was started at the north edge of Highway M-lor at Location
$I$ following the course of Cuyahoga Creek and continued northward to Iake Superior.
Outcrops were confined to the bed of the creek.
Following is a tabulated record of the traverse along section I:

| 1 | Location 1 | - | - | - | Amygdaloidal lavas (Sample 11) | (Lake Shore Traps) |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 1-2 | 102 | N. $0^{\circ} \mathrm{W}$ | $15^{\circ}$ | - | Amygajoloidal lavas | (Lake Shore Traps) |
| $2-3$ | 98 | N. $28^{\circ} \mathrm{E}$. | $10^{\circ}$ | - | Amygdaloidal lavas | (Iake Shore Traps) |
| $3-4$ | 100 | N. $3^{\circ} \mathrm{W}$ | $5^{0}$ | - | Amygdaloidal lavas (Sample 13) | (Lake Shore Traps) |
|  |  |  |  |  | Concealed interval |  |



| Slope $\begin{array}{l}\text { Inickness } \\ \text { in feet }\end{array}$ |
| :--- | :--- |

0
$\infty$

+ $0 \quad 0$ Conglomeratic sandstone
Bedding： $\begin{aligned} & \text { strike } \mathrm{N} \cdot 87^{\circ} \mathrm{E} . \\ & \\ & \operatorname{dip} 26^{\circ}, \mathrm{N}, 3^{\circ} \mathrm{W} .\end{aligned}$


## Ten土equt peteəouop

Concealed interval． Sandstone and conglomeratic sandstone， mud cracks present （Samples l to 10），N． 40 W．

$$
\begin{aligned}
& \circ \\
& \hline \\
& \not+ \\
& \text { H }
\end{aligned}
$$

$$
31.0
$$

Remarks Conglomerate
Bedding：strike N． $86^{\circ} \mathrm{W}$.


$$
\underset{\mathrm{r}}{0}
$$

$$
\begin{array}{c|c}
0 & 0 \\
0 & 0 \\
0 & i \\
-1 & r
\end{array}
$$

$$
\begin{aligned}
& 0 \\
& 0 \\
& -1
\end{aligned}
$$

$$
1
$$

Bearing
N． $20^{\circ} \mathrm{E}$.

$$
\begin{aligned}
& \dot{4} \\
& 0 \\
& \infty \\
& \dot{Z}
\end{aligned}
$$

Position
0
0
0
0

$$
\begin{aligned}
& \text { Distance } \\
& \text { in feet } \\
& 17
\end{aligned}
$$

Distance
in feet
17
POSition
0
-1
1
0
10－11
11－12

> 太
> $\begin{aligned} & \text { Gig } \\ & 0 \\ & \cdot-1 \\ & + \\ & 0 \\ & 0 \\ & 0 \\ & -1\end{aligned}$
Section II

| M-107 to Location 4 on the sh conducted along this section. |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
|  |  |  |  |  |  |
| Following is a tabulated record of the traverse along section II: |  |  |  |  |  |
| Position | Distance in feet | Bearing | Slope | Thickness in feet | Remarks |
| 1 | $\begin{array}{r} \text { Location } \\ 3 \end{array}$ | - | - | 9.0 | Sandstone cut by calcite veinlets Bedding: strike N. $75^{\circ} \mathrm{E}$. (Sample 1) $\operatorname{dip} 27^{\circ}$, N. $15^{\circ} \mathrm{W}$. <br> Joints: N. 80 W., vertical <br> S. $45^{\circ}$ W., vertical |
|  |  |  |  | 640 | Concealed interval |
| 1-2 | 2,541 | $\mathrm{N} \cdot 0^{\circ} \mathrm{W}$. | $10^{\circ}$ | 2.5 | Conglomeratic sandstone <br> Bedding: strike $1 . \bar{J} \cdot 73^{\circ} \mathrm{E}$. <br> $\operatorname{dip} 26^{\circ}, N .17^{\circ} W$. <br> Joints: N. $50^{\circ}$ E., vertical (Sample 2) <br> S. $20^{\circ} \mathrm{E} .$, dip $80^{\circ} \mathrm{W}$. <br> S. $30^{\circ}$ W., vertical |
|  |  |  |  | 70 | Concealed interval |
| 2-3 | 180 | $\mathrm{N} \cdot \mathrm{O}^{\circ} \mathrm{W}$. | $0^{\circ}$ | 3.0 | Sandstone cut by calcite veins (Sample 3) |

Position

100

Remarks Concealed interval
Sandstone, irregular leaching
Joints: N. 60 E., vertical
(Samples 4, 5, and 6)
Concealed interval Conglomeratic sandstone
Conglomerate
Sandstone
Joints: S. $20^{\circ}$ E., vertical
(Samples 7,8 , and 9 )
Concealed interval Sandstone
Joints: $N .70^{\circ} \mathrm{W} .$, vertical
(Sample lo) $16^{\circ} \mathrm{E}$. , vertical
Concealed interval Concealed interval
Sandstone
(Sample 11)
Concealed interval
Conglomeratic sandstone
Bedding: strike N. $71^{\circ} \mathrm{E}$.
(Sample 12 ) $270^{\circ}, \mathrm{N} \cdot 19^{\circ} \mathrm{W}$.




$\stackrel{\circ}{\circ}$
$\bigcirc$
$\stackrel{\circ}{\circ}$
in

|  |
| :---: |

N. $0^{\circ} \mathrm{W}$.

$$
\text { N. } 0^{\circ} \mathrm{W}
$$

$$
\begin{aligned}
& \dot{3} \\
& \circ \\
& 0 \\
& \vdots
\end{aligned}
$$

$$
\begin{array}{cc}
\dot{3} & \dot{z} \\
0 & \circ \\
0 & \dot{Z} \\
\dot{z} & \dot{z}
\end{array}
$$

$$
\left.\begin{gathered}
+ \\
0 \\
0 \\
0 \\
4-1 \\
9 \\
.-1
\end{gathered} \right\rvert\,
$$

$$
\begin{aligned}
& \text { Distance } \\
& \text { in feet }
\end{aligned}
$$

$\infty$
$\stackrel{N}{N}$
$\sim$
O
8
$\stackrel{8}{1}$
-1
0
0
0
0
0


| $H$ |
| :---: |
| $\infty$ |

10
$d$
8
0
1
1
$\hat{1}$
$\vdots$
$7-8$

Remarks

Thickness
in feet

Slope

Bearing

Distance
in feet
Position
Position
Concealed interval
Conglomeratic sandstone

$$
\begin{array}{ll}
\text { Bedding: } & \text { strike } N .70^{\circ} \\
& \operatorname{dip} 25^{\circ}, N . \\
0^{\circ}
\end{array}
$$

-W.
TBOTqIIOA ' "TH

$$
.0_{8}^{08}
$$

$$
\begin{aligned}
& \text { Sandstone showing mud cracks and } \\
& \text { ripple marks }
\end{aligned}
$$

$G \cdot \square$
$G T$
$\begin{array}{llll}0 & 0 & 0 \\ -1 & 0 & N & 0\end{array}$
O
O
$\stackrel{1}{1}$
0
10
-1
앙

Sandstone showing mud cracks and
ripple marks
Bedding: strike N. $70^{\circ} \mathrm{F}$.
Joints: S. $50^{\circ} \mathrm{E}, \mathrm{N}$, vertical
(Samples 17 to 25 )
Calcite vein six inches thick,
trends S. $16^{\circ} \mathrm{E} .$, vertical
trends S. $16^{\circ}$ E., vertical
008
dip $28^{\circ}, \mathrm{N}$.
(Sample I3)
Concealed interval
Conglomeratic sandstone
(Sample 14)
Concealed interval
Conglomerate
Joints: S. $250 \mathrm{E} .$, vertical
(Sample 15) Concealed interval
Sandstone, massive, cut by calcite veins
(Sample 16)
Concealed interval
Concealed interval
Section III
Section III extends from Location 5 on Highway M-107 to Location 8 on the shore
of Lake Superior as shown on Map 2 .
Following is a tabulated record of the pace and compass traverse along Section III:
Following is a tabulated record of the pace and compass traverse along Section III:

$$
\begin{aligned}
& \text { Slope Thickness Remarks } \\
& \text { in feet }
\end{aligned}
$$


730
0.5
2.6
O
15


20
2.7
70

15
$\stackrel{+}{+}$

| 0 | 1 | 0 |
| :--- | :--- | :--- |

$\stackrel{0}{\circ}$
$\stackrel{-1}{\circ}$
$\stackrel{1}{2}$
오
$\stackrel{\circ}{\circ}$
$10^{\circ}$
O

| $\dot{B}$ | $\dot{B}$ |
| :--- | :--- |
| $\dot{\circ}$ | $\dot{0}$ |
| $\dot{\sim}$ | $\dot{z}$ |

8
8
8
8
3

8
Position
$\frac{\text { Bearing }}{\text { N. } 0^{\circ} \mathrm{W} .}$

is


$$
\begin{aligned}
& \dot{z} \\
& \stackrel{\circ}{\circ} \\
& \dot{z}
\end{aligned}
$$



$$
\stackrel{7}{6}
$$



$$
\mathscr{\infty}
$$

$$
\stackrel{o}{\stackrel{\rightharpoonup}{-1}} \underset{\sim}{\circ}
$$


1
0
$i$
$i$
$\hat{1}$
$i$
$\begin{array}{ll}\infty & 0 \\ 1 & 1 \\ & \infty\end{array}$

管



$$
\underset{\sim}{\infty}
$$

$-H$
-1
-1
-1
11-12
12-13
$H$
$H$
1
1
-1

| Position | Distance in feet | Bearing | Slope | Thickness <br> in feet | Remarks |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 14-15 | 130 | N. $70^{\circ} \mathrm{E}$. | $0^{\circ}$ | 6.0 | Conglomeratic sandstone cut by quartz veins, trending $N .49^{\circ}$ W., up to four inches thick <br> (Samples 20, 21, 22, and 23) |
|  |  |  |  | 10 | Concealed interval |
| 15-16 | 69 | $\mathrm{N} \cdot \mathrm{O}^{\circ} \mathrm{W}$ | $15^{\circ}$ | 1.6 | Sandstone (Sample 24) |
|  |  |  |  | 10 | Concealed interval |
| 16-17 | 150 | $\mathrm{N} \cdot \mathrm{O}^{\circ} \mathrm{W}$ | $20^{\circ}$ | 4.0 | Conglomeratic sandstone cut by quartz veins, trending $N .45^{\circ} \mathrm{W}$. (Sample 25) |
|  |  |  |  | 13 | Concealed interval |
| 17-18 | 90 | $\mathrm{N} .0^{\circ} \mathrm{W}$. | $15^{\circ}$ | 4.5 | Sandstone cut by quartz veins, trending N. $5^{\circ}$ W., N. $15^{\circ} \mathrm{E} .$, and N. $60^{\circ} \mathrm{E}$. (Sample 26) |
|  |  |  |  | 10 | Concealed interval |
| 18-19 | 168 | N. $0^{\circ} \mathrm{W}$ | $20^{\circ}$ | 2.0 | Sandstone <br> Bedding: strike N. $69^{\circ} \mathrm{E}$. (Sample 27) dip $21^{\circ}, N .21^{\circ} \mathrm{W}$. |
|  |  |  |  | 20 | Concealed interval |



| Position | Distance in feet | Bearing | Slope | Thickness in feet | Remarks |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 24-25 | 123 | N. $0^{\circ} \mathrm{W}$. | $20^{\circ}$ | 2.0 | Conglomeratic sandstone, crossbedded (Sample 36) |
|  |  |  |  | 6 | Concealed interval |
| 25-26 | 90 | N. $0^{\circ} \mathrm{W}$. | $20^{\circ}$ | 3.2 | Sandstone <br> Bedding: strike $\mathbb{N} .68^{\circ} \mathrm{E}$. (Sample 37) dip $25^{\circ}, N .22^{\circ} \mathrm{W}$. |
|  |  |  |  | 4 | Concealed interval |
| 26-27 | 65 | N. $0^{\circ} \mathrm{W}$. | $20^{\circ}$ | 1.0 | Conglomerate (Sample 38) |
|  |  |  |  | 4 | Concealed interval |
| 27-28 | 60 | IN. $0^{\circ} \mathrm{W}$. | $20^{\circ}$ | 2.2 | Sandstone (Sample 39) |
|  |  |  |  | 6 | Concealed interval |
| 28-29 | 92 | N. $0^{\circ} \mathrm{W}$. | $20^{\circ}$ | 3.7 | Sandstone <br> Bedding: strike $\mathbb{N} .72^{\circ} \mathrm{E}$. (Sample 40) $\operatorname{dip} 24^{\circ}, \mathrm{N} .18^{\circ} \mathrm{W}$. |
|  |  |  |  | 180 | Concealed interval |
| 29-30 | Location 8 | N. $0^{\circ} \mathrm{W}$. | $0^{\circ}$ | - | Outcrops of conglomeratic sandstone 25 feet off shore |

Section IV

| Section IV extends from Location 9 on the shore of Lake Superior to Location 10, thence in a $S .25^{\circ} \mathrm{E}$. direction as shown on Map 2. <br> Following is a tabulated record of the pace and compass traverse along Section IV |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |
| Position | Distance in feet | Bearing | Slope | Thickness in feet | Remarks |
| 1 | Location 9 | - | - | $\begin{aligned} & 3.0 \\ & 1.0 \\ & 1.0 \end{aligned}$ | Sandstone, cross-bedded <br> Conglomeratic sandstone, ripple marks <br> Conglomerate <br> Joints: $\mathbb{N} .50^{\circ}$ W., vertical <br> N. $50^{\circ}$ W., dip $75^{\circ}, N .40^{\circ}$ E. <br> Bedding: strike $\mathrm{N}^{\prime} .66^{\circ} \mathrm{E}$. $\operatorname{dip} 25^{\circ}, \mathrm{N} .24^{\circ} \mathrm{W}$ <br> (Samples 1 to 5) |
|  |  |  |  | 750 | Concealed interval |
| 1-2 | $\begin{array}{r} \text { Location } \\ 10 \end{array}$ | S. $25^{\circ} \mathrm{E}$. | - | 2.2 | ```Sandstone Bedding: strike N. 660 E. (Sample 6) dip 24%,N. 240 W.``` |
|  |  |  |  | 4 | Concealed interval |
| 2-3 | 50 | S. $25^{\circ} \mathrm{E}$. | $20^{\circ}$ | 3.5 | Sandstone (Sample 7) |
|  |  |  |  | 5 | Concealed interval |

POSit£On
$3-4$
4-5
0
1
1
6-7
$7-8$
Concealed interval

앙


$\stackrel{\circ}{\circ}$

$\therefore \stackrel{\circ}{\circ}$
$\stackrel{\circ}{\infty}$
Slope


64
宇
웅
i
$\dot{2}$

ค
-
$\underset{-}{-}$
$\stackrel{\circ}{\sim}$
$\infty$
$\stackrel{\infty}{-}$
-
$\stackrel{\circ}{\circ}$

S. $25^{\circ} \mathbb{E}$.
$\stackrel{\circ}{\circ}$

$\circ$
$\circ$
0
E.

O
$\stackrel{1}{2}$
$i$
$i$
0
10
$H$
$\stackrel{0}{i}$

9
$\circ$
$\stackrel{1}{\infty}$
ふ

21


$\begin{array}{lll} & \infty & \\ 0 & 0 & \end{array}$

$\stackrel{0}{i}$

|  | 过 |
| :---: | :---: |
| * | $\stackrel{\circ}{\circ}$ |
| ¢ | - |


Position
$8-9$
$9-10$
Section V
Section V extends from Location 11 on the shore of Lake Superior to Location 16

$$
\begin{aligned}
& \text { as shown on Map } 2 . \\
& \text { Following is a }
\end{aligned}
$$

$$
\text { the pace and compass traverse along section } V \text { : }
$$ Sandstone

Conglomeratic sandstone
Joints: N. $45^{\circ}$ W., vertical
(Samples 2 and 3 )
Sandstone, parallel and cross-bedding
Bedding: strike N. $68^{\circ}{ }^{\text {E. }} \mathrm{W}$.

(Sample 1) | Thickness |
| :--- |
| in feet |


Slope

| $\stackrel{0}{4}$ |
| :---: |
| $\stackrel{4}{4}$ |
| $\stackrel{4}{0}$ |
| $\stackrel{0}{0}$ |
| $\stackrel{a}{4}$ |

Section $V$ extends from Location 11 on the shore of Lake Superior to Location 16
Following is a tabulated record of the pace and compass traverse along section $V$ :

$$
\begin{aligned}
& \text { Conglomeratic sandstone } \\
& \text { Sandstone, cross-bedded } \\
& \text { Conglomerate } \\
& \text { Joints: N. } 50^{\circ} \text { W., vertical } \\
& \text { N. } 10^{\circ} \text { W., vertical } \\
& \text { S. } 80^{\circ} \text { W., vertical } \\
& \text { (Samples 5, 6, and'7) } \\
& \text { Conglomeratic sandstone } \\
& \text { Joints: N. } 20^{\circ} \text { E., dip } 75^{\circ}, \text { N. } 70^{\circ} \text { W. } \\
& \text { (Sample 4) } 45^{\circ} \text { W., vertical }
\end{aligned}
$$

| Position | Distance <br> in feet | Bearing | Slope | Thickness in feet | Remarks |
| :---: | :---: | :---: | :---: | :---: | :---: |
| 1 | Location 11 | - | - | 1.6 | Conglomeratic sandstone |
|  |  |  |  | 1.2 | Sandstone, cross-bedded |
|  |  |  |  | 0.5 | Conglomerate |
|  |  |  |  |  | Joints: $\begin{aligned} & \text { N. } 50^{\circ} \text { W., vertical } \\ & \\ & \text { N. } 10^{\circ} \text { W., vertical }\end{aligned}$ |
|  |  |  |  |  | $\text { (Samples } 5,6 \text {, and } 7 \text { ) }$ |
| 1-2 | 100 | S. $65^{\circ} \mathrm{W}$. | - | 3.1 | Conglomeratic sandstone <br> N. $45^{\circ}$ W., vertical <br> Joints: N. $20^{\circ} \mathrm{E} ., \operatorname{dip} 75^{\circ}, \mathrm{N} .70^{\circ} \mathrm{W}$. |
|  |  |  |  |  | (Sample 4) |
| 2-3 | 50 | S. $65^{\circ} \mathrm{W}$. | - | 2.5 | Sandstone |
|  |  |  |  | 1.8 | Conglomeratic sandstone |
|  |  |  |  |  | Joints: N. $45^{\circ}$ W., vertical (Samples 2 and 3) |
| 3-4 | 50 | S. $65^{\circ} \mathrm{W}$. | - | 4.5 | Sandstone, parallel and cross-bedding |
|  |  |  |  |  | Bedding: strike N. $68^{\circ} \mathrm{F} . \mathrm{W}$. |
|  |  |  |  |  | (Sample 1) |

Position

15
$\stackrel{N}{2}$
$\begin{array}{llll}10 & 0 & 0 & 1 \\ \cdots & 0 & 0 & 0\end{array}$

Sandstone
Conglomeratic sandstone
Conglomerate
Conglomeratic sandstone
(Samples $10,11,12$, and 13) Concealea interval

Concealed interval
Conglomeratic sandstone
(Sample 16 )
Slope Thickness

| 0 |
| :--- |
| $0_{1}$ |
| 0 |
| -1 |
| 0 |





## 1

$\underset{-1}{O}$
1-5

0
1
1
5
1
6
$\begin{array}{cc}\infty & \infty \\ 1 & \infty \\ \boldsymbol{1} & \infty\end{array}$
$\begin{array}{cc}\infty & \infty \\ 1 & \infty \\ \boldsymbol{1} & \infty\end{array}$
$\dot{9}$
0
$\stackrel{0}{2}$
$\dot{0}$

$\stackrel{10}{2}$
$\stackrel{+}{\circ}$
$\stackrel{\circ}{\circ}$
$\stackrel{\circ}{-1}$

|  | 田 - 立 |
| :---: | :---: |
|  | $\begin{array}{ll} \circ & \stackrel{\circ}{\circ} \\ \underset{\sim}{0} & \end{array}$ |
|  | $\dot{j}$ j |

$\dot{9}$
0
0
0
0
$i 2$

8
○ $\quad \underset{\infty}{\text { - }}$

$4 \quad \stackrel{\Gamma}{-i}$
LO
4io
:

$\stackrel{\circ}{\circ}$

O
O
O



| 1 |
| :---: |
| 0 |
| 0 |
| +1 |
| +-1 |
| 0 |
| 0 |
| 0 |
| 0 |

$$
0
$$

$\circ$
$\infty$
$\infty$
44
$\underset{\infty}{\infty}$
0
0
0
0
0
0
0
R
0
0
0
0
0
S. $25^{\circ} \mathrm{E}$.
0
0
0
10
0
0
$\dot{0}$

Slope Thickness
in feet


| Distance |
| :--- |
| in feet |

$\stackrel{10}{\leftrightarrow}$
$\stackrel{1}{2}$

| $\circ$ |
| :--- |
| $\circ$ |
|  |

$\dot{1}$
$\stackrel{1}{1}$
$\stackrel{1}{2}$
$\dot{4}$
0
0
0
0
0
0
0


$\begin{array}{lll}0 & 0 & 0 \\ \Gamma & \infty \\ \Gamma & 0 \\ \Gamma & \infty\end{array}$
$\stackrel{\text { 옹 }}{\stackrel{\circ}{-1}}$
1

|  | 0 | 0 |
| :--- | :--- | :--- |
|  | 0 | - |
|  |  |  |

$\begin{array}{ll}0 \\ i \\ i & 0 \\ i\end{array}$

1 | g |
| :---: |
| 0 |
| +-7 |
| +3 |
| +-1 |
| 0 |
| 0 |
| 0 |
| 0 |

$$
13-14
$$


0
1
1
-1
16-17
17-18
Remarks
Concealed interval
Conglomerate, boulders up to 1.0 feet
(Samples $37,38,39$, and 40 )
Concealed interval
Thickness
in feet
140
11.0
-

| 0 |
| :---: |
| 0 |
| 0 |
| 0 |
| 0 |
| 0 |
| 0 |$|$

I

|  | $\square^{\circ}$ |
| :---: | :---: |
| 80 | $\bigcirc$ |
| $\stackrel{-7}{4}$ |  |
| $\stackrel{\widetilde{\otimes}}{\oplus}$ | $\pm 0$ |


|  | Clic |
| :---: | :---: |
| ${ }_{\square}$ |  |
| $\stackrel{\sim}{0}$ | $+$ |
|  | O |
| $\stackrel{\sim}{A} \cdot \underline{H}$ | $\bigcirc$ |


| $\begin{aligned} & \text { gin } \\ & \text {-r } \end{aligned}$ |  |
| :---: | :---: |
|  |  |
|  |  |
|  |  |
| O |  |

0
$\cdots$
1
0
$\cdots$
-1
Section VI



Slope


$\xrightarrow{\text { Position }}$
0

$\begin{array}{llll}0 & 0 & 1 & 0 \\ 0 & 0 & 0 & 0 \\ -1 & -1 & -1 & 0\end{array}$

## General Statement

Following is a general description of the rock outcrops along the sections (traverses):

1. Lithologic types
a. Conglomeratic sandstone: These are exposed along all the sections, and constitute the majority of the outcrops. They exhibit mainly parallel bedding, although some zones are massive. They vary in color from brick red to dark brown, and in grain size range from fine to coarse. Pebbles disseminated throughout the body of the sandstones do not compose more than ten per cent of the total volume of the rock. The sandstones grade into non-conglomeratic phases along both the strike and dip.
b. Non-conglomeratic sandstone: These are exposed along all the sections. In abundance they are next to the conglomeratic sandstone. The sandstones mainly exhibit parallel bedding, but some zones are massive. They vary in color from brick red to brown, and in grain size range from fine to coarse. The non-conglomeratic sandstones grade into conglomeratic phases along both the strike and dip.
c. Coarse conglomerate: These are exposed along all the sections, but constitute a minority of the outcrops. Their thickness varies from a few inches to 18 feet. The conglomerates are brown and are composed of well rounded, poorly sorted pebbles with an average diameter
of three inches, but some boulders of over one foot in diameter were observed. The conglomerate zones are more abundant in Sections $V$ and VI than in the other sections. This indicates that there is a change in lithology from a sandy to a coarse conglomerate phase westwards along the strike of the formation; but even in these two westerly sections, the conglomerate zones are subordinate to the sandstones.
d. Basic lava flows: These are exposed only in the southern part of Section $I$. They are dark greenish gray in color and are mostly amygdaloidal. The amygdules are filled with epidote. These rocks represent the upper section of the Lake Shore Traps formation. A concealed interval of 29 feet obscures their actual contact with the sandstone.
2. Structural features
A. Strike and dip: Strike and dip measurements were taken on zones of sandstones and conglomeratic sandstones exhibiting parallel bedding. Following is a tabulated record of the average strikes and dips along each of the six sections:

Section number Average strike Average dip
N. $88^{\circ}$ E.
N. $70^{\circ}$ 玉.

IN. $70^{\circ} \mathrm{E}$.
N. $65^{\circ} \mathrm{E}$.
N. $65^{\circ} \mathrm{E}$.
N. $56^{\circ} \mathrm{E}$.
$24^{\circ}, \mathbb{N} .2^{\circ} \mathrm{W}$. $26^{\circ}$, N. $20^{\circ} \mathrm{W}$. $23^{\circ}, N .20^{\circ} \mathrm{N}$. $24^{\circ}$, N. $25^{\circ} \mathrm{N}$. $24^{\circ}$, N. $25^{\circ}$ W. $27^{\circ}, N .34^{\circ} W$.

The regional strike of the outcrops in the area is approximately parallel to the shore of Lake Superior (Map 2); i.e., it swings northwards to the southwest. The dip increases to the southwest.

## B. Fractures:

1. Faults: Evidences of cross faulting occur along Sections III and VI in the form of brecciated zones (Ihap 2), but no actual displacement of beds was observed. In their geologic map of the Porcupine Mountains, F.E. Wright and A.C. Lane show many cross faults cutting the Lake Shore Traps formation (Map 2). These probably extend into the sandstone to the north.
2. Joints: The outcrops of conglomeratic sandstones, sandstones, and conglomerates are jointed in places (See descriptions of sections). The majority of the joints are vertical; a few possess steep dips ranging from 75 to 85 degrees. There are three main joint systems with trends respectively:
a. Oblique joints that trend in a northeastsouthwest direction.
b. Oblique joints that trend in a northwestsoutheast direction.
c. Dip joints that trend in an approxinate north-south direction.

It seems very logical that the joints are closely related to the faulting reported in this area.

If this is true, the joints would be classified as feather joints; i.e., tension fractures. This is only an inferred classification, and more extensive field study is necessary before any definite conclusion could be reached regarding the true genetic classification of these joints.
C. Primary features:
a. Cross-bedding: These features are not extensive but occur locally. They are not very well exposed; the tangential parts, if any, and the truncations could not be seen. Thus, it could not be determined whether the cross-bedding was normal or torrential.
b. Ripple marks: These were observed in but a few localities. Nost of them are very badly weathered and indistinct, but in places it was determined that they were mainly current ripple marks.
c. Mū cracks: "Horseshoe" type mud cracks were observed in few localities. They vary from a fraction of one inch in extent to about six inches for each "horseshoe".
3. Sand grain, sphericity, and roundness analysis The average sphericity and roundness of each sandstone sample was determined by the visuel methods previousiy described. From an examination of the results a few significant facts are noted:
a. Coarse grains show the highest roundness value with an average of .6 .
b. Liedium grains show an average roundness value of .5.
c. Fine grains show an average roundness value of .4 to . 5.
d. The sphericity figures are not significant since they vary haphazardly and not according to grain size. They range in value from . 69 to . 81.
4. Thickness of the sections The inferred thickness of the concealed and exposed sandstones along each of the sections is reported below. It was calculated by measuring the map distances between the highest and lowest stratigraphic exposed zones perpendicular to the average strike along each section. The average dips (p. 44) were used in the calculation.

Section number
I
II
III
IV
V
VI

Total thickness in feet

$$
1,600
$$

$$
1,700
$$

$$
1,150
$$

$$
1,050
$$

$$
1,650
$$

1,900

The above would be the true thicknesses only if no repetition of beds occur due to faulting. The thickness of the concealed intervals between outcrops as recorded in the description of the sections (pp. 21-44),
was computed by the use of Mertie's* chart. In each case,

* Mertie, J.B., "Graphic and IEchanical Computation of Thickness of Strata and Distance to a Stratum," United States Geological Survey, Prof. Paper 129-C, plate VI (1921).
the strike and dip used in the estimation of the thickness is the average of the strikes and dips obtained on the north side and the south side of the concealed section.

5. Secondary inclusions

The sandstone outcrops are cut in places by irregular quartz and calcite veinlets. The thickest calcite veinlet observed is six inches, and the thickest quartz veinlet is four inches. Their exact locations an $\begin{aligned} & \text { trends are }\end{aligned}$ recorded in the description of the sections (pp. 21-44).

## MINERALOGY AND PETROGRAPHY

1. The Sandstones

Pight thin sections of the sandstones were examined:
Section I
Sample 1. Sandstone (Orthoauartzite)
Megascopic description: The sandstone is brown, medium grained, and bedded. The grains are well sorted. Quartz and feldspars are visible to the eye.

Microscopic description: Clastic, well sorted, angular, and sub-angular grains.

| Stable <br> "primary" <br> minerals* | Stable <br> "secondary" <br> minerals** | Precipitates* <br> "cement" | Minor con- <br> Stituents |
| :--- | :--- | :--- | :--- |
| Oligoclase-10\% | Chlorite | Hernatite |  |

[^14]Sample 14. Arkose

Megascopic description: The arkose is brown. Poorly sorted, coarse quartz and feldspar grains are visible to the eye.

Microscopic description: Clastic, poorly sorted, subangular and rounded grains. The arkose contains about five per cent of igneous rock fragments. These are mainly composed of microlitic feldspars possessing a felty texture.

| Stable <br> "primary" <br> minerals* | Stable <br> "secondary" <br> minerals* | Precipitates* <br> "cement" | Minor con- <br> Stituents |
| :--- | :--- | :--- | :--- |
| Oligoclase- $30 \%$ | Kaolin-5\% |  |  |
| Orthoclase-5\% |  |  |  |

* Pettijohn, F.J., Sedimentary Rocks, p. 86 (1940).


## Section II

Sample l. Sandstone (Orthoquartzite)

Megascopic description: The sandstone is brown, fine grained, and is cut by calcite veinlets. Quartz, feldspars, and mica are visible to the eye.

Microscopic description: Clastic, well sorted, angular and sub-angular grains.

| Stable <br> "primary" <br> minerals* | Stable <br> "secondary" <br> minerals* | Precipitates* <br> "cement" | Minor con- <br> Stituents |
| :--- | :--- | :--- | :--- |
| Oligartz-85\% | Kaolin |  | Calcite |

* Pettijohn, F.J., Sedimentary Rocks, p. 86 (1940).

Sample 17. Arkosic sandstone

Megascopic description: The sandstone is brown, medium grained, and bedded. The grains are poorly sorted. Brick red, parallel, platy, discontinuous zones about 0.5 millimeters thick (ferruginous clay zones?) are found. Microscopic description: Clastic, poorly sorted, angular and sub-angular grains. The sanastone contains five per cent of igneous rock fragments. These are mainly composed of microlitic feldspars, and possess a felty texture.

Stable
"primary"
minerals*
Quartz-70\%
Oligoclase-15\%
Microcline-20\%
Perthite- $3 \%$
Muscovite
Tuscovite

Stable "secondary" Precipitates*

Minor con"cement" stituents minerals*

Chlorite
Calcite
Hematite
Magnetite$2 \%$
Kaolin

## Section III

Sample 28. Sandstone (Orthoquartzite)

Megascopic description: The sandstone is brown, fine grained, and brecciated.

Microscopic description: Clastic, well sorted, angular and sub-angular grains. The spaces between the brecciated fragments are filled with cryptocrystalline quartz.

| Stable <br> "primary" <br> minerals* | Stable <br> "secondary" <br> minerals* | Precipitates* <br> Quartz-90\% | Chlorite |  |
| :--- | :--- | :--- | :--- | :--- | | Minor con- |
| :--- |
| Stituents |

[^15]
## Section IV

Sample 3. Arkosic sandstone

Megascopic description: The sandstone is brown, medium grained, bedded, and cut by calcite and quartz veinlets. Microscopic description: Clastic, poorly sorted, angular and sub-angular grains. The sandstone contains about five per cent of igneous rock fragments. These are mainly composed of microlitic feldspars, and possess a felty texture.

| Stable "primary" minerals* | Stable <br> "secondary" minerals* | Precipitates* <br> "cement" | Minor constituents |
| :---: | :---: | :---: | :---: |
| Quartz-75\% | Kaolin | Calcite-2\% | Zircon |
| Oligoclase-15\% | Chlorite | Hematite | Epidote |
| Microcline-3\% | Sericite |  | Magnetite |
| Orthoclase-2\% |  |  |  |
| Muscovite |  |  |  |
| Biotite |  |  |  |

[^16]
## Section V

Sample 14. Sandstone (Orthoquartzite)

Megascopic description: The sandstone is brown and fine grained. It is bedded and micaceous.

Microscopic description: Clastic, well sorted, angular and sub-angular grains.

| Stable <br> "primary" <br> minerals* | Stable <br> "secondary" <br> minerals* | Precipitates* <br> "cement" | Minor con- <br> Qtituents |
| :--- | :--- | :--- | :--- |
| Oligoclase-10\% <br> Microcline-2\% | Chlorite |  | Calcite |

[^17]
## Section VI

> Sample 3. Sandstone (Orthoquartzite)

Megascopic description: The sandstone is brown and fine grained.

Microscopic description: Clastic, well sorted, angular and sub-angular grains.

| Stable <br> "primary" <br> minerals* | Stable <br> "secondary" <br> minerals* | Precipitates* <br> "cement" | Minor con- <br> Stituents |
| :--- | :--- | :--- | :--- |
| Oligoclase-5 $\%$ | Sericite |  | Calcite |

[^18]2. The Igneous Pebbles

Six thin sections of type samples of the igneous pebbles from the coarse conglomerates and the conglomeratic sandstones were examined.

## Pebble 1. Rhyolite

Megascopic description: Reddish brown, fine grained, and aphanitic. Quartz and feldspars are identified with a hand lens.

Microscopic description: Holocrystalline, hypautomor-phic--granular texture.

| Primary minerals |  | Secondary minerals |  |
| :--- | :--- | :--- | :--- |
| Quartz-35\% | Hornblende- $5 \%$ |  | Sericite |
| Oligoclase-15\% | Zircon |  | Kaolinite-10\% |
| Perthite-30\% | Magnetite |  |  |
|  | Epidote |  |  |

Pebble 2. Basalt

Megascopic description: Greenish gray, very fine grained, and aphanitic.

Microscopic description: Holocrystalline, diabasic texture.

Primary minerals Secondary minerals Alterations
Labradorite-60\% Magnetite-7\% Chlorite-3\%

Augite-30\% Epidote
Hornblende

Pebble 3. Amygdaloidal basalt

Megascopic description: Greenish gray, very fine grained, aphanitic, and amygdaloidal.

Microscopic description: Holocrystalline, diabasic texture. Most of the amygdules are filled with cryptocrystalline quartz. One amygdule is filled with skeleton hematite crystals.

| Primary minerals | Secondary minerals |  |
| :--- | :--- | :--- |
| Labradorite-60\% | Magnetite $-5 \%$ | Chlorite- $4 \%$ |
| Augite $-30 \%$ | Epidote |  |
|  | Biotite |  |
|  | Apatite |  |

## Pebble 4. Trachyte porphyry

Megascopic description: Light brown, very fine grained, and porphyritic. The phenocrysts are composed of feldspars.

Microscopic description: Hypocrystalline, microlitic texture, and trachitic structure. The phenocrysts are subhedral.

Primary minerals Secondary minerals Alterations
Phenocrysts:
Quartz-5\% Kaolinite Albite-5\% Magnetite-2\% Chlorite Orthoclase-15\% Biotite shreds-3\%

Microlites:
Orthoclase (?)-60\%
Glassy matrix-10\%

Pebble 5. Trachyte porphyry

Megascopic description: Light brown, very fine grained, and porphyritic. The phenocrysts are composed of feldspars.

Microscopic description: Hypocrystalline, microlitic texture, and trachitic structure. The phenocrysts are subhedral.

Primary minerals Secondary minerals Alterations
Phenocrysts: Quartz-5\% Chlorite
Albite-3\% Magnetite
Orthoclase-20\% Epidote
Microlites:
Orthoclase (?) - $60 \%$
Glassy matrix-10\%

Pebble 6. Rhyolite porphyry

Megascopic description: Reddish brown, very fine grained, and porphyritic. The phenocrysts are composed of feldspar and quartz. The pebble is fractured.

Microscopic description: Hypocrystalline texture. The phenocrysts are subhedral. The fractures in the pebble are filled with cryptocrystalline quartz.

## Pebble 6 (cont.)

Primary minerals Secondary minerals Alterations
Phenocrysts:Kaolinite
Quartz-5\%Perthite-8\%
Orthoclase-3\%
Sanidine-2\%
Glassy matrix:
Feldspars and
quartz-82\%
3. General Statement
A. The Sandstones
The sandstones along all of the sections show avery similar mineralogy and petrography. This is an in-dication that they have come from the same source orsources.
The fine grained sandstones are mostly orthoquartzites, the medium grained are arkosic sandstones, and the coarse grained are arkoses with more than 25 per cent of feldspars. As most of the outcrops are medium grained, it is inferred that the Porcupine Mountains Sandstone in the area studied, is mainly a conglomeratic arkosic sandstone with lenses of fine and coarse grained sandstone and arkose.
The cement of the orthoouartzites, arkosic sandstones, and arkoses is composed of secondary calcite and
hematite. The latter is the mineral from which the sandstones derive their characteristic reddish brown color.
B. The Igneous Pebbles

A total of one hundred and eighty-eight pebbles were collected from the different outcrops of coarse conglomerates and conglomeratic sandstone. These are classified into four main groups:

1. Trachyte porphyry: Seventy-eight pebbles of this group were collected.
2. Rhyolite porphyry: Forty-three pebbles of this group were collected.
3. Basalt: Thirty-four pebbles of this group were collected. Some are amygdaloidal.
4. Rhyolite: Sixteen pebbles of this group were collected.

The remaining seventeen pebbles are classified into two additional minor groups:
5. Quartz: Ten pebbles of this group were collected. One pebble showed typical agate zoning.
6. Granite: Seven pebbles of this group vere collected. One pebble showed gneissic banding.

The above described pebbles were found to have a roundness value of .6 to .9, and a sphericity of . 67 to . 87 (See appendix).

The average size of the pebbles is about three inches, but a few boulders of more than one foot in
diameter were observed. The pebbles are poorly sorted in their respective zones (Figures 1 and 5).

The majority of the pebbles composing the conglomerate zones are acidic and intermediate igneous rocks, which can be classified generally as felsites. This is in conformity with the character of most of the pebbles of the Keweenawan conglomerates that outcrop in the Keweenaw Peninsula and along the Black River southwest of the Porcupine lountains.

It is inferred that the felsite pebbles which make up most of the material of the conglomerate zones of the Porcupine Mountains sandstone in the area studied, were derived from the lower Keweenawan rocks. Gordon* be-

[^19]lieves that these rocks were the only possible source of felsites. Other pebbles found, such as granites, quartz, and agates were most probably derived from the pre-Keweenawan rocks, which include the Huronian and the Archeozoic.

## CORRELATIONS

Correlation of the outcrops from one section to another was attempted by the following methods:

1. By lithology: The outcrops in the area are composed of conglomeratic sandstones, sandstones, and thin conglomerate zones. The first two lithologic types do not occur in distinct zones, but grade into each other along both the strike and dip, making it impossible to trace any particular zone from one section to another. The only zones that possess distinct lithology and sharp contacts are those of the coarse conglomerates. Along Sections I, II, III, and IV, all of the conglomerate zones observed were very thin and irregular; but along Section $V$ and $V I$, relatively large conglomerate zones outcrop in the form of prominent ridges. Along Section V, these outcrop at Locations 14, 15, and 16; and along Section VI, similar ridges occur at Locations 21, 22, and 23 (Rap 2).

The topographic expression of the area of Sections $V$ and VI indicates that these conglomerate ridges described above are a part of one large conglomerate zone that extends from Sections $V$ to VI (Plate l). It is not known whether this zone extends to the east, or pinches out somewhere in the area between sections IV and $V$. The area west of Section VI was not traversed, and it could not be determined if the conglomerate zone extends to the west.

In an unpublished paper describing outcrops along the

Carp and Little Carp Rivers in the Porcupine Mountains, southwest of Section VI, Hornstein* reported the existence

```
* Hornstein, O.M., "A Field and Petrographic Study of Some Extrusive and Sedimentary Rocks Along the Carp and Little Carp Rivers in Ontonagon and Gogebic Counties, Michigan." Unpublished Master's thesis, Department of Geology and Geography, Michigan State College, 1950.
```

of conglomerate zones of very similar character stratigraphically above the Lake Shore Traps formation. It is very probable that one of these zones may be directly related to the conglomerate zone along Sections $V$ and $V I$. On a map of the area the regional strike of the Porcupine Mountains sandstone was projected from the position of the conglomerate zone of Section $V$ and $V I$, to the area of the Little Carp and Carp Rivers. At the locations where this projection cuts the two river courses, Hornstein*

## Ibid.

reported that outcrops of poorly sorted, coarse, felsite conglomerate occur, demonstrating the very probable relation existing between these zones.

It is not within the scope of this investigation to correlate between the area described in this paper, and the area of the Little Carp and Carp Rivers. Nevertheless, it has been shown that relations do exist, both on the basis of lithology and stratigraphic position. More detailed
field work in the area between the Iittle Carp and Carp Rivers and Section VI is, however, necessary before any definite correlations can be established.
2. By the sphericity and roundness of the sand grains and pebbles: Correlation by this method is not possible, because no distinct zones with a constant sphericity and roundness were noted. On the contrary, variations in sphericity and roundness figures were observed along the strike and dip of the sandstones. This is primarily due to variations in the grain size along the strike and dip. Correlation by this method would have been possible if distinct zones of constant grain sizes occurred. In this case, the sandstone grains would have a constant roundness. This assumption is based on results demonstrated by this investigation; namely, that equal-sized grains of the sandstones in the area studied, possess equal roundness.

Upon inspection of the roundness and sphericity figures of the pebbles from the different conglomerate zones, it can be seen that they possess a very similar sphericity and roundness, regardless of the zone from which they came (See appendix). Thus, these figures could not be used as a means of correlating conglomerate zones.

## CONCLUSION

The rocks that outcrop along the six sections traversed are composed mainly of conglomeratic sandstones that exhibit parallel bedding with local cross-bedding (Tap 2). Some zones of non-conglomeratic sandstone and coarse conglomerate also occur, probably as lenses. The rocks vary in color from brick red to dark brown, and range in grain size from fine to coarse. They are jointed, possibly faulted, and are cut by irregular veinlets of quartz and calcite.

Attempts to correlate the outcrops:

1. By lithology: Correlation was only possible between the conglomerate zones along Sections $V$ and $V I$ (Plate l).
2. By the sphericity and roundness of the sand grains and pebbles: Correlation was not possible.

The sandstones are composed essentially of quartz with minor amounts of oligoclase. Most of the pebbles are composed of intermediate and acidic igneous rocks with minor amounts of basic igneous pebbles. This indicates that the main source of the sediments was an area of outcrops of intermediate and acidic igneous rocks. The most likely area in the vicinity from which these sediments could have been derived is the main Red Rock area in the heart of the Porcupine llountains.

The basalt pebbles found in the conglomerate zones and the conglomeratic sandstones indicate that not all of the sediments were derived from acidic and intermediate igneous rocks. In the vicinity of the Porcupine Ilountains there are large
areas where basaltic flows are located*. These are probably

* Wright, F.E., and Lane, A.C., "Preliminary Geological Map of the Porcupine Mountains and Vicinity," Michigan Geological Survey Rept. for 1908, pl. 1 (1909).
the source rocks from which these basic igneous pebbles were derived.

The pebbles of the conglomerates and the congloneratic sandstones are well rounded indicating that they are waterworn. The writer is of the opinion that these sandstones and conglomerates are stream deposits for the following reasons:

1. Thin conglomerate zones are repeated haphazardly
through more than 1,500 feet of sandstone.
2. The sandstones are poorly sorted and feldspathic.
3. Ferruginous clay zones are founa in the sandstone.
4. Mud cracks are present.


Figure 3 . Location I7. In the fore-
ground is the brecciated sandstone zone
crisscrossed by calcite veinlets. In
the background is Lone Island.

Outcrop of conglomeratic



APPENDIX



Good



 Bedded, platy discontinuous
inches thick

## DESCRIPTION OF SANPIES OF SECPION I


ค.
10
$\infty$
$\bigcirc$

> !
$\not+$
Good

| Grain size |
| :---: |
| Medium |
| Medium |
| Fine |
| Medium |
| Very fine |
| Medium |
| Medium |
| Medium |

Color
$\qquad$
brown
Red-
dish
brown
Red$\stackrel{s}{n}$
r-1
rob
Red-
dish
brown
Red-
dish
brown
Red-
dish
brown
Brown


0
1
0
0
0
0
0
1
10
0
0
Sandstone

0
$\stackrel{1}{-1}$
0
0
0
0
0

$\boldsymbol{H} \quad \omega$
$\infty$
$\omega$
7 Sandstone
0
0
0
+3
0
0
0
0
0
0
Other char-
acteristics

1



| ro |  |
| :--- | :--- |
| 8 | 0 |
| 0 | 0 |

1
1
1
'8
8
8

| o |
| :--- |
| 8 |
| 8 |

Average Average
sphericity roundness

| $\begin{gathered} \underset{\sim}{N} \\ \underset{\sim}{\sim} \end{gathered}$ |  |  | $\underset{\substack{\text { ¢ }}}{\substack{\text { ¢ }}}$ | $\underset{\substack{\text { ¢ }}}{\text { ¢ }}$ | $\underset{\substack{0 \\ \underset{H}{4} \\ \hline \\ \hline}}{1}$ |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{aligned} & \text {. } \\ & \text { ๙゙ } \\ & \text { dy } \end{aligned}$ | $\begin{aligned} & \text { E. } \\ & \stackrel{\rightharpoonup}{-1} \\ & 0 \\ & 0 \\ & 0 \end{aligned}$ | 1 | $\stackrel{41}{8}$ |  | $\begin{aligned} & H \\ & \text { H } \\ & 0 \\ & \hline \end{aligned}$ |  | $\begin{aligned} & \mathbb{Q} \\ & \text { O } \\ & \text { O } \\ & 0 \\ & \hline 0 \end{aligned}$ | 1 |  |



| Sample number | Lithologic type |
| :---: | :---: |
| 9 | Sandstone |
| 10 | Conglomerate |
| Pebble | Basalt |
| 11 | Basalt |
| 12 | Basalt |
| 13 | Basalt |
| 14 | Sandstone |
| 15 | Conglomerate |
| 16 | Conglomeratic sandstone |

Other char-
acteristics
Vesicular
0
0
0
0
0
0
0
0
0
0
0
0
0
0 Pebbles up to
0.4 inches
Cross-bedded,
micaceous
 Average
sorting $\begin{array}{cc}\begin{array}{l}\text { Average } \\ \text { sphericity }\end{array} & \begin{array}{c}\text { Average } \\ \text { roundness }\end{array} \\ .75 & .8 \\ .83 & .7 \\ .83 & .4\end{array}$



## 

әtqqed
0
0
0
0
0
0
0

sandstone
Conglom-


Conglom-
eratic
sandstone
Conglom-
erate
Rhyolite
porphyry
H
Øo
On
®
0
$\sim$
0
0
0
0
0
0

| 0 |
| :---: |
| $\sim$ |
| 1 |
| 0 |
| 0 |
| 0 |
| 0 |
| 0 |
| 1 |

Other char-
acteristics
-
-
Quartz
phenocrysts
Bedded,
miceceous


| Average |
| :---: |
| roundness |

.9
.8
.8
Average
sphericity
.65
.81
.63
.75


Lithologic
type
Basalt
Rhyolite

Phyolite
porphyry
Conglom-
eratic
sandstone
Sample
number
Pebble
Pebble
Pebble
21

| DESCRIPTION OF SAMPLES OF SECTION II |  |  |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Sample number | Lithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| 1 | Sandstone | Brown | Fine | .77 | . 4 | Good | Bedded, cut by calcite veinlets |
| 2 | Conglomeratic sandstone | Brown | Fine | . 83 | . 4 | Good | Bedded, micaceous |
| 3 | Conglomeratic sandstone | Reddish brown | Fine | . 79 | . 4 | Good | Bedding indistinct, cut by calcite veins |
| 4 | Conglomeratic sandstone | Brown | Fine | . 83 | . 5 | Good | Leaching in irregular zones |
| 5 | Sandstone | Brown | Medium | . 79 | . 5 | Poor | Bedded |
| 6 | Sandstone | Reddish brown | Coarse | . 81 | . 6 | Poor | Bedded |
| 7 | Conglomeratic sandstone | Brown | Coarse | .79 | . 6 | Poor | Indistinct bedding |
| 8 | Conglomerate | - | - | - | - | Poor | Average sige of pebbles 3 inches |


| Sample number | Lithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pebble | Rhyolite | Reddish brown | Fine | . 65 | . 8 | - | - |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | .79 | . 7 | - | Quartz phenocrysts |
| Pebble | Basalt | Green- <br> ish <br> gray | Very fine | . 81 | . 9 | - | - |
| 9 | Sandstone | Reddish brown | Fine | . 83 | . 4 | Good | Bedded, micaceous |
| 10 | Sandstone | Dark brown | Medium | . 81 | . 5 | Good | Bedded, micaceous |
| 11 | Sandstone | Dark brown | Medium | .83 | . 5 | Poor | Massive |
| 12 | $\begin{aligned} & \text { Conglom- } \\ & \text { eratic } \\ & \text { sandstone } \end{aligned}$ | Dark brown | Coarse | .77 | .6 | Poor | Bedded |
| Pebble | Basalt | Green- <br> ish <br> gray | Very fine | .77 | .7 | - | - |
| 13 | Sandstone | Reddish brown | Medium | .79 | . 6 | Poor | Bedded |

Other char－
acteristics
Bedded

$$
\begin{aligned}
& \text { Massive, cut } \\
& \text { by calcite } \\
& \text { veinlets }
\end{aligned}
$$五

$$
\left.\begin{array}{ll}
0 & 8 \\
0 & 0 \\
0 & -1 \\
0 & -1 \\
4 & 6 \\
0 & 0 \\
B & 0 \\
4 & 0
\end{array} \right\rvert\,
$$

$$
\begin{array}{ll}
\text { Average } & \text { Average } \\
\text { sphericity } & \text { roundnes }
\end{array}
$$$\begin{array}{ll}\text { Average } & \text { Average } \\ \text { sphericity roundness }\end{array}$

Poor

$$
\begin{aligned}
& 0 \\
& 0 \\
& 0 \\
& 0
\end{aligned}
$$

$$
\begin{aligned}
& 0 \\
& 0 \\
& 0 \\
& 0 \\
& \hline
\end{aligned}
$$

$$
\begin{aligned}
& \mu_{1} \\
& 0 \\
& 0 \\
& \beta_{1}
\end{aligned}
$$H

O
O
H

$$
\begin{aligned}
& E_{0} \\
& 0 \\
& 0 \\
& 0_{1}
\end{aligned}
$$

\[
$$
\begin{aligned}
& 0 \\
& 0 \\
& 0 \\
& 0
\end{aligned}
$$

\]| o |
| :--- |
| 0 |
| 0 |
| 0 |

$$
\begin{gathered}
\kappa_{1} \\
\mathrm{O} \\
0 \\
0
\end{gathered}
$$

$$
\varrho
$$

?

$$
\underset{\sim}{H}
$$

$$
\underset{~+~}{+}
$$

!
-صـ$\stackrel{+}{\infty}$ミ
Brown.77

## Grain size

Grain size

$$
\begin{aligned}
& 0 \\
& 0 \\
& 8 \\
& \tilde{0} \\
& 0 \\
& 0
\end{aligned}
$$

$$
\begin{aligned}
& \text { 目 } \\
& \stackrel{1}{-1} \\
& \underset{\sim}{0} \\
& 0 \\
& 0
\end{aligned}
$$

Fine Bedded，platy
 zones $2-6$
inches thick

.73 .77
micaceous

$$
\begin{aligned}
& \text { Bedded, mica- } \\
& \text { ceous, exhi- } \\
& \text { bits mud } \\
& \text { cracks }
\end{aligned}
$$


Brown

$$
\begin{aligned}
& \text { Red- } \\
& \text { dish } \\
& \text { brown } \\
& \text { Brown }
\end{aligned}
$$


1
8
0
-1
－1
영
0
0 sandstoneConglom－
eraticsandstone
Sandstone
Medium
Sandstone0
0
0
+
0
0
0
－
a
00
0
0
+
0
0
0
0
a
0

$H$
$H$
$\stackrel{\sim}{r}$
0
-1
17
$\infty$
+
$\square$
$\square$
$\bigcirc$

| Sample number | Iithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 21 | Calcite vein in sandstone | White | - | - | - | - | Vein 6 inches thick, trends S. $16^{6}$ E., vertical |
| 22 | Sandstone | Brown | Medium | . 79 | . 5 | Poor | Bedded, micaceous |
| 23 | Sandstone | Brown | Medium | . 79 | . 5 | Poor | Bedded, micaceous |
| 24 | Sandstone | Brown | Fine | . 73 | . 4 | Good | Mud cracks, ripple marks |
| 25 | Sandstone | Brown | Fine | . 83 | . 5 | Good | Bedded |

$$
\begin{aligned}
& \text { Average Average } \\
& \text { sphericity roundness }
\end{aligned}
$$

Other char-
acteristics
Bedded,
micaceous
Pebbles up to
3 inches
$\quad-$
Bedded,
micaceous
Pebbles up to
4 inches
Average
sorting
Good
Poor
-
Pood
-
-

$$
\begin{array}{ll}
4 & 0
\end{array}
$$

$$
\varrho \quad \infty
$$

$$
\infty \quad \stackrel{\square}{0}
$$

$$
1 \quad \infty
$$

$$
\infty
$$

$$
.75
$$


$\infty$


.79

| $\begin{gathered} 0 \\ N \\ \sim \\ \sim-1 \\ 0 \\ \underset{\sim}{-1} \\ 0 \\ \substack{0-1 \\ -5} \end{gathered}$ |
| :---: |
|  |  |

## Medium

Coarse Medium
Very fine
Very fine
Fine



 dish brown Red-
dish
brown Brown


 | Sample | Lithologic |
| :--- | :--- |
| number | type |

0 0 0 + + 0 0 $\boxed{0}$ 0 0


sandstone
Rhyolite
porphyry

Conglom-
Basalt
Basalt
Rhyolite

| 0 |
| :--- |
| $\underset{\sim}{-1}$ |
| 0 |
| 0 |
| 0 |

$H$
0
0
$\sim$
0
0
0
0
0
0
$\infty$

| Sample number | Lithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | .77 | . 8 | - | Quartz phenocrysts |
| 5 | Conglomeratic sandstone | Dark brown | Coarse | . 75 | . 6 | Poor | Pebbles up to 4 inches |
| Pebble | Trachyte porphyry | Light brown | Very fine | .79 | . 7 | - | Vesicular, feldspar phenocrysts |
| Pebble | Rhyolite | Reddish brown | Fine | . 75 | . 8 | - | - |
| 6 | Conglomeratic sandstone | Dark brown | Coarse | . 75 | . 6 | Poor | Bedded, pebbles up to 1.0 inch |
| Pebble | Trachyte porphyry | Light brown | Very fine | .73 | . 8 | - | Feldspar phenocrysts |
| Pebble | Rhyolite | Reddish brown | Fine | . 73 | . 7 | - | - |
| 7 | Sandstone | Red- <br> dish <br> brown | Medium | . 83 | . 6 | Good | Bedded, micaceous |
| 8 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 3-4 in. |


| Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: |
| .73 | . 8 | - | Feldspar phenocrysts |
| .71 | . 8 | - | - |
| .73 | . 8 | - | - |
| .75 | . 8 | - | - |
| .73 | . 5 | Good | Bedded |
| .73 | . 6 | Poor | Bedded |
| . 75 | . 5 | Poor | Indistinct bedding |
| - | - | Poor | Pebbles up to 5 inches |
| .73 | . 9 | - | - |
| .73 | . 9 | - | Feldspar phenocrysts |


| $\begin{gathered} \mathbb{N} \\ \mathbf{N} \\ \cdot-1 \\ 0 \end{gathered}$ |  |  |  | $\begin{aligned} & \stackrel{0}{4} \\ & \stackrel{y}{+1-1} \\ & 4-1 \end{aligned}$ | E | $\pm$ | 日 |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| $\begin{aligned} & \text { dy } \\ & .{ }_{-1}^{0} \\ & \text { Hy } \end{aligned}$ | $\begin{aligned} & B \\ & 0 \\ & \stackrel{B}{8} \end{aligned}$ |  | $\begin{aligned} & \mathbb{1} \\ & \underset{\sim}{E} \\ & \stackrel{-1}{E} \end{aligned}$ | $\begin{aligned} & B \\ & 0 \\ & 0 \\ & 8 \end{aligned}$ | $\begin{aligned} & \overrightarrow{3} \\ & \stackrel{\rightharpoonup}{H} \\ & \stackrel{\sim}{0} \\ & \underset{y y y}{*} \end{aligned}$ | $\begin{aligned} & \text { N} \\ & \text { \& } \\ & 0 \\ & 0 \\ & 0 \end{aligned}$ |  | 1 | $\begin{aligned} & \text { B } \\ & 0 \\ & 0 \end{aligned}$ | $\begin{aligned} & B \\ & \underset{8}{D} \\ & \stackrel{y}{*} \end{aligned}$ |


Lithologic
type
Trachyte
porphyry
Rhyolite

Sandstone
Sandstone

Conglom-
erate
Basalt
Trachyte
porphyry

Pebble
Pebble
Pebble Basalt
Pebble
Pebbie
0
-1
0
0
0
0
0
0

| Sample <br> number | Iithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting: | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pebble | Trachyte porphyry | Light brown | Very fine | . 69 | . 9 | - | Feldspar phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | . 71 | . 9 | - | Feldspar phenocrysts |
| 13 | Sandstone | Brown | Coarse | . 69 | . 6 | Poor | Bedded |
| 14 | Conglomerate | Brown | - | - | - | Poor | Pebbles up to 5 inches |
| Pebble | Trachyte porphyry | Iight brown | Very fine | . 71 | . 8 | - | Vesicular, feldspar phenocrysts |
| Pebble | Rhyolite | Reddish brown | Fine | . 83 | . 8 | - | Vesicular |
| Pebble | Basalt | $\begin{aligned} & \text { Green- } \\ & \text { ish } \\ & \text { gray } \end{aligned}$ | Very fine | . 81 | : 8 | - | Vesicular |
| Pebble | Trachyte porphyry | Light brown | Very fine | .71 | . 9 | - | Feldspar phenocrysts |
| 15 | Conglomeratic sandstone | Brown | Coarse | . 73 | . 6 | Poor | Bedded, micaceous |
| 16 | Sandstone | Brown | Coarse | .73 | . 5 | Poor | Bedded, micaceous |


| Sample number | Lithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 17 | Conglomerate | Brown | - | - | - | Poor | Pebbies up to 6 inches |
| Pebble | Trachyte porphyry | Light brown | Very fine | .79 | .9 | - | Feldspar phenocrysts |
| Pebble | Rhyolite | Reddish brown | Very fine | .75 | . 9 | - | Vesicular |
| Pebble | Basalt | Green- <br> ish gr | Very fine | .69 | . 8 | - | $\cdots$ |
| Pebble | Trachyte porphyry | Light brown | Very fine | .83 | . 9 | - | Feldspar phenocrysts |
| 18 | $\begin{aligned} & \text { Conglom- } \\ & \text { eratic } \\ & \text { sandstone } \end{aligned}$ | Brown | Coarse | .69 | . 6 | Poor | Cross-bedded |
| 19 | Conglomeratic sandstone | Red- <br> dish <br> brown | Medium | .79 | . 5 | Good | Bedded, micaceous |
| Pebble | Trachyte porphyry | Light brown | - | . 83 | .7 | - | Feldspar phenocrysts |
| 20 | Sandstone | Reddish brown | Fine | .79 | . 5 | Good | Bedded, cut by quartz veins .5-1. 5 inches |
| 21 | Conglomerate | Brown | - | - | - | Poor | Cut by quartz veins up to 4 inches thick |

Other char-
0
0
0
+-1
0
0
$r-1$
$F$
0
0
0
0
0
0
Feldspar
phenocrysts
Vesicular
Vesicular
Bedded
4 inches,
vertical,
trends $N .49^{\circ}$ V.
Bedded
Bedded


Good
$\begin{array}{r}0 \\ 0 \\ 0 \\ \hline 8\end{array}$
0
0
0
8
Poor
Good
Average
sorting
Average Average
sphericity roundness
.79
.81
$\% \quad \%$
.6
$\begin{array}{ll}8_{1} & \\ 0 & \\ 0 & 1 \\ 0 & \end{array}$
1
1
I

0
$\bullet$
I
$\stackrel{\square}{\circ}$
1
.73
1
.81
.79
8
1
$\stackrel{\pi}{8} \quad 1$ Grain size


Very fine

| $\pm$ |
| :--- |
| -1 |
| .-1 |
| - |
| 1 |

Coarse
Coarse
Medium


Red-
dish
brown
Red-
dish

| 물 |
| :--- |
| 0 |
| $H$ |
| 0 |

Red-
brown
Brown
Red-
dish
brown


## Trachyte porphyry <br> 

Pebble
$\overline{\text { Iəquinu }}$
erdures
Iəqunu

| Sample number | Lithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 29 | Sandstone | Reddish brown | Fine | - | - | Good | Brecciated zone, cut by quartz veinlets |
| 30 | Sandstone | Reddish brown | Fine | - | - | Good | Massive, cut by quartz veinlets |
| 31 | $\begin{aligned} & \text { Conglom- } \\ & \text { eratic } \\ & \text { sandstone } \end{aligned}$ | Brown | Medium | . 81 | . 4 | Poor | Pebbles up to 3 feet |
| 32 | Sandstone | Reddish brown | Medium | .79 | . 5 | Good | Bedded, micaceous |
| 33 | Conglomeratic sandstone | Reddish brown | Medium | . 75 | . 5 | Poor | Bedded, pebbles up to 4 inches |
| 34 | Conglomerate | Brown | - | - | - | Poor | Pebbles up to 6 inches |
| Pebble | Trachyte porphyry | Light brown | Very fine | .77 | . 7 | - | Feldspar phenocrysts |
| Pebble | Basalt | ```Green- ish gray``` | Very fine | .79 | . 8 | - | - |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | .79 | . 7 | - | Quartz phenocrysts |

Other char-
acteristics
Pebbles up to
0.4 inches
Cross-bedded
Bedded
Average peb-
ble size
3-4 inches
Quartz
phenocrysts
Quartz
phenocrysts

1



- !

I
$\stackrel{\square}{0}$
$\stackrel{\rightharpoonup}{ }$
$\sigma$
$\stackrel{\square}{ }$

.77

$$
\begin{array}{r}
.73 \\
.79
\end{array}
$$

I
.87
.73
$\stackrel{H}{-1}$
$\stackrel{\infty}{\infty}$

1


Other char-
acteristics
Bedded
Bedded


| Average <br> Sphericity | Average <br> roundness |
| :---: | :---: |
| .81 | .5 |
| .77 | .5 |


|  |  |
| :---: | :---: |




|  |  |
| :---: | :---: |
|  | 8 |

DESCRIPTION OF SAMPIES OF SECTION IV
 $\begin{array}{ll}0 & 00 \\ 00 & 2 \\ 0 & 0 \\ 0 & 0-1 \\ 4 & + \\ 0 & \varepsilon_{1} \\ 8 & 0 \\ 4 & 0\end{array}$


H
O
0
O


1
1
$\begin{array}{lll} & 0 & 8 \\ & 0 & 1 \\ 0 & 0 \\ 0 & 0 \\ 0 & 1\end{array}$

$\because \quad!\quad!$



| $\begin{aligned} & 8 \\ & 0 \\ & 0 \\ & 0 \\ & 0 \\ & 0 \end{aligned}$ | $\begin{aligned} & \text { d } \\ & \text { B } \\ & \text { R } \\ & \text { N } \end{aligned}$ | $$ | $\begin{aligned} & \text { g } \\ & \mathbf{B} \\ & 0 \\ & \text { H } \\ & \text { M } \end{aligned}$ | $\begin{aligned} & \text { F } \\ & \text { 冡 } \\ & 0 \\ & \text { 品 } \end{aligned}$ | $\begin{aligned} & \text { H } \\ & 0 \\ & 0 \\ & 0 \\ & 0 \end{aligned}$ | $\begin{aligned} & + \\ & \stackrel{\rightharpoonup}{60} \\ & \rightarrow \\ & H \end{aligned}$ |  |  | $\begin{aligned} & \text { C } \\ & \text { B } \\ & 0 \\ & \text { \& } \\ & \text { M } \end{aligned}$ |  |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |


0
0
0
0
0
0
0
0
0
0
$\alpha$
Sandstone
0
H
0
+
0
0
0
®
0
0
Sandstone
Conglom－
erate
Granite
Rhyolite
porphyry
Trachyte
porphyry
0
I
0
+
0
0
－
－
O

$\begin{array}{ll}0 & 0 \\ -1 & \sim \\ 0 & 0 \\ 0 & 0 \\ 0 & 0 \\ 0 & 0\end{array}$
Pebble
$\rightarrow$ N N M H $N$
$\omega$ 人

| Sample number | Lithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 8 | Conglomeratic sandstone | Brown | Coarse | .71 | .6 | Poor | Bedded, pebbles up to 6 inches |
| Pebble | Trachyte porphyry | Light brown | Very fine | .73 | .7 | - | Feldspar phenocrysts |
| 9 | $\begin{aligned} & \text { Conglom- } \\ & \text { eratic } \\ & \text { sandstone } \end{aligned}$ | Reddish brown | Medium | . 81 | . 5 | Good | Bedded |
| 10 | Conglom- <br> eratic sandstone | Brown | Coarse | .75 | . 6 | Poor | Bedded |
| Pebble | Trachyte porphyry | Light brown | Very fine | .77 | . 9 | - | Feldspar phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | . 85 | . 9 | - | Feldspar phenocrysts |
| 11 | Conglomeratic sandstone | Brown | Coarse | .75 | .6 | Poor | Massive |
| 12 | Conglomeratic sandstone | Brown | Coarse | .73 | . 6 | Poor | Cross-bedded |
| Pebble | Trachyte porphyry | Iight brown | Very fine | .73 | .8 | - | Feldspar phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | . 85 | . 9 | - | Feldspar phenocrysts |

Other char-
acteristics
Feldspar
phenocrysts
Bedded
Bedded
Average
Sorting 1 8
0
0
0
0 ..... $5_{1}$
0
0
0
0
Average $\quad$ Average
Sphericity roundness $\stackrel{\bullet}{0}$

$$
\varrho
$$

$$
\begin{aligned}
& .83 \\
& .73
\end{aligned}
$$

$$
.75
$$

Grain size
Very fine
Coarse
Coarse
Color
Light
brown
Brown
Brown

| Lithologic <br> type |
| :--- |
| Trachyte |
| porphyry |
| Conglom- |
| eratic |
| sandstone |
| Conglom- |
| eratic |
| sandstone |


| Sample |
| :--- |
| number |

Pebble
13
14
DESGRIPTION OF SAMPLES OF SECTION $V$

$$
\begin{aligned}
& \text { Average } \\
& \text { sorting }
\end{aligned}
$$

Grain size
Coarse
Medium


 Conglom-
eratic
sandstone
sandstone
Sandstone
Conglom-
eratic
sandstone
Trachyte
porphyry
Trachyte
porphyry
Conglom-
eratic
sandstone

әTqqә्ट
Pebble
0
$\sim$
0
0
0
0
0
0
Sample
number
-
UMOIq
74371
Brown
Green-
ish
gray
1

$$
\begin{aligned}
& \text { Average Average } \\
& \text { sphericity roundness }
\end{aligned}
$$

Poor
Poor
Poor
Poor
Other char-
acteristics
Cross-bedded,
micaceous
Pebbles up to
0.7 inches
Bedded,
micaceous
Micaceous

Vesicular,
feldspar
phenocrysts
Feldspar
phenocrysts
Pebbles up to
0.7 inches

1




1
1
$\infty$.

ค ค ค ค ค ล ค
.85
.73
.75
.77
-7

$$
\begin{aligned}
& .5 \\
& .5
\end{aligned}
$$

$\sim$
$\infty$

| Sample number | Lithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | .77 | . 8 | - | Quartz phenocrysts |
| 6 | Sandstone | Brown | Coarse | .77 | . 6 | Poor | Cross-bedded |
| 7 | Conglomerate | Brown | - | - | - | Poor | Pebbles up to 5 inches |
| Pebble | Granite | Light | Medium | - | - | - | - |
| Pebble | Trachyte porphyry | Light <br> brown | Very fine | .87 | .9 | - | Feldspar <br> phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | . 75 | . 7 | - | Feldspar phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | .71 | .7 | - | Feldspar phenocrysts |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | . 75 | . 8 | - | Feldspar and quartz phenocrysts |
| 8 | Sandstone | Brown | Coarse | .77 | . 6 | Poor | Bedded |
| 9 | Conglomerate | Brown | - | - | - | Poor | Pebbles up to 4 inches |
| Pebble | Basalt | Green- <br> ish <br> gray | Very fine | - | - | - | Epidote amygdules |


| Sample number | Lithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pebble | Trachyte porphyry | Light brown | Very fine | .77 | . 8 | - | Feldspar phenocrysts |
| Pebble | Granite | Light | Medium | - | - | - | - |
| Pebble | Basalt | Green- <br> ish <br> gray | Very fine | .81 | . 8 | - | Epidote amygdules |
| 10 | Sandstone | Brown | Coarse | . 75 | . 6 | Poor | Bedded |
| 11 | Conglomeratic sandstone | Brown | Coarse | .73 | . 6 | Poor | Cross-bedded |
| 12 | Conglornerate | Brovn | - | - | - | Poor | Pebbles up to 4 inches |
| Pebble | Trachyte porphyry | Light brown | Very fine | . 73 | . 8 | - | Vesicular, feldspar phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | .71 | . 8 | - | Feldspar phenocrysts |
| Pebble | Granite | Light | Medium | - | - | - | - |
| Pebble | Prachyte porphyry | Light brown | Very fine | . 75 | . 6 | - | Feldspar phenocrysts |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | - | - | - | Quartz phenocrysts |

Other char-
acteristics Bedded,
micaceous
Indistinct
bedding
Feldspar
phenocrysts
Pebbles up to
2 inches Feldspar
phenocrysts
Vesicular,
feldspar
phenocrysts
Bedded 0
80
0
0
0
0
$B$
4

Poor

号

1
$\begin{array}{ll}\text { Average } & \text { Average } \\ \text { sphericity roundness }\end{array}$

| Sample number | Lithologic type | Color | Grain size |
| :---: | :---: | :---: | :---: |
| 13 | $\begin{aligned} & \text { Conglom- } \\ & \text { eratic } \\ & \text { sandstone } \end{aligned}$ | Brown | Medium |
| 14 | Sandstone | Brown | Fine |
| 15 | $\begin{aligned} & \text { Conglom- } \\ & \text { eratic } \\ & \text { sandstone } \end{aligned}$ | Brown | Medium |
| Pebble | Trachyte porphyry | Light brown | Very fine |
| 17 | $\begin{aligned} & \text { Conglom- } \\ & \text { eratic } \\ & \text { sandstone } \end{aligned}$ | Brown | Coarse |
| Pebble | Basalt | $\begin{aligned} & \text { Green- } \\ & \text { ish } \\ & \text { gray } \end{aligned}$ | Very fine |
| Pebble | Trachyte porphyry | Light brown | Very fine |
| Pebble | Trachyte porphyry | Iight brown | Very fine |
| 18 | $\begin{aligned} & \text { Conglom- } \\ & \text { eratic } \\ & \text { sandstone } \end{aligned}$ | Brown | Coarse |


| Sample number | Lithologic type | Color | Grain size | Average sphericity | A verage roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pebble | Rhyolite porphyry | Red- <br> dish <br> brown | Very fine | . 83 | . 8 | - | Quartz phenocrysts |
| Pebble | Rhyolite porphyry | Red- <br> dish <br> brown | Very fine | .73 | . 8 | - | Quartz and feldspar phenocrysts |
| Pebble | Rhyolite | Reddish brown | Fine | .71 | . 6 | - | - |
| 19 | Conglomerate | Brown | - | - | - | Poor | Pebbles up to 5 inches |
| Pebble | Trachyte porphyry | Light brown | Very fine | . 79 | . 7 | - | Feldspar phenocrysts |
| Pebble | Trachyte porphyry | Light <br> brown | Very fine | .77 | . 8 | - | Feldspar phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | . 83 | . 8 | - | Feldspar phenocrysts |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | . 81 | . 8 | - | Quartz and feldspar phenocrysts |
| Pebble | Rhyolite | Reddish brown | Fine | .75 | .7 | - | - |
| Pebble | Trachyte porphyry | Iight <br> brown | Very fine | . 75 | . 8 | - | Feldspar phenocrysts |


| Sample number | Lithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 20 | Conglomerate | Brown | - | - | - | Poor | Pebbles up to 5 inches |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | .79 | . 6 | - | Quartz and feldspar phenocrysts |
| Pebble | Granite | Ijght | Medium | - | - | - | - |
| Pebble | Trachyte porphyry | Light brown | Very fine | .71 | . 9 | - | Feldspar phenocrysts |
| Pebble | Quartz | Light | Fine | - | - | - | Agate zoning: |
| 21 | Conglom- | Brown | - | - | - | Poor | Pebbles up to 5 inches |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | .73 | .7 | - | Quartz and feldspar phenocrysts |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | .79 | . 8 | - | Quartz and feldspar phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | $\cdots$ | - | - | Feldspar phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | - | - | - | Feldspar phenocrysts |
| 22 | Conglomerate | Brown | - | - | - | Poor | Pebbles up to 4 inches |

Other char-
acteristics



1 Quartz and
feldspar
phenocrysts

$$
\begin{aligned}
& \text { Quartz and } \\
& \text { feldspar } \\
& \text { phenocrysts }
\end{aligned}
$$


Poor
Poor

$$
1
$$

I
1





5 Grain size

Very fine
Very fine
Very fine





Lithologic
type

Pebble $\begin{aligned} & \text { Trachyte } \\ & \text { porphyry }\end{aligned}$
Pebble Rhyolite
Sandstone
Conglom-
erate
Rhyolite
porphyry
Rhyolite
porphyry
0
0
0
0
0
0
0

Pebble $\begin{aligned} & \text { Rhyolite } \\ & \text { porphyry }\end{aligned}$
Pebble Basalt
Sample
number
Pebble

| Sample number | Lithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 25 | Conglomeratic sandstone | Brown | Medium | . 75 | . 5 | Poor | Bedded, micaceous |
| 26 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 4 inches |
| Pebble | Rhyolite | Reddish brown | Fine | - | - | - | Vesicular |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | . 73 | . 8 | - | Quartz and feldspar phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | . 85 | . 9 | - | Feldspar phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | . 81 | . 9 | - | Feldspar phenocrysts |
| Pebble | Basalt | ```Green- ish gray``` | Very fine | .77 | . 8 | - | - |
| 27 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 4 inches |
| Pebble | Granite | Light | Medium | . 65 | . 8 | - | - |
| Pebble | Trachyte porphyry | Light brown | Very fine | - | - | - | Feldspar phenocrysts |


| Sample number | Iithologic type | Golor | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pebble | Trachyte porphyry | Light brown | Very fine | . 81 | . 9 | - | Feldspar phenocrysts |
| 28 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 4 inches |
| Pebble | Trachyte | Light brown | Very fine | - | - | - | Vesicular |
| Pebble | Granite | Light | Medium | . 76 | . 8 | - | - |
| Pebble | Basalt | ```Green- ish gray``` | Very fine | - | - | - | - |
| Pebble | Quartz | Light gray | Fine | - | - | - | Quartzite (?) |
| 29 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 4 inches |
| Pebble | Trachyte porphyry | Light brown | Very fine | .74 | . 8 | - | Feldspar phenocrysts |
| Pebble | Basalt | Greenish gray | Very fine | .73 | . 7 | - | - |
| Pebble | Trachyte porphyry | Light brown | Very fine | - | - | - | Feld.spar phenocrysts |
| 30 | Conglomerate | Brown | - | - | - | Poor | Pebbles up to 1.0 feet |


| Sample number | Lithologic type | Color | Grain size | Average sphericity | A verage roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pebble | Trachyte porphyry | Iight <br> brown | Very fine | .74 | . 8 | - | Feldspar phenocrysts |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | . 81 | . 6 | - | Quartz and feldspar phenocrysts |
| Pebble | Quartz | White | Fine | - | - | - | Vein quartz |
| Pebble | Basalt | Green- <br> ish <br> gray | Very fine | .73 | . 7 | - | - |
| 31 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 5 inches |
| Pebble | Trachyte porphyry | Light brown | Very fine | .77 | . 8 | - | Feldspar phenocrysts |
| Pebble | Basalt | Green- <br> ish <br> gray | Very fine | .79 | . 8 | - | - |
| Pebble | Trachyte porphyry | Iight brown | Very fine | .73 | .7 | $\cdots$ | Feldspar phenocrysts |
| 32 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 5 inches |
| Pebble | Basalt | Green- <br> ish <br> gray | Very fine | . 81 | . 8 | - | - |


| Sample number | Iithologic type | Color | Grain size | Average sphericity | A verage roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pebole | Trachyte porphyry | Light brown | Very fine | .77 | . 6 | - | Feldspar phenocrysts |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | .73 | .7 | - | Quartz and feldspar phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | .79 | . 8 | - | Feldspar phenocrysts |
| Pebble | Trachyte | Light brown | Very fine | . 77 | . 9 | - | Vesicular |
| 33 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 5 inches |
| Pebble | Trachyte porphyry | Iight brown | Very fine | . 81 | . 8 | - | Feldspar phenocrysts |
| Pebble | Basalt | Green- <br> ish <br> gray | Very fine | . 73 | . 9 | - | - |
| Pebble | Trachyte porphyry | Light brown | Very fine | - | - | - | Feldspar phenocrysts |
| Pebble | Quartz | White | Fine | - | - | - | Vein quartz(?) |
| Pebble | Trachyte porphyry | Iight brown | Very fine | . 73 | . 8 | - | Feldspar phenocrysts |


| Sample number | Lithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| 34 | $\begin{aligned} & \text { Conglom- } \\ & \text { eratic } \\ & \text { sandstone } \end{aligned}$ | Reddish brown | Coarse | .73 | . 6 | Poor | Bedded |
| 35 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 3 inches |
| Pebble | Trachyte porphyry | Light brown | Very fine | .77 | . 8 | - | Feldspar phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | .73 | . 7 | - | Feldspar phenocrysts |
| Pebble | Basalt | Green- <br> ish <br> gray | Very fine | .71 | .7 | - | - |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | - | - | - | Quartz phenocrysts |
| 36 | Sandstone | Brown | Coarse | .77 | .6 | Poor | Bedded |
| 37 | $\begin{aligned} & \text { Conglom- } \\ & \text { erate } \end{aligned}$ | Brown | - | - | - | Poor | Average size of pebbles 3 inches |
| Pebble | Quartz | Light green | Fine | - | - | - | Vein quartz(?) |
| Pebble | Trachyte porphyry | Light brown | Very fine | .71 | . 8 | - | Feldspar phenocrysts |


| Sample number | Lithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | . 65 | . 8 | - | Quartz phenocrysts |
| 38 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 3-4 inches |
| Pebble | Trachyte porphyry | Light <br> brown | Very fine | .81 | . 9 | - | Feldspar phenocrysts |
| Pebble | Rhyolite porphyry | Red- <br> dish <br> brown | Very fine | .73 | .7 | - | Quertz phenocrysts |
| Pebble | Trachyte porphyry | Iight <br> brown | Very fine | .75 | . 6 | - | Feldspar phenocrysts |
| 39 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 3-4 inches |
| Pebble | Basalt | Green- <br> ish <br> gray | Very fine | .73 | .9 | - | Epidote amygdules |
| Pebble | Rhyolite porphyry | Red- <br> dish <br> brown | Very fine | - | - | - | Quartz phenocrysts |
| Pebble | Trachyte porphyry | Light brown | Very fine | .73 | .9 | - | Feldspar phenocrysts |

Other char-
acteristics
Average size
of pebbles
$3-4$ inches
Feldspar
and quartz
phenocrysts
Feldspar
and quartz
phenocrysts
Banded dark
minerals
Parallel
flow lines

1
111

| Average <br> sphericity | Average <br> roundness |
| :---: | :---: |
| - | - |
| .73 | .8 |
| .71 | .8 |
| - | - |
| - | - |

Grain size
Very fine
Very fine
Medium
Very fine

| $\begin{aligned} & 8 \\ & 0 \\ & 0 \\ & 0 \\ & 0 \\ & 0 \end{aligned}$ |  |  |  |  |  |
| :---: | :---: | :---: | :---: | :---: | :---: |


|  |  |
| :---: | :---: |
|  |  |
|  |  |
|  |  |
|  |  |


| $\begin{aligned} & 1 \\ & 0 \\ & 0 \\ & 0 \\ & -10 \\ & 0 \\ & 0 \\ & 0 \\ & 0 \\ & 0 \\ & 0 \\ & 0 \\ & \hline \end{aligned}$ |
| :---: |
|  |  |
|  |  |


Rhyolite
porphyry

Sample
number

| 0 |
| :--- |
| $\sim$ |
| 0 |
| 0 |
| 0 |
| 0 |
| $\sim$ |

Pebble
$\begin{array}{cc}0 & 0 \\ 0 & 0 \\ 0 & 0 \\ 0 & 0 \\ 0 & 0 \\ 0 & 0 \\ 0 & 0\end{array}$
Other char-
acteristics
Brecciated, cut
 lets. Massive Bedded
Feldspar
phenocry
$o$
on
on
0
0
0
0
0
0
0
0
0
0
0
Feldspar
phenocrysts
Quartz
phenocrysts

Average
Sorting

| r |
| :--- |
| 8 |
| 8 |

1
Bedded DESCRIPTION OF SAMPLES OF SEGTION VI
$\underline{\underline{G r a i n ~ s i z e}}$

| $c$ |
| :---: |
| 0 |
| 0 |
| -1 |
| 0 |
| 0 |$|$

I
$\bullet$
․
1
$\therefore \quad \bullet$
? $\because$
Grain size Average Average
sphericity roundness
1

$\stackrel{-}{-}$
$\stackrel{H}{\leftrightarrows} \stackrel{\sim}{\approx}$
$\stackrel{20}{\sim}$
1
Fine
Coarse
Very fine
Very fine



Light
brown


Lithologic
type
əuozspues
Conglomeratic
auozspues
Trachyte
porphyry
+
H
区
on
®
Trachyte
porphyry
Rhyolite
porphyry
Sandstone
Trachyte
porphyry

$\begin{array}{cc}0 & 0 \\ \sim & -1 \\ -1 & -1 \\ 0 & 0 \\ 0 & 0 \\ 0 & 0 \\ & \end{array}$


| $\stackrel{0}{1}$ |
| :--- |
| 0 |
| 0 |
| 0 |


| Sample number | Iithologic type | Color | Grain size | Average sphericity | A verage roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | .77 | . 9 | - | Quartz phenocrysts |
| 4 | Conglomeratic sandstone | Brown | Coarse | . 73 | . 6 | Poor | Cross-bedded |
| Pebble | Trachyte porphyry | Hight brown | Very fine | .79 | .7 | - | Feldspar phenocrysts |
| Pebole | Basalt | Green- <br> ish <br> gray | Very fine | .77 | .7 | - | - |
| 5 | $\begin{aligned} & \text { Conglom- } \\ & \text { eratic } \\ & \text { sandstone } \end{aligned}$ | Brown | Medium | .73 | . 5 | Poor | Cross-bedded, micaceous |
| Pebble | Rhyolite porphyry | Red- <br> dish <br> brown | Very fine | .77 | .7 | - | Quartz phenocrysts |
| Pebble | Trachyte porphyry | $\begin{aligned} & \text { Red- } \\ & \text { dish } \\ & \text { brown } \end{aligned}$ | Very fine | . 81 | . 8 | - | Feldspar phenocrysts |
| 6 | Conglomeratic sandstone | Brown | Coarse | .71 | . 6 | Poor | Indistinct bedding |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | .71 | . 8 | - | Quartz phenocrysts |


| Sample number | Lithologic type | Color | Grain size | Average sphericity | Average roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | .79 | . 8 | - | Quartz and feldspar phenocrysts |
| 7 | Conglomeratic sandstone | Brown | Medium | . 77 | . 5 | Poor | Bedded |
| 8 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles $2-3$ in. |
| Pebble | Basalt | $\begin{aligned} & \text { Green- } \\ & \text { ish } \\ & \text { gray } \end{aligned}$ | Very fine | - | - | - | - |
| Pebble | Trachyte porphyry | Light brown | Very fine | . 87 | . 8 | - | Feldspar phenocrysts |
| 9 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 2 inches |
| Pebble | Quartz | Milky white | Very fine | .79 | . 7 | - | Vein quartz(?) |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | . 77 | . 7 | - | Feldspar and quartz phenocrysts |
| Pebble | Basalt | $\begin{aligned} & \text { Green- } \\ & \text { ish } \\ & \text { gray } \end{aligned}$ | Very fine | .73 | .7 | - | - |
| 10 | Conglom- erate | Brown | - | - | - | Poor | Average size of pebbles 3-4 in. |


| Sample number | Lithologic type | Color | Grain size | Average sphericity | A verage roundness | Average sorting | Other characteristics |
| :---: | :---: | :---: | :---: | :---: | :---: | :---: | :---: |
| Pebble | Trachyte porphyry | Light <br> brown | Very fine | .67 | .7 | - | Feldspar phenocrysts |
| Pebble | Basalt | Green- <br> ish <br> gray | Very fine | .71 | . 8 | - | - |
| Pebble | Rhyolite porphyry | Red- <br> dish brown | Very fine | . 73 | .7 | - | Feldspar and quartz phenocrysts |
| 11 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 3-4 in. |
| Pebble | Quartz | Light gray | Very fine | .71 | . 6 | - | Vein quartz(?) |
| Pebble | Trachyte porphyry | Light brown | Very fine | .87 | .7 | - | Feldspar phenocrysts |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine | . 83 | . 8 | - | Quartz phenocrysts. Flow lines |
| 12 | Conglomerate | Brown | - | - | - | Poor | Average size of pebbles 3-4 in. |
| Pebble | Trachyte porphyry | Light brown | Very fine | .71 | .7 | - | Feldspar phenocrysts |
| Pebble | Rhyolite porphyry | Red- <br> dish <br> brown | Very fine | .77 | . 9 | $\cdots$ | Feldspar and quartz phenocrysts |



| Sample number | Iithologic type | Color | Grain size |
| :---: | :---: | :---: | :---: |
| Pebble | Trachyte porphyry | Light brown | Very fine |
| 13 | Conglomerate | Brown | - |
| Pebble | Trachyte porphyry | Light brown | Very fine |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine |
| Pebble | Quartz | Light gray | Very fine |
| Pebble | Basalt | ```Green- ish gray``` | Very fine |
| 14 | Conglomerate | Brown | - |
| Pebble | Trachyte porphyry | Light brown | Very fine |
| Pebble | Trachyte porphyry | Light brown | Very fine |
| Pebble | Rhyolite porphyry | Reddish brown | Very fine |

## BIBLIOGRAPHY

## Books

Billings, M.P. Structural Geology. New York: Prentice-Hall, Inc., 1947.

Pettijohn, F.J. Sedimentary Rocks. New York: McGraw-Hill Book Co., 19 90.

Wahlstrom, E.E. Igneous Minerals and Rocks. New York: John Wiley \& Sons, Inc., 1948.

## Reports

Butler, B.S., Burbank, W.S., and Collaborators. "The Copper Deposits of Michigan." United States Geological Survey Prof. Paper 144 (1929).

Gordon, W.C., and Lane, A.C. "A Geological Section From Bessemer Down Black River." Michigan Geological Survey Rept. for 1906, pp. 397-507, map (1907).

Irving, Roland D. "The Copper-Bearing Rocks of Lake Superior." United States Geological Survey Monograph 5, 464 pages, maps (1883).

Lane, A.C. "The Keweenaw Series of Lichigan." Michigan Geological Survey, Publication 6 (Geol. Series 4), 2 vol., 983 pages, maps (1911).

Leith, C.K., Lund, R.J., and Leith, A. "Pre-Cambrian Rocks of the Lake Superior Region." United States Geological Survey Prof. Paper 184, (1935).

Mertie, J.B. "Graphic and Mechanical Computation of Thickness of Strata and Distance to a Stratum." United States Geological Survey Prof. Paper 129-C (1921).
Ruthven, A.G. "An Ecological Survey in the Porcupine Mountains and Isle Royale." Michigan Geological Survey Rept. for 1905 (1906).

Van Hise, C.R., and Leith, C.K. "The Geology of the Lake Superior Region." United States Geological Survey Monograph 52, 641 pages, maps (1911).

Wright, F.E. "Report of Progress in the Porcupines." Michigan Geological Survey Rept. for 1903, pp. 33-44 (1905).



MAP OF THE NORTHERN PART OF THE PORCUPINE MOUNTAINS, ONTONAGON COUNTY, MICHIGAN

Wright, F.F., and Lane, A.C. "Preliminary Geological Map of the Porcupine Mountains and Vicinity." Michigan Geological Survey Rept. for 1908, plate 1 (1909).

## Articles

Krumbein, W.C. "Measurement and Geological Significance of Shape and Roundness of Sedimentary Particles." Journal of Sedimentary Petrology, Vol. 11 (1941), pp. 68-69, plate l, abstract.

Rittenhouse, G. "A Visual Irethod of Estimating IwoDimensional Sphericity." Journal of Sedimentary Petrology, Vol. 13 (1943), pp. 79-81, abstract.

## Unpublished Material

Hornstein, O.M. "A Field and Petrographic Study of Some Extrusive and Sedimentary Rocks Along the Carp and Iittle Carp Rivers in Ontonagon." Master's thesis, Department of Geology, Hichigan State College (1950).

Kelly, W.A. Personal communication.
Thaden, R.E. Personal cormunication.


[^0]:    * Lane, A.C., "The Keweenawan Series of Michigan," Michigan Geological Survey, Publication 6 (Geol. Series 4), 2 Vol., $983 \mathrm{pp} .(1911)$.

[^1]:    * Irving, R.D., "The Copper Bearing Rocks of Lake Superior," United States Geological Survey, Monograph 5 (1883).

[^2]:    * Lane, A.C., "The Keweenawan Series of Michigan," Michigan Geological Survey, Publication 6 (Geol. Series 4), 2 Vol., 983 pp. (1911).

[^3]:    *Butler, B.S., Burbank, W.S., and Collaborators, "The Copper Deposits of Michigan," United States Geological Survey, Prof. Paper 144 (1929).

[^4]:    * Irving, R.D., "The Copper Bearing Rocks of Lake Superior," United States Geological Survey, Monograph 5, p. 139 (1883).

[^5]:    * Butler, B.S., Burbank, W.S., and Collaborators, "The Copper Deposits of Michigan," United States Geological Survey, Prof. Paper 144 (1929).

[^6]:    * Butler, B.S., Burbank, W.S., and Collaborators, "The Copper Deposits of Michigan," United. States Geological Survey, Prof. Paper 144 (1929).

[^7]:    * Wright, F.E., and Lane, A.C., Preliminary Geological Map of the Porcupine Mountains and Vicinity, Nichigan Geological Survey Rept. for 1908, pl. 1 (1909).

[^8]:    * Wright, F.E., and Lane, A.C., Preliminary Geological Map of the Porcupine Mountains and Vicinity, Michigan Geological Survey Rept. for 1908, pl. 1 (1909).

[^9]:    * Irving, R.D., "The Copper Bearing Rocks of Lake Superior," United States Geological Survey, Monograph 5 (1883).

[^10]:    * Irving, R.D., "The Copper Bearing Rocks of Lake Superior," United States Geological Survey, Monograph 5 (1883).

[^11]:    * Irving, R.D., "The Copper Bearing Rocks of Lake Superior," United States Geological Survey, Monograph 5 (1883).

[^12]:    * Kelly, W.A., Personal communication

[^13]:    * Rittenhouse, G., "A Visual Method of Estimating Two-dimensional Sphericity," Journal of Sedimentary Petrology, Vol. 13, pp. 79-81, abstract (1943).

[^14]:    * Pettijohn, F.J., Sedimentary Rocks, p. 86 (1940).

[^15]:    * Pettijohn, F.J., Sedimentary Rocks, p. 86 (1940).

[^16]:    * Pettijohn, F.J., Sedimentary Rocks, p. 86 (1940).

[^17]:    * Pettijohn, F.J., Sedimentary Rocks, p. 86 (1940).

[^18]:    * Pettijohn, F.J., Sedimentary Rocks, p. 86 (1940).

[^19]:    Gordon, W.C., and Lane, A.C., "A Geological Section From Bessemer Down Black River," Michigan Geological Survey Rept. for 1906, p. 429 (1907).

